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CLIMATE, OCEAN AND SEA ICE MODELING PROGRAM

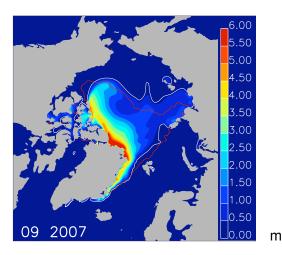
# The state of affairs in modelling sea ice and its atmosphere/ocean interconnections

#### Elizabeth Hunke June 2013



## Monthly Sea Ice Thickness

#### - SSM/I 15% ice concentration



## Interannual variability:

#### wind

air temperature humidity

Ocean Mod. 34, 2010

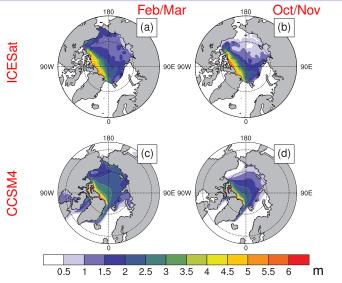
Calculated/ Feedbacks:

> turbulent fluxes radiation SST air temperature

#### 1958–2007 CICE simulation

## Community Climate System Model 4





Jahn et al., "Late 20th century simulation of Arctic sea ice and ocean properties in the CCSM4," J. Clim., 2012

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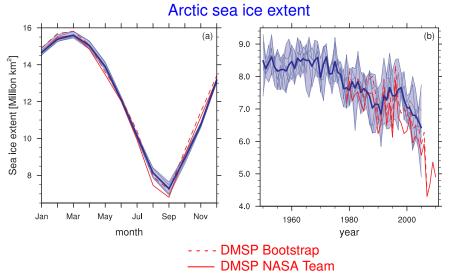
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## **Community Climate System Model 4**





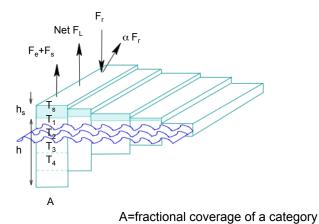
Jahn et al., "Late 20th century simulation of Arctic sea ice and ocean properties in the CCSM4," J. Clim., 2012

Ice-ocean coupling

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## Ice Thickness Distribution g

## Schematic of model representation of g(H) in five ice "categories"



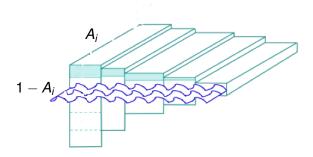
Slide courtesy Dave Bailey, NCAR

Ice-ocean coupling

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## Ice Thickness Distribution g

Schematic of model representation of g(H) in five ice "categories"

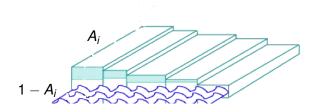


Ice-ocean coupling

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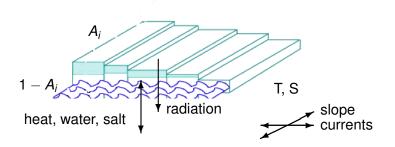


Ice-ocean coupling

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## Ice Thickness Distribution g

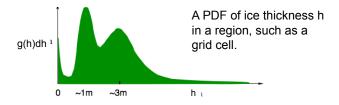
Schematic of model representation of g(H) in five ice "categories"



## Ice Thickness Distribution g

 $g(\mathbf{x}, h, t) dh$  = the fractional area covered by ice in the thickness range (h, h + dh) at a given time *t* and location **x** 

$$rac{\partial oldsymbol{g}}{\partial t} = - 
abla \cdot (oldsymbol{gu}) + \psi - rac{\partial}{\partial oldsymbol{h}}(oldsymbol{f}oldsymbol{g}) + oldsymbol{L},$$



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## Ice Thickness Distribution g

 $g(\mathbf{x}, h, t) dh$  = the fractional area covered by ice in the thickness range (h, h + dh) at a given time *t* and location **x** 

$$\frac{\partial \boldsymbol{g}}{\partial t} = -\nabla \cdot (\boldsymbol{g} \boldsymbol{\mathsf{u}}) + \psi - \frac{\partial}{\partial \boldsymbol{h}} (\boldsymbol{f} \boldsymbol{g}) + \boldsymbol{\mathsf{L}},$$

- $\nabla = \left(\frac{\partial}{\partial x}, \frac{\partial}{\partial y}\right)$
- u = horizontal ice velocity
- $\psi =$  mechanical redistribution function
- f = rate of thermodynamic ice growth
- L = lateral melting

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### What Still Needs to be Done in Sea Ice Models?

