

## **Estimating uncertainty** with the ECMWF ensembles

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# Why do we need ensembles?



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## Why do we need ensembles?



## We need ENS even in the short range (US Storm)

Single HRES fcs failed to positioned correctly the storm, and this led to snowfall overestimation for NY of in the 24-36-48h forecasts.

MLSP+TP maps show a 150-200 km eastward shift in the storm centre.







ENS-based probabilistic forecasts can be used to estimate the level of confidence (predictability) of single forecasts. They show that NY was closer to the edge of the area of high probability of +30mm of precipitation, indicating higher uncertainty.



# US Storm, 27-28/01/2015: ENS CDF to estimate confidence

ENS-based probabilistic forecasts expressed in terms of CDF shows that the fcs for NY were more uncertain (the slope of the CDF curves is steeper) than the fcs for Boston.



## A necessary property for ENS to be valuable: reliability

 $M_1$ 

 $M_2$ 

A reliable ensemble has, on average over many cases M, spread measured by the ensemble standard deviation  $\sigma$ , equal to the average error of the ensemble mean  $e_{EM}$ :  $<\sigma>_{M} = <e_{EM}>_{M}$ 

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**e**<sub>EM</sub>

## In a reliable ENS, small spread >> high predictability



In a reliable ensemble, small ensemble standard deviation indicates a more predictable case, i.e. a small error of the ensemble mean  $e_{EM}$ .



# Reliability: <fc-prob>~<obs-prob>

One way to check the ensemble reliability is to assess whether the average forecast and observed probabilities of a certain event are similar.

These plots compare the two probabilities at t+144h for the event '24h precipitation in excess of 1/5/10/20 mm' (top) and '2mT gt/lt 4/8 degrees' over Europe for FMA15 (verified against observations).



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# Reliability: <spread>~<rmse(EM)>

One way to check the ensemble reliability is to assess whether the seasonal average of the ensemble standard deviation and of the ensemble-mean error are similar.

This plot compares these statistics for Z500 over NH in FMA15.



500hPa geopotential NHem Extratropics (lat 20.01090.0, lon -180.0 to 180.0) FebMarApr







### PDF forecast skill: how do we measure it?



- Measure the average distance between the climatological and the observed PDFs with CRPS
- CRPSS = [<CRPS(cli)> <CRPS(fc)>] / <CRPS(cli)>







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# ENS has skill in the monthly forecast range





## **ENS** has skill in the monthly forecast range

The **S2S** (WWRP & WCRP) project is helping us to understand subseasonal to seasonal predictability.

(S2S @ ECMWF web: https://software.ecm wf.int/wiki/display/S2 S/Home)

(from Frederic Vitart)





#### The MJO is important because it has impact on NAO



(from Frederic Vitart)

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#### The seasonal ensemble S4 provides probabilities up to 1y

The tropics remain the area where seasonal prediction has the highest skill, as indicated e.g. by the accuracy of 1-year forecasts of SST anomaly in the Nino3.4 area.



#### How does S4 perform in terms of reliability? 2mT

On average (30 years), 4-6 month probabilistic predictions of 2mT over North America and Europe started in Feb for MJJ (t+4-6m) are reliable and skilful compared to climatology (BSS>0).



BSS PR(2mT>U3)	EU	NA
1 Feb > MJJ (t+ 4-6m)	0.064	0.050
1 Apr > MJJ (t+2-4m)	0.058	0.074

#### ... but for precipitation, reliability is very poor!

On average (30 years), even shorter-range 2-4 month probabilistic predictions of TP over North America and Europe started in Apr for MJJ (t+2-4m) are **not** reliable and **less** skilful than climatology.



