

# Model uncertainties in climate prediction: Don't forget the oceans!

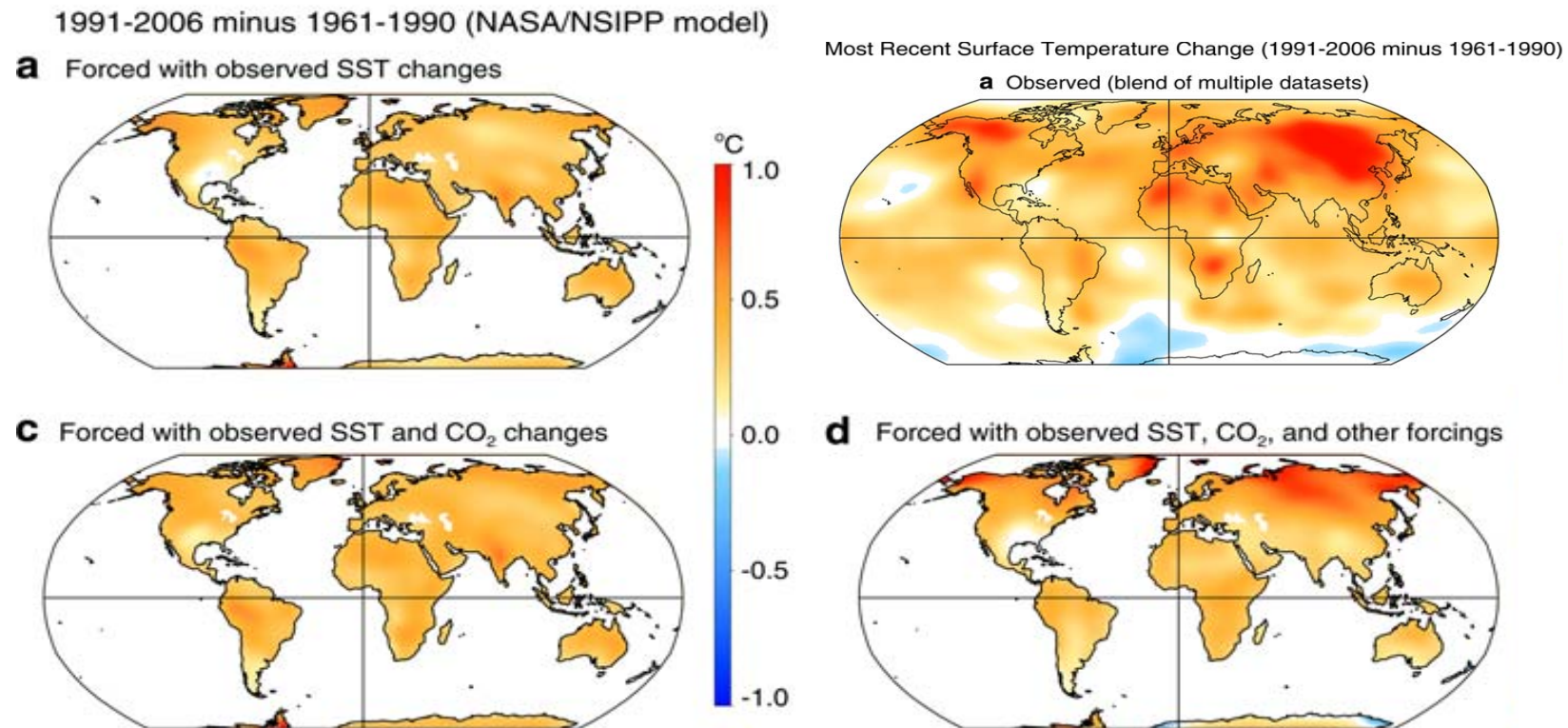
Laure Zanna (Oxford Physics)

ECMWF, June 21<sup>st</sup> 2011

# Outline

- Role of ocean in recent continental warming
- Examples of uncertainties & errors in ocean models
- Recent studies dealing with uncertainties, high-latitudes processes
- Possible Future directions: Observations to reduce uncertainties, Stochastic physics
- → not enough to rely on global climate models, stronger links between theory/obs and modeling centers

