



# Assessing and Improving the Numerical Solution of Atmospheric Physics in an Earth System Model

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on behalf of a SciDAC project team:

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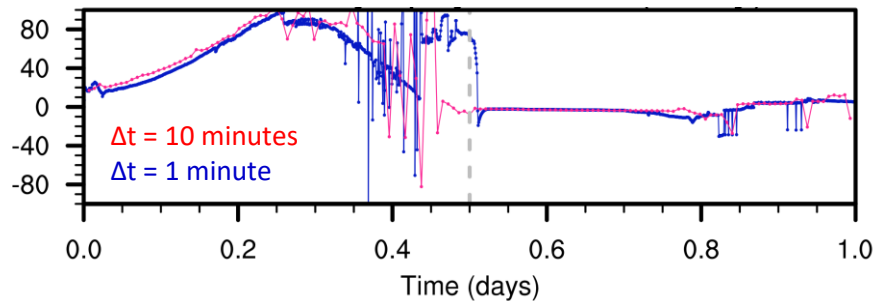
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# Parameterizations in AGCMs

- ▶ Traditional focus on conceptualization of understanding + handling of spatial scales
- ▶ Practical motivation to use longest possible time step
- ▶ Ubiquitous use of clipping, limiters etc.

## 1-day time series of T tendency (K/day) at 700 hPa from E3SM v0 physics



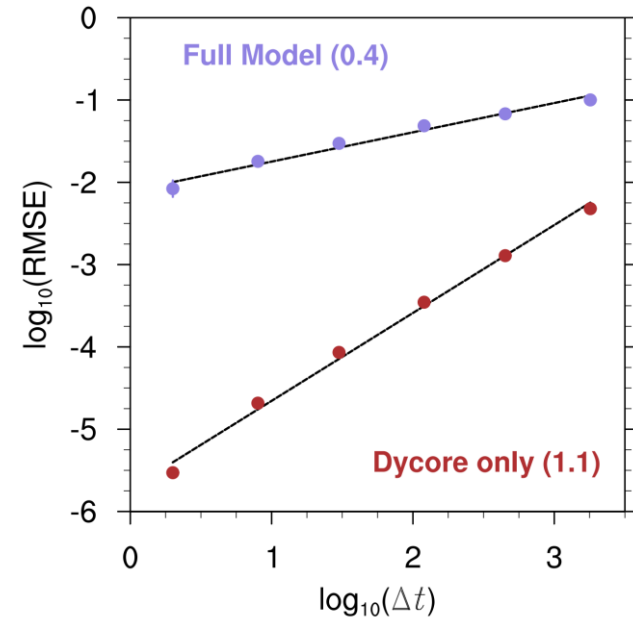
- ▶ Physics parameterizations
  - are known to be noisy in time (i.e., varying fast compare to dynamics)
  - can be very sensitive to perturbation
- ▶ Are these expected for deterministic PDE systems?



# The Time-step Convergence Puzzle

- ▶ Experimental design
  - Very short (1-hour) simulations
  - A wide range of time step sizes
  - Solution with shortest step size as reference (i.e. self-convergence)
  - Ensemble runs to take into account possible flow-dependency
- ▶ Convergence rate of the **full model**
  - Expected: 1.0
  - Diagnosed: 0.4
- ▶ Contrast between **dynamical-core-only** and **full-model** results
- ▶ Slower convergence is associated with larger time stepping error

**Time Stepping Error and Convergence Rate in CAM5**





# Value of Convergence Testing

- ▶ For a full-fledged model with complex physics, does it make sense to talk about time-step convergence at all?
  