BSS PR(TP <l3)< th=""><th>EU</th><th>NA</th></l3)<>	EU	NA
1 Feb > MJJ (t+ 4-6m)	-0.049	-0.052
1 Apr > MJJ (t+2-4m)	-0.072	-0.052





## The EDA is used to estimate flow dependent stats

The 25-member **Ensemble of Data Assimilations** provides the 4DV-HRES with flow dependent background error statistics. Background error correlation length scale for long( $p_{msl}$ ) and  $p_{msl}$ 



## **EDA flow-dependent stats are key to assimilate obs**

This figure shows the differences between EDA-based temperature correlations at ~28 hPa for two points, (45°N;0°E) (blue) and (45°S;0°E) (red) for a day in January (dashed) and June (solid).

The plot shows, e.g., that the SH winter (red solid) and summer (red dashed) temperature correlations are significantly different, a feature that cannot be accurately presented by climatological correlations.



(from Massimo Bonavita)

## How can we keep improving the ensembles?

Work is progressing on many areas to further improve the ECMWF ensembles:

- 1. **Modelling** (including model uncertainty simulation): improve all model components (land, atmosphere and ocean) and increase resolution; upgrade the stochastic schemes that simulate model uncertainty
- 2. Initial Conditions estimation: integrate further the analysis and forecast ensembles (EDA and ENS) and re-assess the potential benefit of starting ENS directly from EDA analyses; assess the impact of using a more strongly coupled DA
- **3. Predictability:** identify sources of predictability, and ways to extract predictable signals for the ensemble PDF
- 4. Ensemble methods: assess whether different ensemble configurations (IC/model unc, membership, truncation, refc suite, ...) could lead to more accurate and reliable PDF fcs

## Forthcoming resolution increase: all IFS components

	<b>Operational suite</b>	Future resolution (preferred option)
HRES	T <sub>L</sub> 1279 (~16km) / L137	T <sub>co</sub> 1279 (~9 km) / L137
4DVAR	T <sub>L</sub> 1279+3xT <sub>L</sub> 255 / L137	T <sub>co</sub> 1279+T <sub>co</sub> 255/319/399 / L137
EDA	T <sub>L</sub> 399 / L137 (25 members)	T <sub>CO</sub> 639 / L137 (25 members)
ENS/MOFC IFS:	51 members TL639 (~ 32 km) (0 - 10d) +T <sub>L</sub> 319 (10 - 46d) / L91	51 members T <sub>CO</sub> 639 (~ 18km) (0 - 10d) +T <sub>CO</sub> 319 (10 - 46d) / L91
OCEAN:	NEMO ORCA100z42	NEMO ORCA025z75

T<sub>co</sub> – Cubic octahedral Gaussian reduced grids

### Predictability: impact of using a higher-resolution EDA

An increase in the resolution of the EDA outer loops from T399 to T639 is expected to improve the estimation of the analysis uncertainty and thus the ENS initial conditions. This figure shows the impact on the ENS spread (std) in the case of typhoon Bolaven (26 Aug 2012).



#### **Ensemble with perturbations from 399 EDA**

Ensemble with perturbations from 639 EDA

#### Very recent results: impact of increasing EDA & ENS resol?



(From Simon Lang)

#### Too small initial spread and deepening not captured

#### Very recent results: impact of increasing EDA & ENS resol?



(From Simon Lang)

#### **Deepening is captured and spread reflects uncertainty**



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#### .. to conclude .. think ensembles!!!





#### .... extra slides



### Forthcoming resolution upgrade: octahedral cubic grids

It is a reduced Gaussian grid with the same number of latitude circles (*NDGL*) than the standard Gaussian grid ( $\leftrightarrow$  Gaussian weights) but with a new rule to compute the number of points per latitude circle.

Number of points per latitude NLOEN( $lat_N$ )=20  $\rightarrow$  Poles NLOEN( $lat_i$ )=NLOEN( $lat_{i-1}$ )+4

TL1279 :2.14 Mpoints TC1023 :5.45 Mpoints TC1279 :8.51 Mpoints TC01279 :6.59 Mpoints



# Forthcoming resolution increase: orographic variance



(from Nils Wedi)