# Continental warming influenced by ocean temperatures



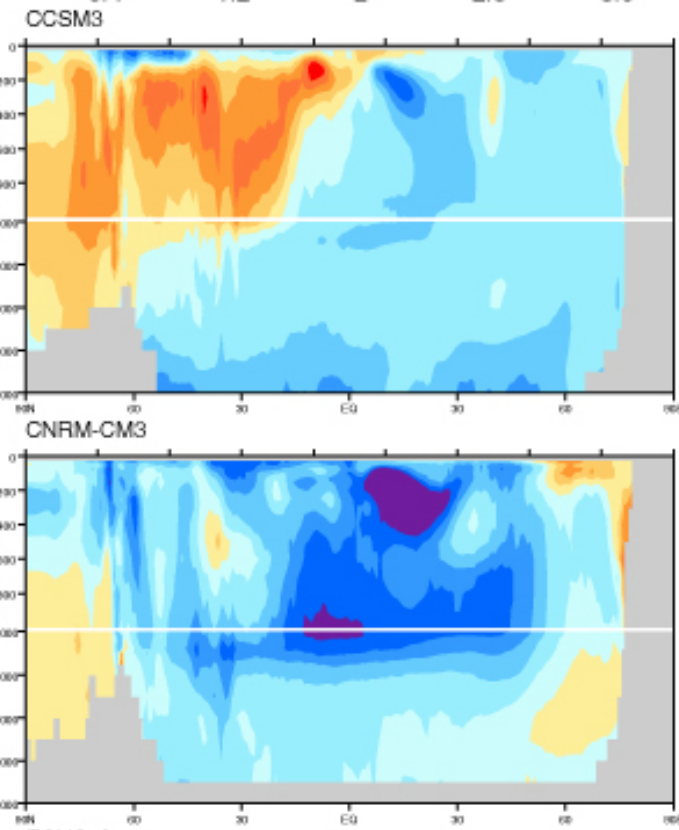
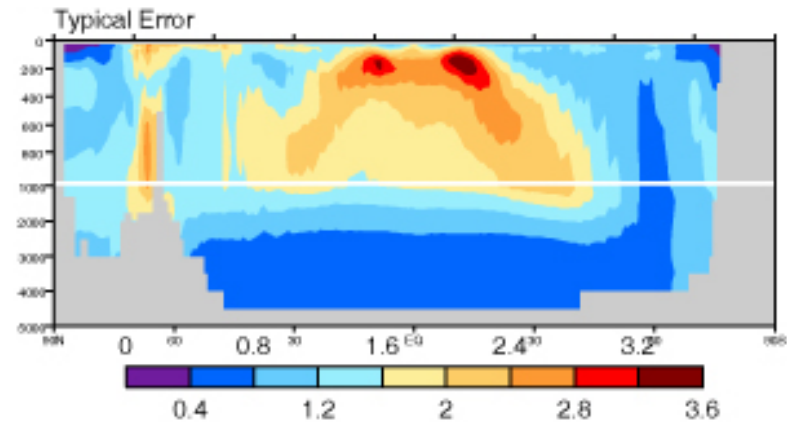
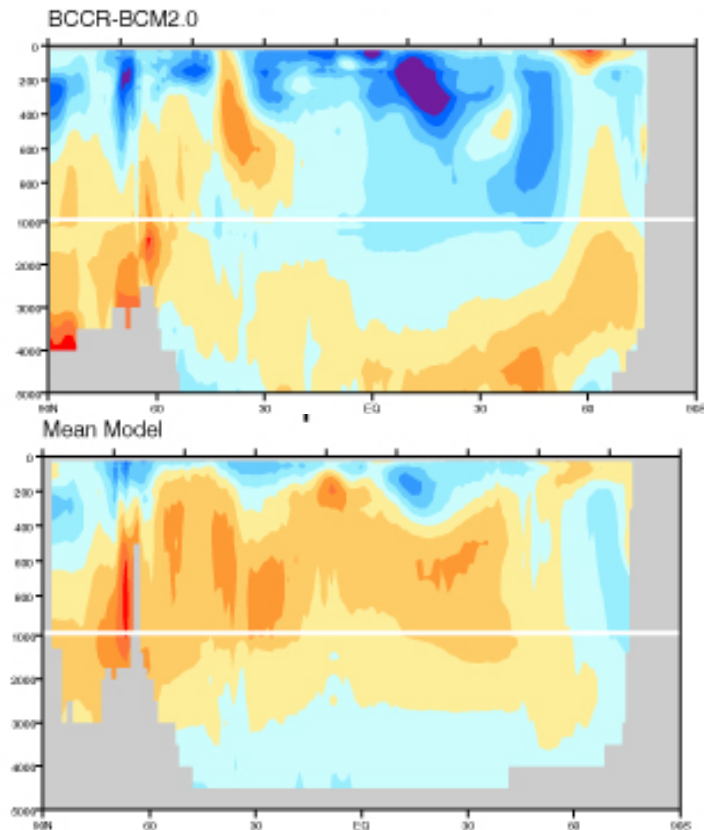
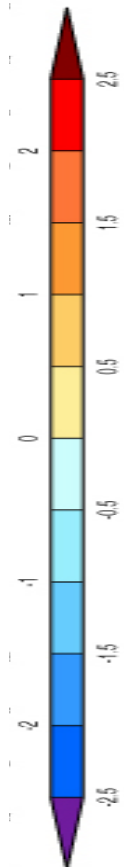
*Winton et al, 2010, Compo & Sardeshmukh, 2009*

Surface & ocean interior properties are important including circulation

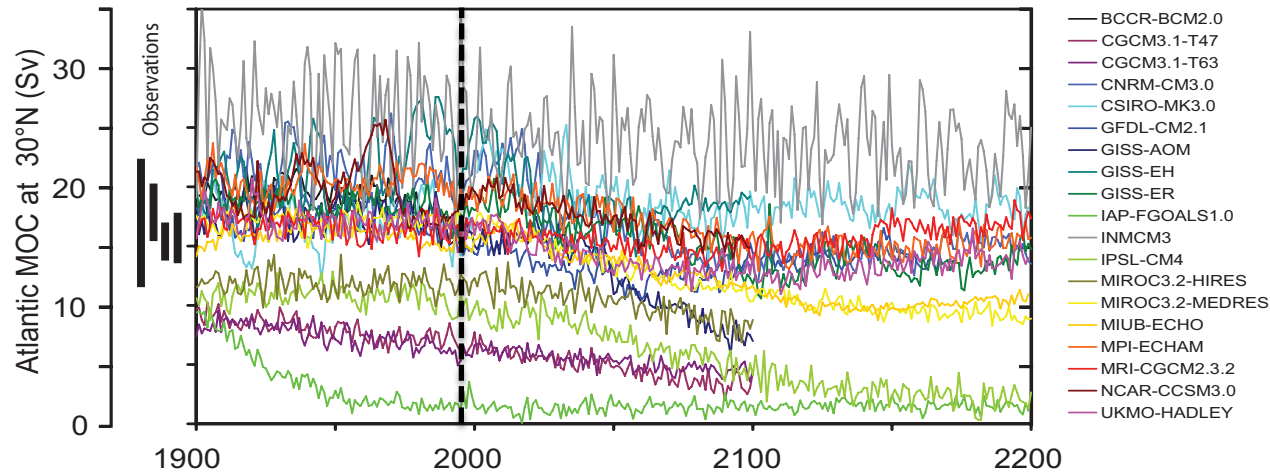
# Multi-model Error in Temperature

Zonal Mean global ocean potential temperature difference (C)