- ▶ **Our opinion:** For very short simulations, if the solutions do not converge or converge to an unexpected state, then the equations, the discretization, and the coding need to be revisited
  
- ▶ The next slides demonstrate that convergence testing can help identify issues in
  - **Model's continuous formulation**
  - **Physics-dynamics coupling (splitting)**
  - **Time stepping in physics**



# A Test Problem

- ▶ E3SM's dynamical core + a very simple parameterization
- ▶ Equations directly affected by parameterizations are:

Temperature:  $\frac{\partial \bar{T}}{\partial t} = \overline{A_T} + \frac{L_v}{C_p} \overline{Q}$

Water vapor:  $\frac{\partial \overline{q_v}}{\partial t} = \overline{A_v} - \overline{Q}$

Cloud liquid:  $\frac{\partial \overline{q_l}}{\partial t} = \overline{A_l} + \overline{Q}$

**Physics:** bare-bone version of the large-scale condensation scheme in CAM2-CAM4

(Zhang et al., 2003; Rasch and Kristjansson, 1998; Sundqvist, 1978)

**Dynamics (advection):** Spectral-element dynamical core on cubed sphere, 1-degree, 30 layers



# Large-scale Condensation Scheme

- ▶ Basic assumptions
  - Instantaneous condensation
  - Fractional cloudiness
- ▶ Grid-box mean condensation rate

$$\overline{Q} = \underbrace{f \widehat{Q}}_{\text{In-cloud condensation}} - \underbrace{\left( \overline{A}_l - f \widehat{A}_l \right)}_{\text{Clear-sky evaporation}} + \underbrace{\tilde{q}_l \frac{\partial f}{\partial t}}_{\text{Phase change associated with cloud fraction change}}$$

- ▶ Closure assumption:  $\tilde{q}_l = \frac{\overline{q}_l}{f}$

*“When the cloud is growing ( $df/dt > 0$ ), the new cloud water increases to match that within the cloudy part of the grid box. Conversely, when the cloud is eroding ( $df/dt < 0$ ), the cloud water goes to zero in that region.”*

— Rasch and Kristjansson (1998); Zhang et al. (2003)

- ▶ Closure assumption  $\tilde{q}_l = \frac{\overline{q_l}}{f}$
- ▶ If  $q_l > 0$  but  $f \sim 0$ , we get “infinitely dense cloud” (singularity)
- ▶ The use of a “**safeguard parameter**” is a common remedy  $\tilde{q}_l = \frac{\overline{q_l}}{\max(f, f_{\min})}$
- ▶ ... but it can hide problems  
(will show on a later slide)
- ▶ Does such singularity actually occur in the simulations?  
**Unfortunately, yes, and it affects convergence**



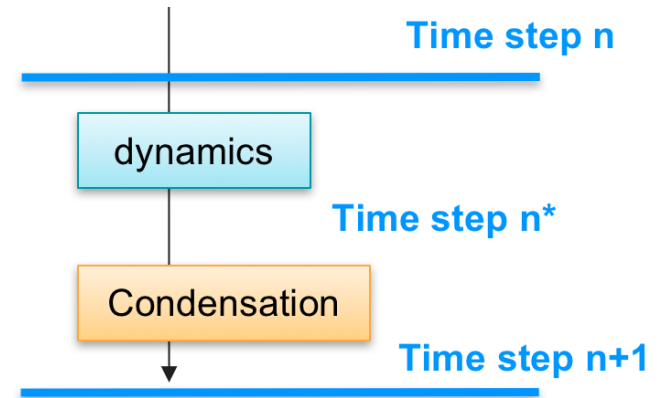
# Impact of Physics-Dynamics Coupling (Splitting)





# Splitting in CAM4

- ▶ Physics and dynamics are sequentially split
- ▶ At the intermediate step  $n^*$ 
  - Model state is out of saturation equilibrium
  - Condensation scheme is expected to bring the state back to equilibrium
  - This is **totally legitimate** (i.e., this is how the parameterization was designed to work)
- ▶ **But** consider this scenario:
  - Advection brings liquid to a very dry cell
  - Condensation scheme evaporates liquid and brings cell back to cloud-free at step  $n+1$
  - The intermediate step  $n^*$  has  $q_l > 0$  and  $f = 0!$





# Revised Splitting for the Closure

- ▶ Original implementation in e.g., CAM4:

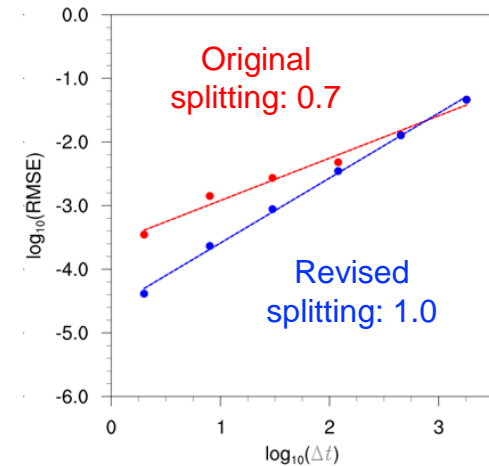
$$\tilde{q}_l = \frac{\overline{q}_l^{(n^*)}}{\max(f^{(n^*)}, f_{\min})}$$

- ▶ Revision: use step  $n$  (in equilibrium)

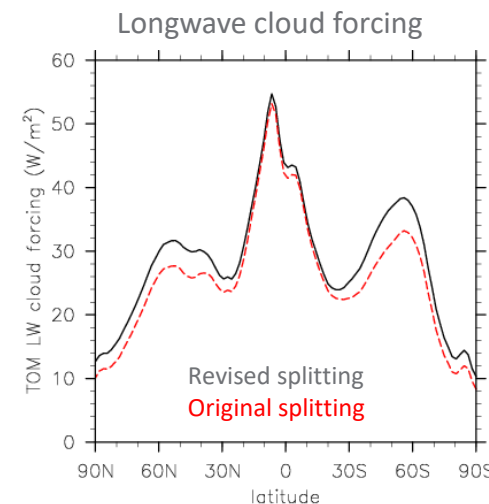
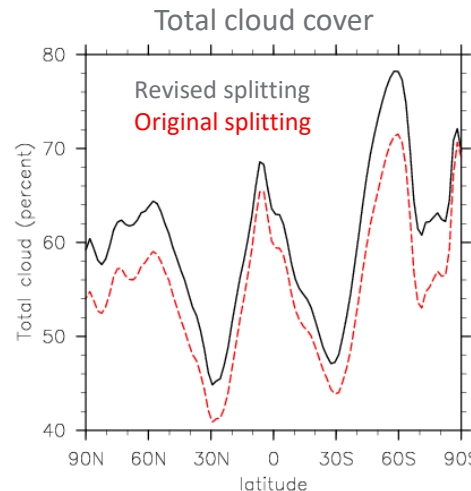
$$\tilde{q}_l = \frac{\overline{q}_l^{(n)}}{\max(f^{(n)}, f_{\min})}$$

- ▶ Helps restore convergence in test problem
- ▶ Also has a substantial impact on model climate in CAM4!

Convergence rate in test problem



Mean climate in full-model simulations with CAM4 physics





# Formal Error Analysis

$$\frac{dy}{dt} = D(y) + P(y)$$

$$|e_n| \leq |\tilde{e}_0|e^{(t_f-t_0)K} + \frac{e^{(t_f-t_0)K} - 1}{2K} \left[ 2K_{f_y} \left\| \frac{y}{f} D(D+P) \right\|_{\infty} + 2K_D \left\| \frac{y}{f} f_y(D+P) \right\|_{\infty} + \|y''\|_{\infty} \right. \\ \left. + 2 \left\| \frac{y}{f} D^2 f_{yy} \right\|_{\infty} + \left\| \frac{y}{f} f'' \right\|_{\infty} + 2 \left\| D \frac{f_y f'}{f^2} \right\|_{\infty} + 2 \left\| D \frac{f'}{f} \right\|_{\infty} \right] \Delta t$$

- Impact of singularity:  $f = 0$  can lead to **unbounded** solution error, hence loss of convergence