Depends on what you want to use the model for

- Notz. Challenges in simulating sea ice in Earth System Models. Wiley Interdisc. Rev.: Climate Change, **3**, 509–526, 2012.
  - radiation melt ponds, zenith angle, soot, extinction coefficient, emissivity
  - Snow improved thermodynamics, redistribution
  - ice-ocean interface stratification effects, friction velocity, false bottoms
  - multiphase physics heat capacity, heat partitioning, porosity, bulk salinity, grease ice
  - floe size distribution lateral melt, form drag
  - dynamics roughness length, rheology, anisotropy, fast ice

#### **Observations needed**

## Ice at the Interface

Atmosphere-Ice-Ocean Boundary Layer Processes and Their Role in Polar Change

Toward better understanding of boundary processes in the atmosphere-ice-ocean system and their contribution to biogeochemical cycling within the climate system.

#### June 25-27, 2012 • Boulder, Colorado USA











## **Scientific Outcomes**

- Theme: Episodic and extreme events fracturing, DMS emissions, clouds, freezing fronts in ice
- Theme: Stratification mobility, exchanges with deeper ocean, stronger currents
- Theme: Precipitation Snow, surface water, clouds/storms, NH/SH differences
- Theme: Marginal ice zone wave action, bio/chemical interactions, floe size
- Model development

priorities: Snow physics, fluid flow within ice, ice-ocean dynamics (especially roughness length), clouds/radiation outlook: under-ice ponds, superimposed ice, flooding and snow-ice formation, floe size distribution

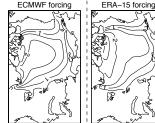
## Predictability: My (sea-ice-centric) View

#### sea ice responds to the forcing applied to it (atm, ocn)









Sep thickness (m)

Bitz, Fyfe, Flato. Sea ice response to wind forcing from AMIP models. J. Clim. 2002.

#### sea ice amplifies the applied change in either direction

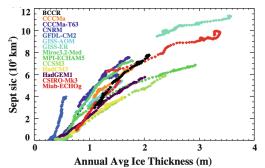


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## Predictability: My (sea-ice-centric) View

sea ice predictability is critically dependent on

- predictability of the applied forcing
- ice equilibrium state associated with the applied forcing

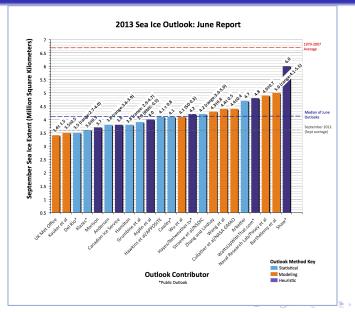


Holland, Serreze, Stroeve. The sea ice mass budget of the Arctic and its future change as simulated by coupled climate models. Clim. Dyn., 2010.

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need to focus sea-ice-specific work on the strength of feedbacks (including atm/ocn fluxes)

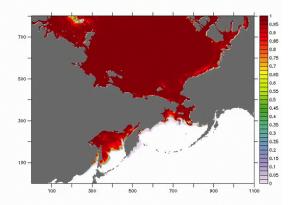
### June 2013 Sea Ice Outlook



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## Ice-Ocean Coupling Issues

#### 15-day fully-coupled high-resolution CESM simulation



atm: ne120 cam4 ocn, ice: 0.1°

ice-atm  $\Delta t_{cpl} = 15$  min ice-ocn  $\Delta t_{cpl} = 2$  hr OK for 1, 6 hr

ice  $\Delta t_{dyn} = 2.5$  min ocn  $\Delta t = 5$  min

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#### courtesy Mathew Maltrud

## Summary

sea ice modeling for climate projection: First-order processes are well captured growth, melt, mobility, deformation, albedo ... for shorter-term prediction, earth system modeling: Second-order processes are being implemented ponds, salinity, vertical transport, snow density, anisotropy ...

#### This Modeler's Perspective on Sea Ice Fragile Relatively passive external forcing balance is critical for simulating base state strength of feedbacks depends on base state Much remains to be understood