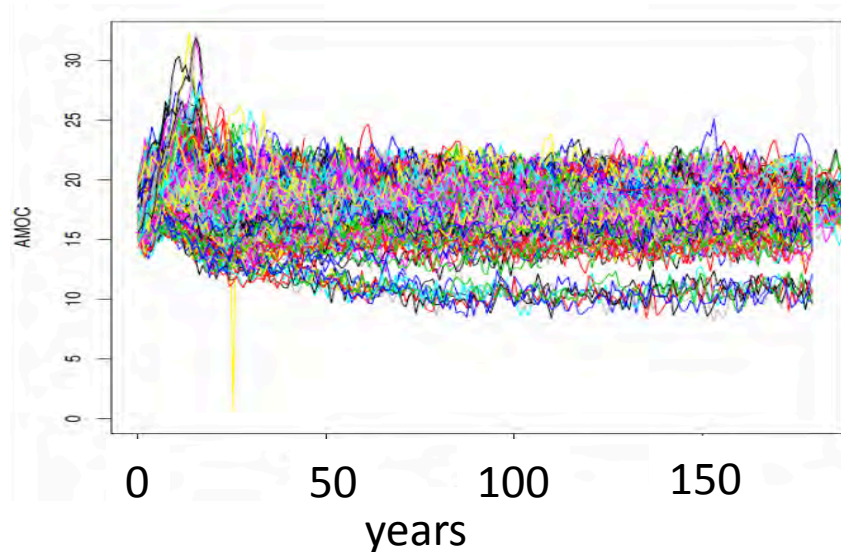
*(IPPC AR4, Ch.8 supp)*



# Climate Projections



- Multi- Model, AR4

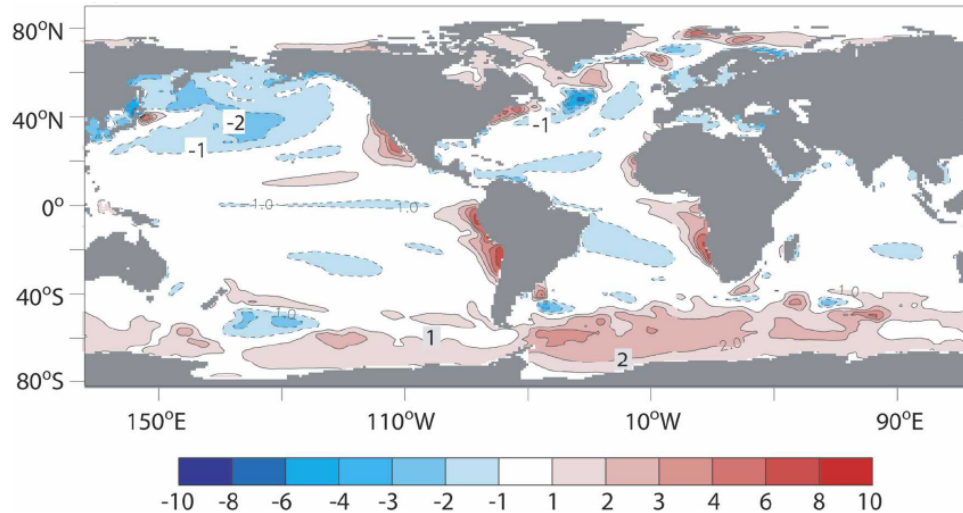


- Climateprediction.net ensemble (~700 members) with FAMOUS (*Yamazaki et al*)
- Seven CO2 emission scenarios

- Not easy to understand the behavior of the models and uncertainties

# Singular Vectors in IPCC AR4 model

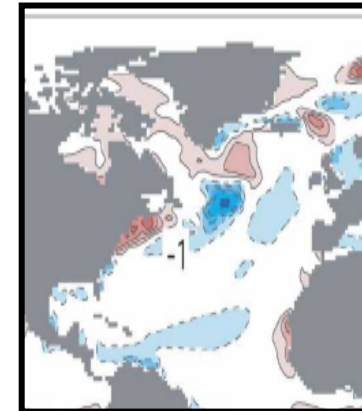
Sea Surface Temperature: Model minus Observations



*GFDL CM2.1*

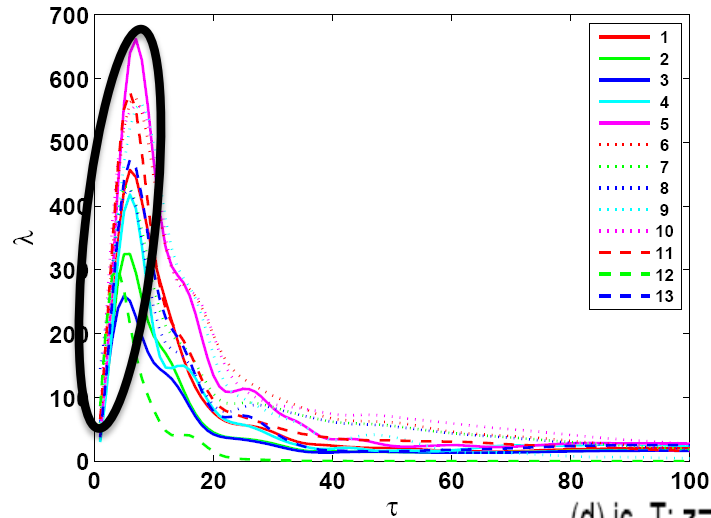
*(Delworth 2006; Delworth et al. 2006; Gnanadesikan et al. 2006; Griffies & Coauthors 2005; Stouffer et al. 2006; Wittenberg et al. 2006)*

- 1000 years of control run from GFDL CM2.1
- North Atlantic annually averaged temperature and salinity fields
- Reduced space based on EOFs



# SVs to detect most sensitive regions

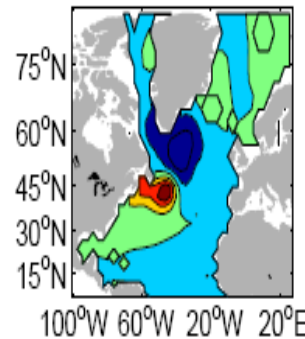
Maximum MOC Amplification Curves



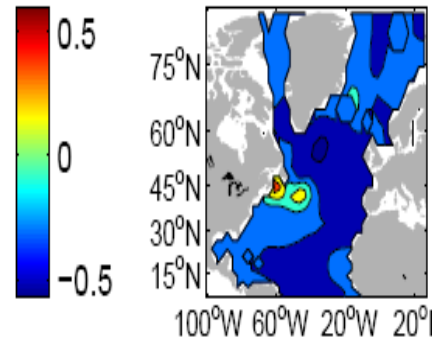
→ Maximum growth of energy and MOC 5-8 yrs  
Can be used to sample uncertainties

(Tziperman, Zanna & Penland 2008)