# Impact of Model's Continuous Formulation



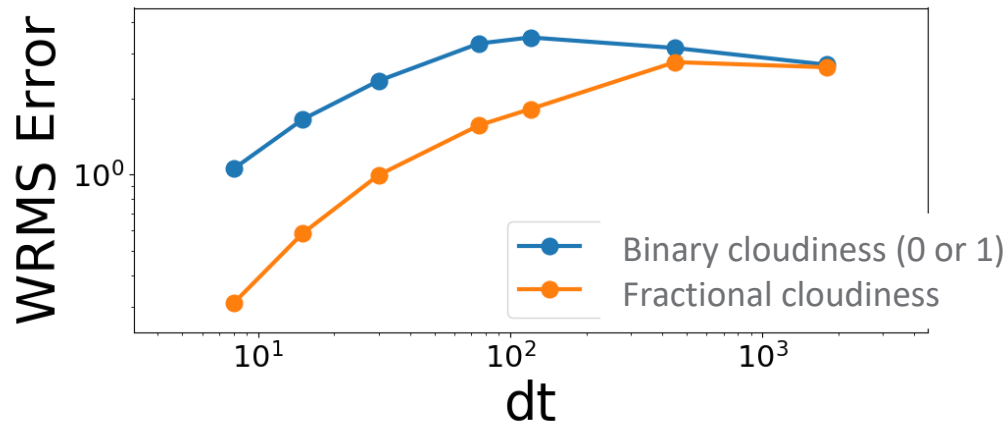
# Cause of Singularity

- ▶ Disconnect between  $q_l$  and  $f$  in basic model setup
  - Cloud fraction based purely on RH (Slingo-type)
  - Liquid concentration predicted by a separate equation
- ▶ Closure assumption  $\tilde{q}_l = \frac{\overline{q_l}}{f}$
- ▶ Revised splitting helps, but convergence can still be lost
  - Within 1 hour -- if initial condition contains singularity
  - In longer simulations -- because singularity can be generated even when it does not occur in the initial conditions
- ▶ Ultimate solution of convergence problem requires revision of model formulation



# A Revised Formulation

- ▶ Keeps the basic setup and assumptions
- ▶ Replaced the closure assumption  $\tilde{q}_l = \frac{\overline{q_l}}{f}$



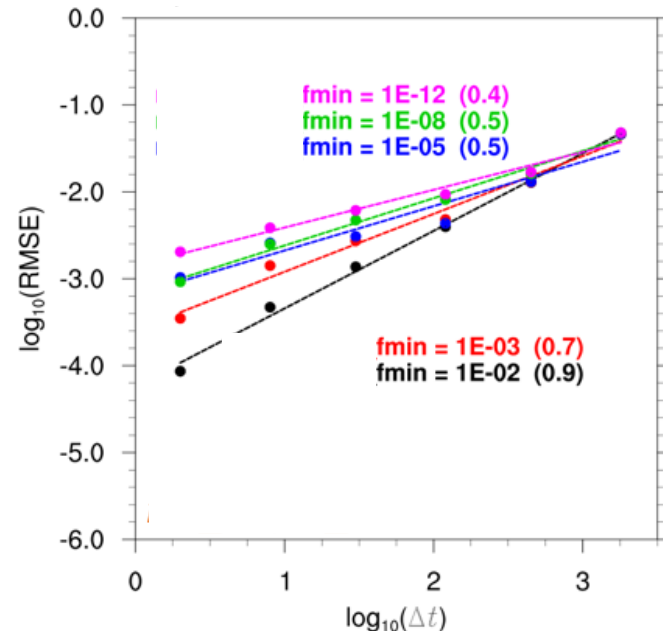
- ▶ Preliminary results:  
Closure without assumption of continuity shows better convergence
- ▶ Another possibility: change the basic setup, choose different prognostic variables (we are interested in exploring this option, too)



# Why Is This Important?

- ▶ What we encountered was essentially a “division by zero” problem, commonly encountered in, e.g.
  - In-cloud hydrometeor concentrations
  - In-cloud aerosol concentrations
  - Skewness of sub-grid PDFs in CLUBB
- ▶ The use of the “**safeguard parameter**” seems to simplify life
- ▶ ... but it can hide problems
- ▶ The revised splitting avoids sensitivity to the artificial parameter
- ▶ The revised formulation does not need the parameter at all.