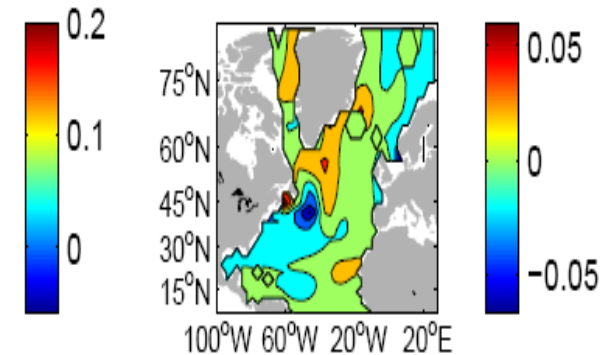
(d) ic, T; z=45m



(e) ic, S; z=45m



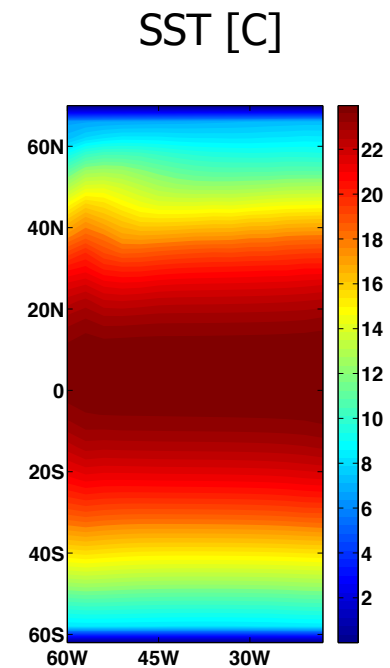
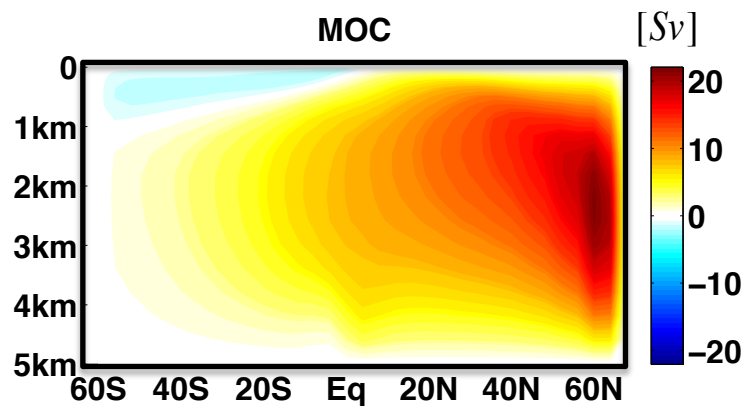
(f) ic, ρ; z=45m



- Build on reduced space; the SVs could potentially project on higher order EOFs (Similar analysis in HadCM3, Hawkins & Sutton, 2009)
- Can be used to initialize climate predictions

# SVs in idealized ocean MITgcm

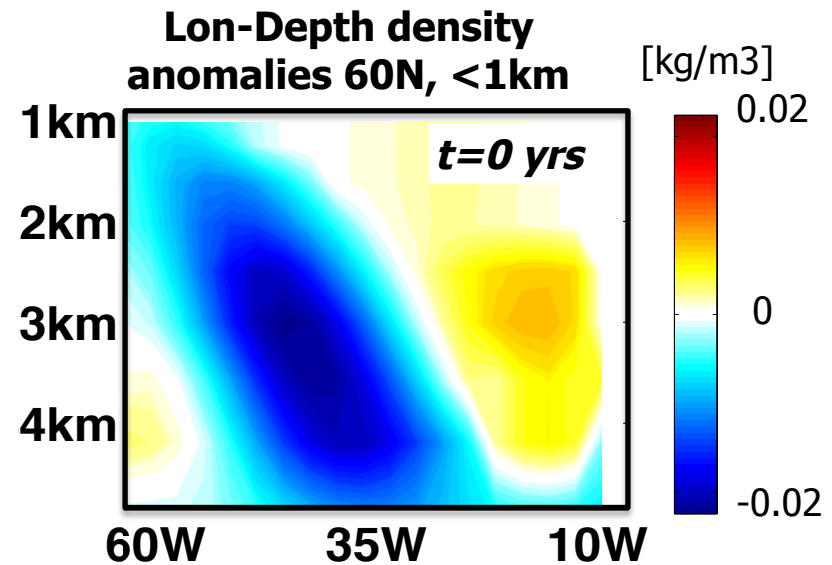
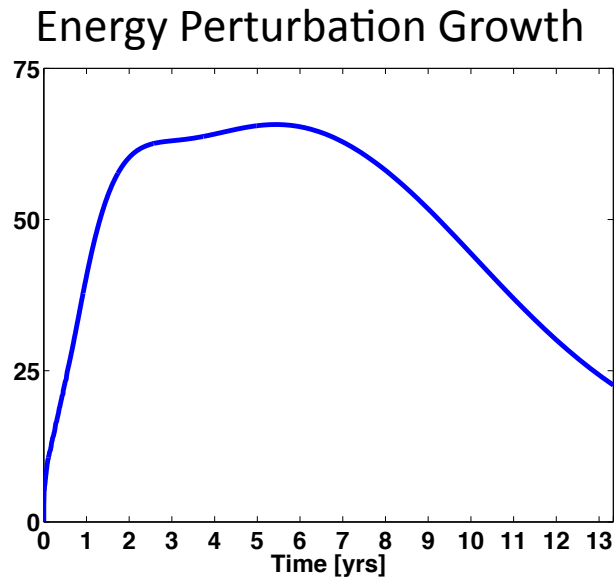
Primitive equations,  $1^\circ \times 1^\circ$ , 15 Levels, Annual averaged Wind & Buoyancy forcing (*Marshall et al, 1996*)





# SVDs in idealized ocean MITgcm

(Zanna et al 2011)

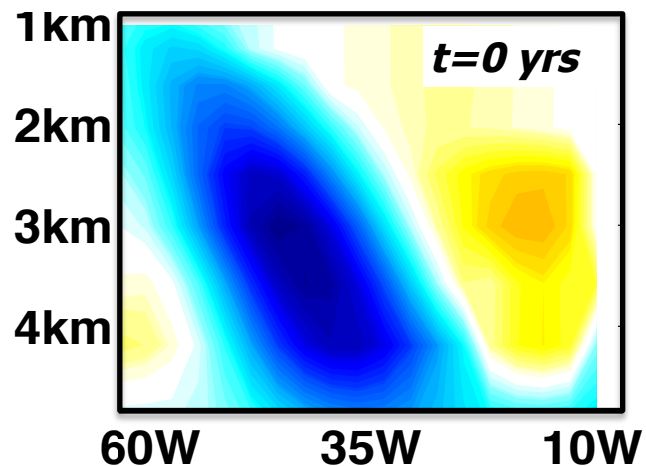


- **Growth** → conversion of mean available potential energy into perturbation kinetic and potential energy
- Perturbations “leaning” against the mean flow (~baroclinic instability)

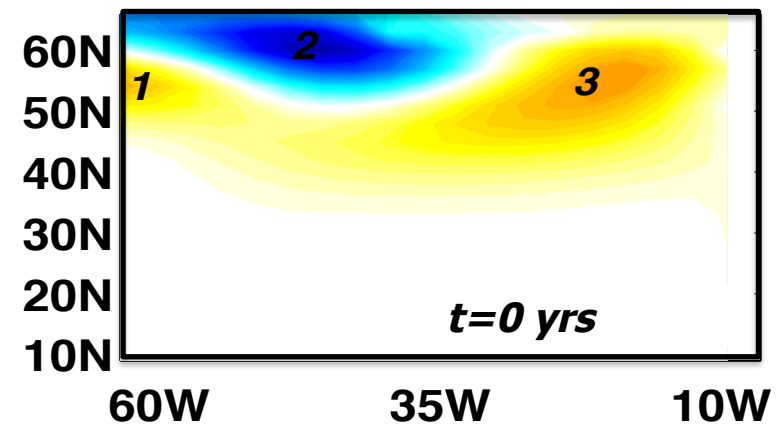
# Leading 3D Singular Vector

- = largest sensitivity of MOC' to perturbations at depth & high-latitudes

Lon-Depth density anomalies 60N, <1km [kg/m<sup>3</sup>]



Density anomalies @ 3000m



- Perturbation estimate:  $\vec{P}'_0 = 0.1C \rightarrow MOC'(7.5\text{ yrs}) = 2.4\text{ Sv}$  (12% mean)
- Errors at high latitudes, at depth in ocean i.c. & model representation (overflows, eddies, deep convection) limit predictability; large impact on the ocean and climate
- Additional observations and better parameterizations are necessary

# High-latitudes ocean processes are important for climate

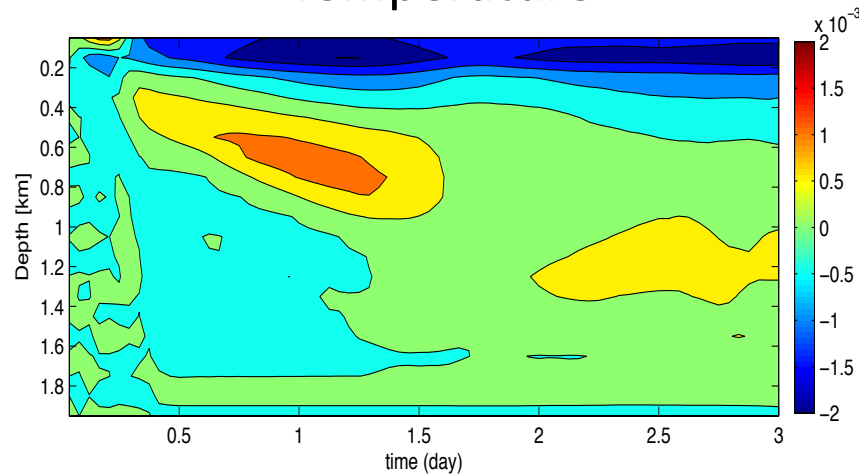
- Upper ocean dynamics = communication between the atmosphere & the oceanic reservoir of heat, freshwater & CO<sub>2</sub>
  - Small-scale & local processes impact the large-scale ocean circulation and uptake of tracers (temperature + carbon)
- Mesoscale & microscale variability (turbulent mixing due to breaking internal waves & convection) are sub-grid scale & are parameterized; most models have similar parameterizations
- Examples of new parameterizations for deep convection and eddy-mixed layer

# Open Ocean Deep Convection

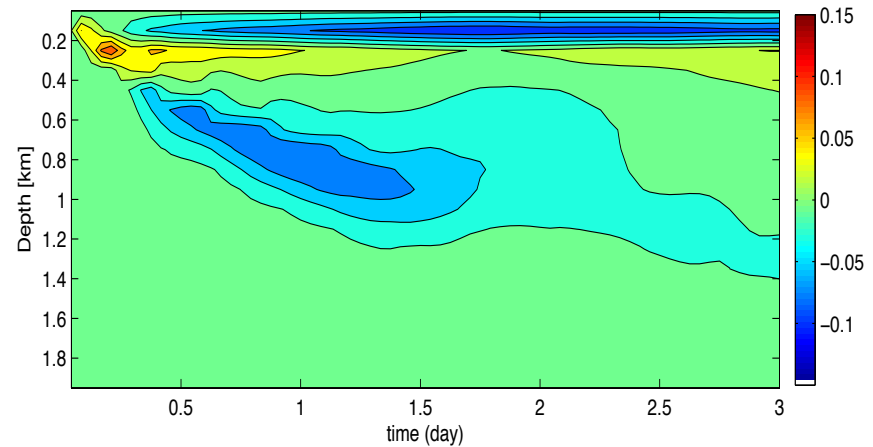
Deep convection: role in global ocean circulation & heat and carbon uptake