Original closure assumption and splitting, poor convergence masked by large  $f_{min}$





# Impact of Time Stepping in Physics



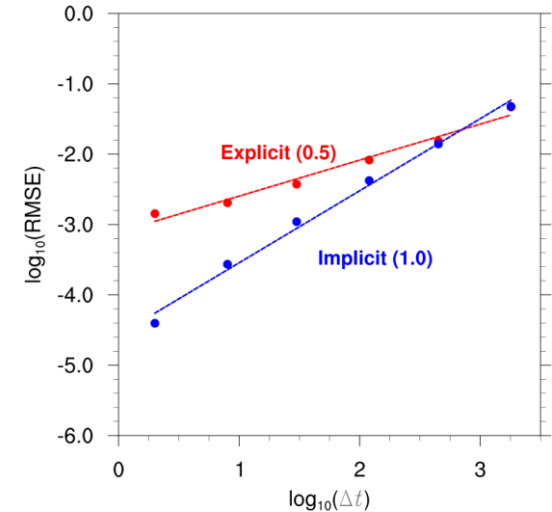


# Cloud Fraction Change

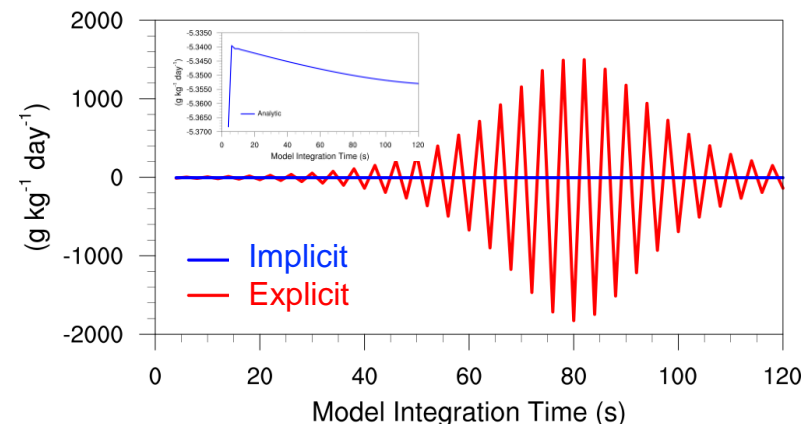
$$\bar{Q} = f\hat{Q} - (\bar{A}_l - f\hat{A}_l) + \tilde{q}_l \frac{\partial f}{\partial t}$$

- ▶ Two options based on a semi-analytic method (Zhang et al. 2003): **explicit** and **implicit**
- ▶ **Explicit** method converges poorly and produces unphysically large oscillations
- ▶ Fast and strongly coupled processes need to be handled with care

### Time Stepping Error and Convergence Rate



### Time Series of Condensation Rate



# Summary

- ▶ E3SM model developers have teamed up with applied mathematicians to address the time-step convergence puzzle in the atmosphere model
- ▶ First results from a simplified model demonstrate that **poor convergence** in short-term simulations **can be understood and improved**
- ▶ Poorly converging and properly converging models can produce **different climate**
- ▶ **Insights** from convergence testing can help improve not only **time integration** but also the **continuous formulation** of a parameterization.
- ▶ We are now working on more complex and realistic equations (i.e., parameterizations in E3SM: CLUBB, cloud microphysics, etc.)



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- ▶ Hydrology and land surface processes
- ▶ Land-atmosphere interactions
- ▶ Cloud observations and remote sensing
- ▶ Climate dynamics and atmospheric feedbacks
- ▶ Terrestrial biogeochemistry
- ▶ Arctic amplification and high-latitude processes
- ▶ Integrated assessment and multi-sector modeling
- ▶ Software engineering to support multi-sector modeling