Difference between Hydrostatic and Non-hydrostatic

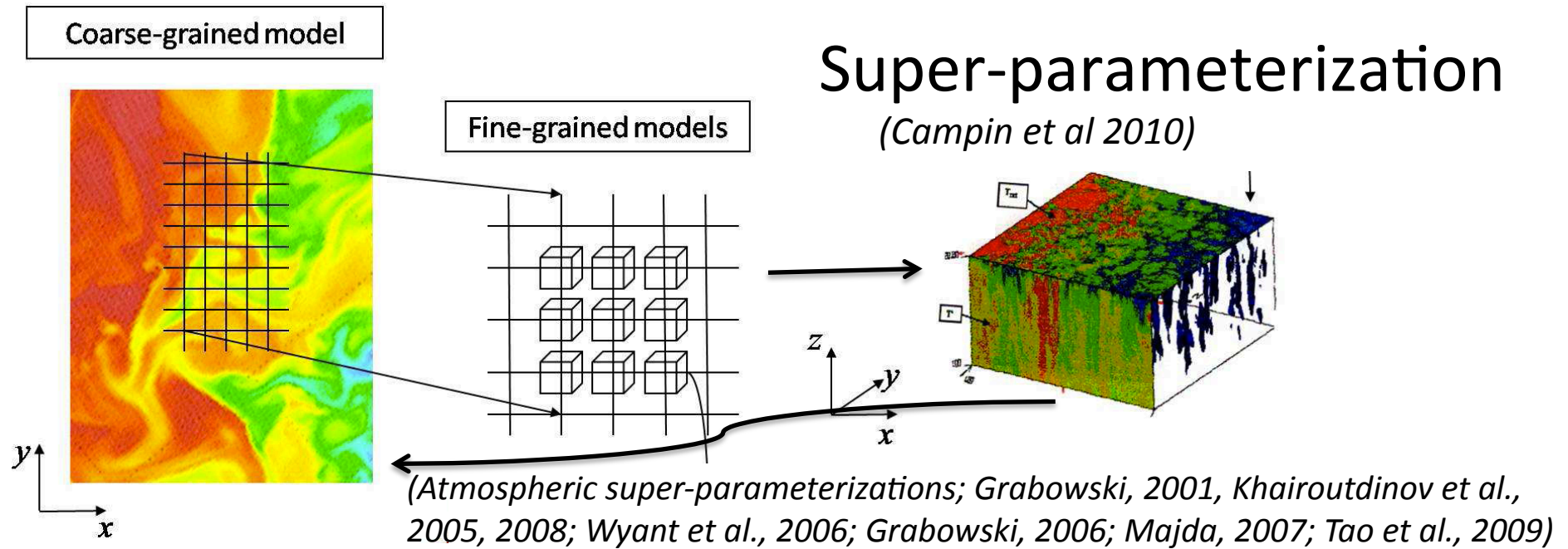
Temperature



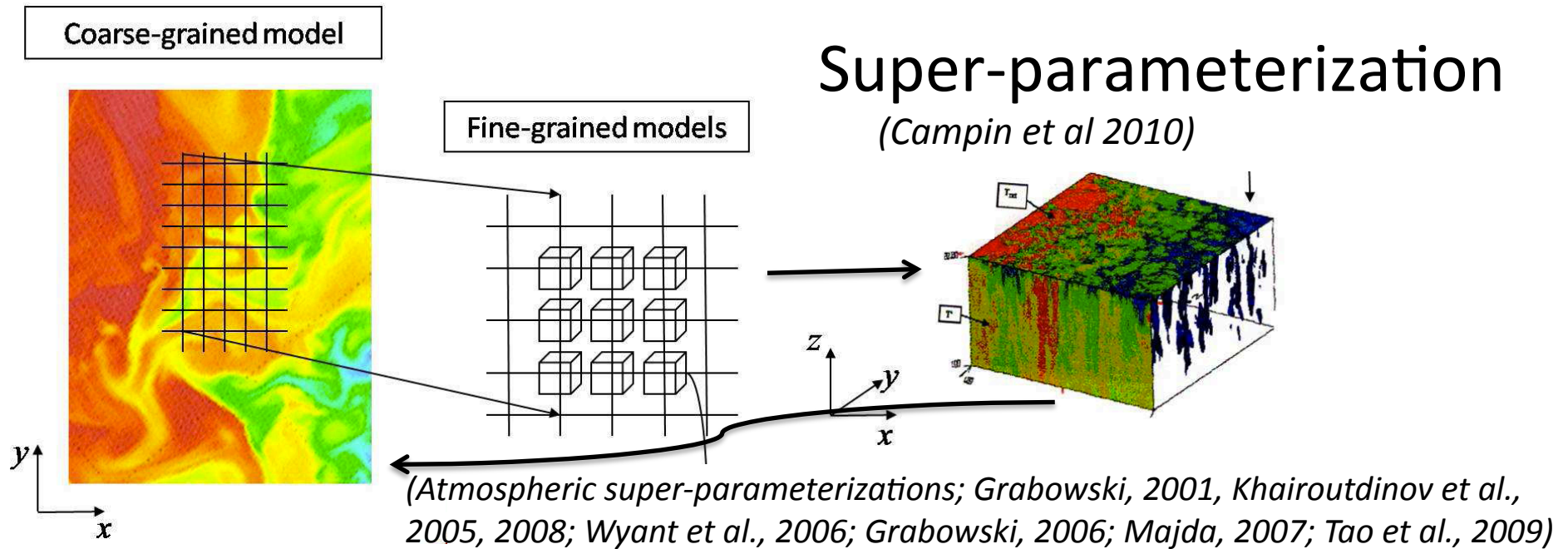
Tracer



# Open Ocean Deep Convection