# Backup slides

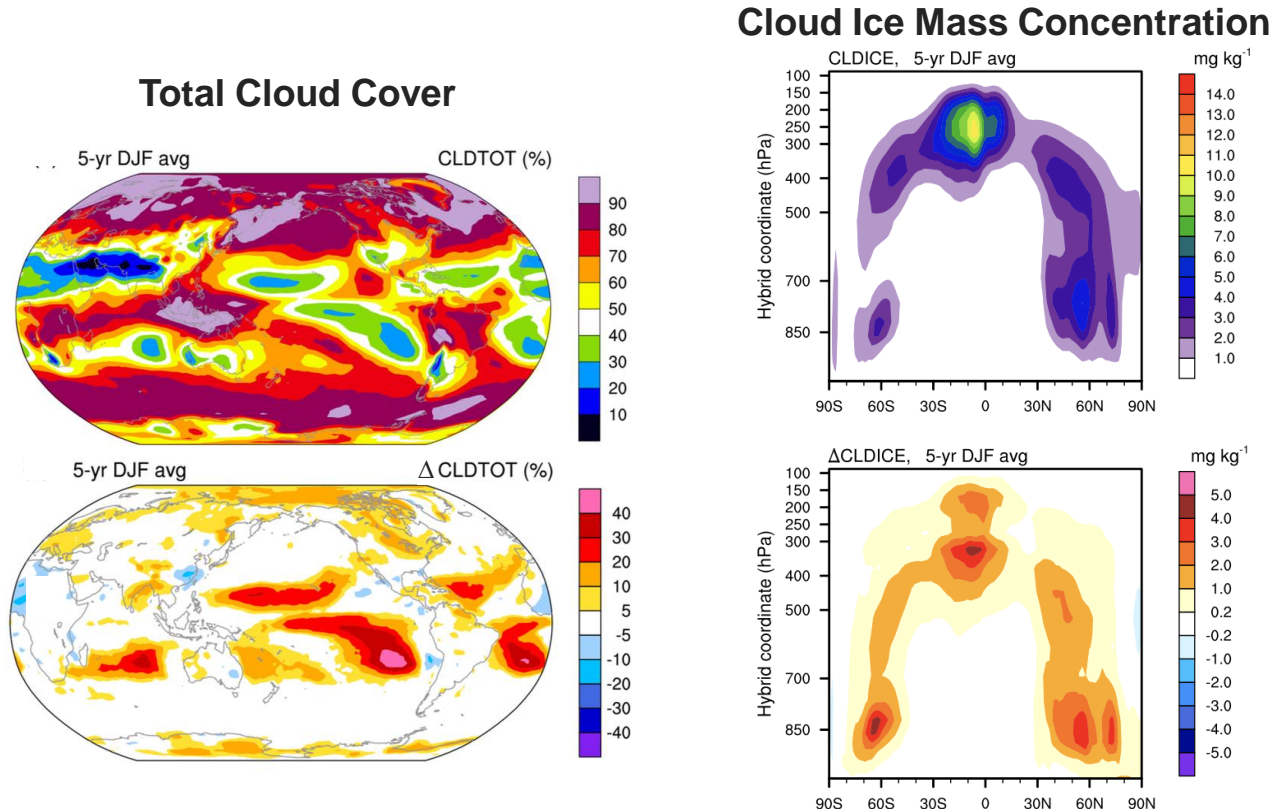


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# Strong Time-step Sensitivity in E3SM and Its Predecessors

- Present-day climate simulated with CAM5



Default model  
(2-degree FV,  
30-min step size)

Difference due to  
reduction of step  
size to 4 min

*Wan, Rasch et al. (2014, GMD)*

- E3SM v0 uses a different dynamical core but shows very similar results, indicating the issues are in the physics package



# Noisy Physics

- ▶ Our strategy for addressing this challenge
  - Identify and remove pathological noisiness
  - Apply deterministic and stochastic PDE theories to handle physical noisiness
- ▶ Initial investigations into prototype problems show promising results
  
- ▶ SDE work inspired by Hodyss et al. (2013, MWR)
- ▶ Our first test problem:
  - 1D advection-diffusion equation with a wide spectrum of fast forcing
  - Generalized Ito correction
  - Improved accuracy and convergence