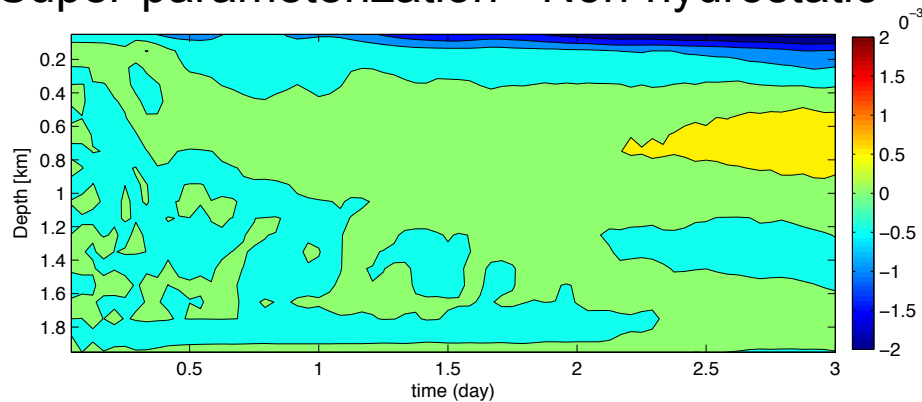


# Open Ocean Deep Convection

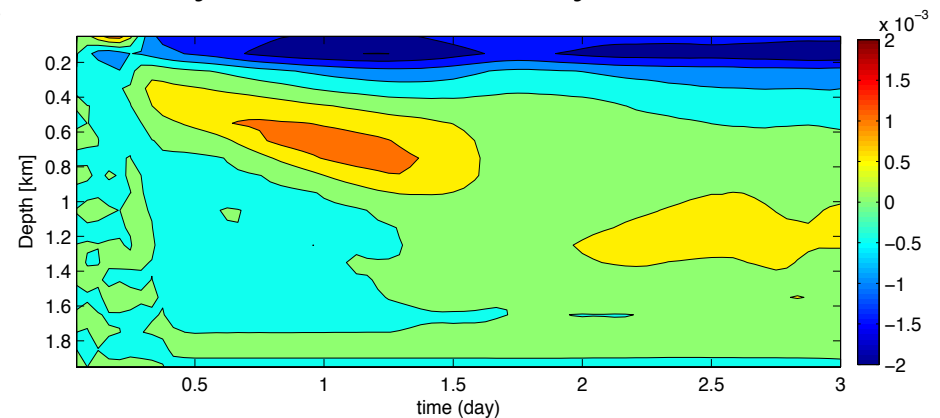


Temperature

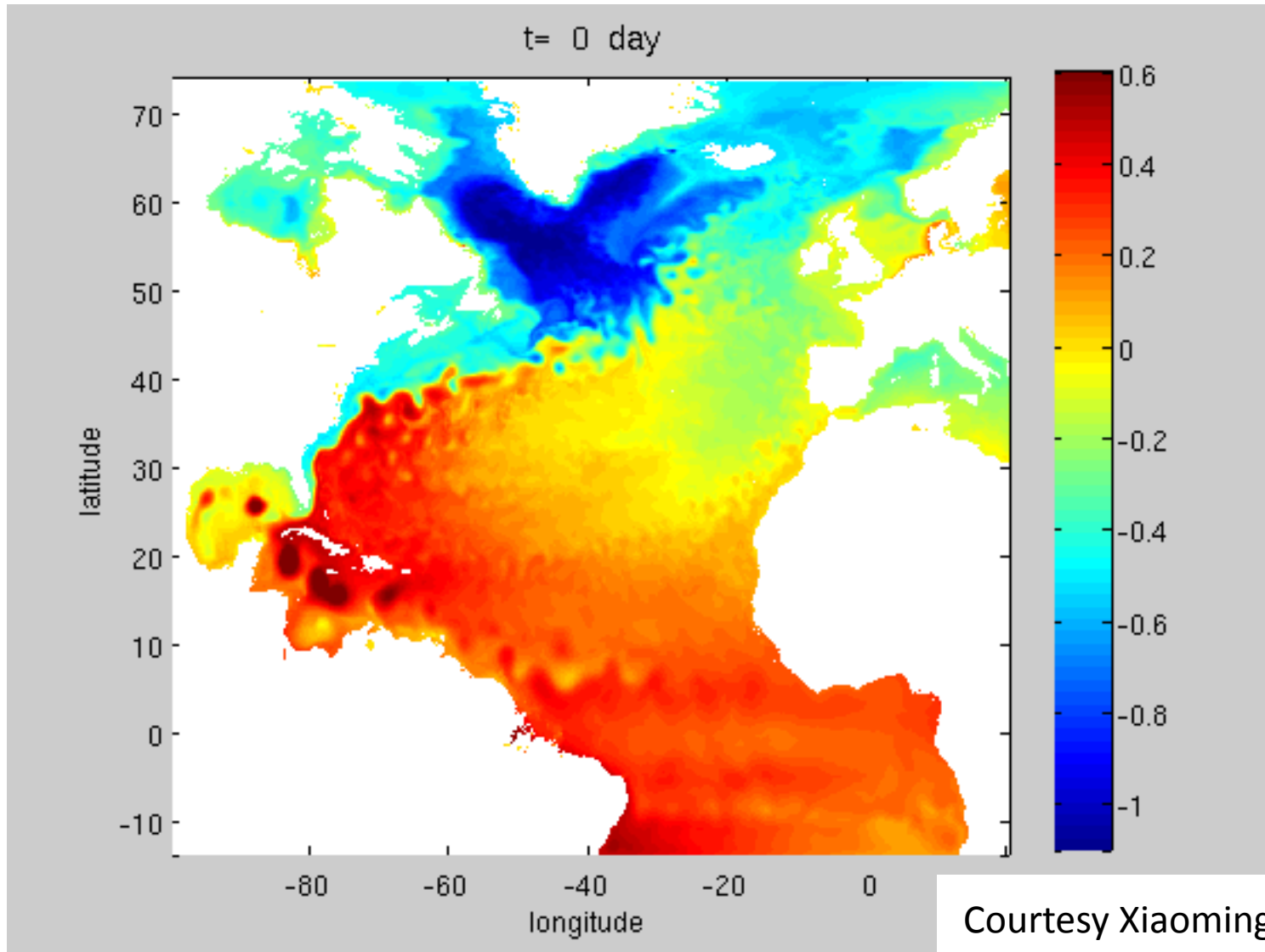
Super-parameterization - Non-hydrostatic



Hydrostatic - Non-hydrostatic



# Sea level height – 1/10° eddy resolving simulation

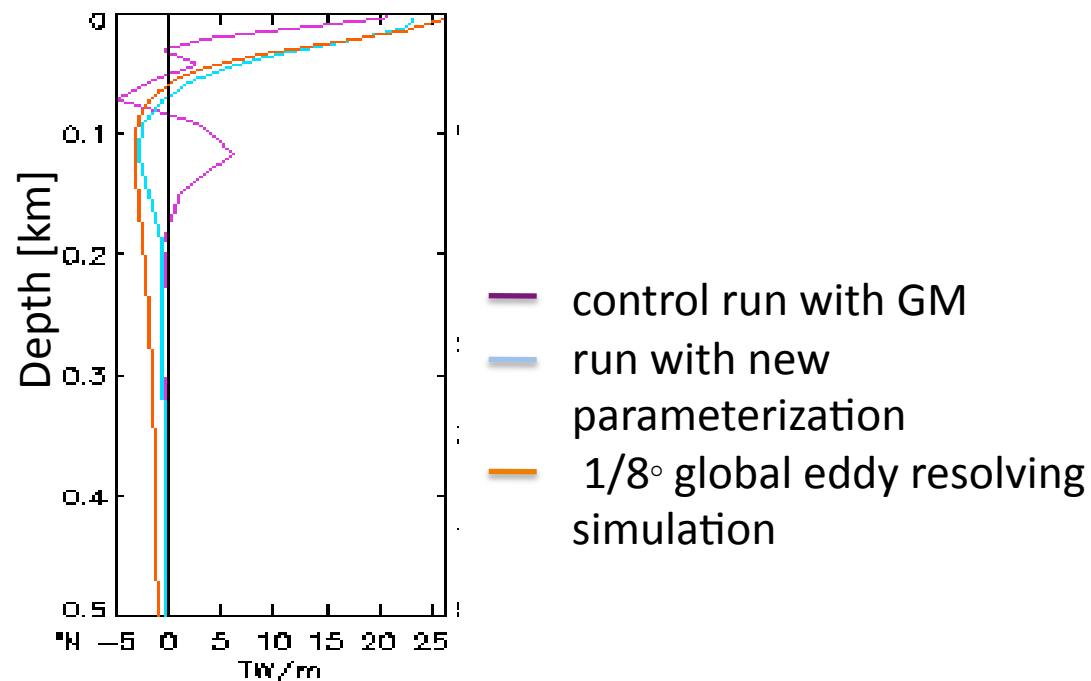


# Eddy-Mixed Layer Interactions

- Mesoscale eddies: Ocean interior = Gent-McWilliams parameterization (adiabatic eddy-induced velocity); *Turbulent BL = eddy induced velocity with zero shear* (well-mixed BL models) + an along-boundary down-gradient flux of density (diabatic mesoscale eddies in the BL)

*Fox-Kemper et al 2010*

zonally averaged heat flux across 47°S

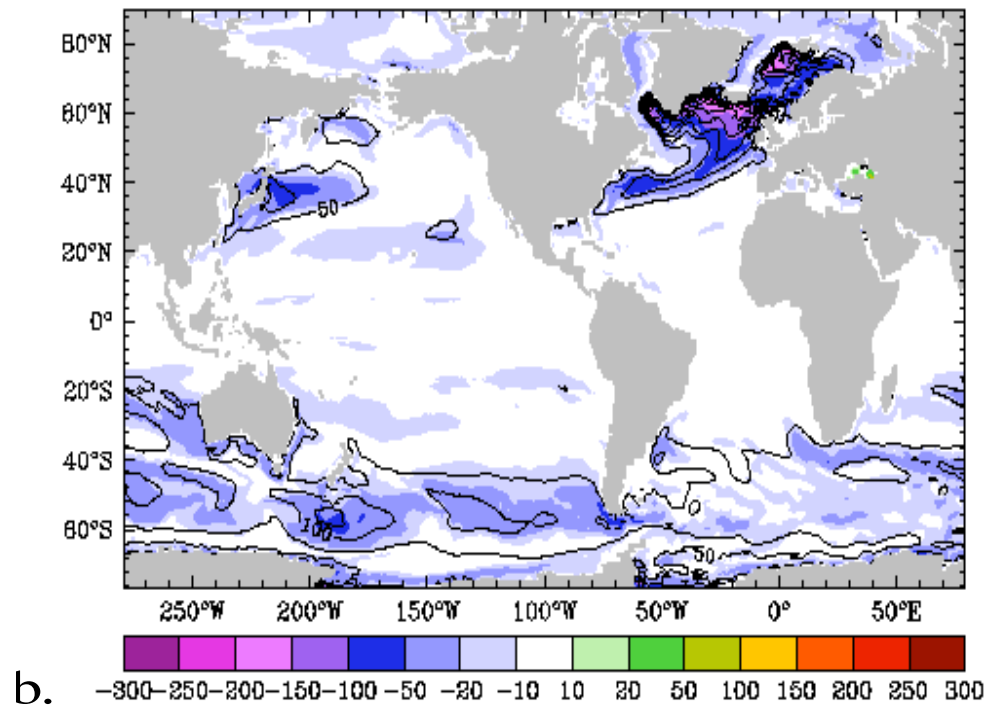




# Eddy-Mixed Layer Interactions

- Submesoscale eddies: buoyancy gradient & front development

Mixed layer depth changes after 10 yrs between control run & run with submesoscale restratification



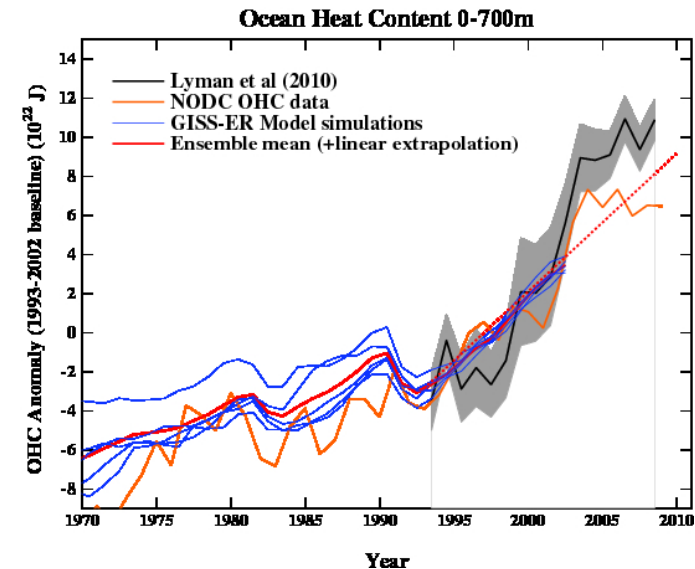
*Ferrari et al 2008*

# Future Directions

- Using observations to constrain & test the models especially on regional scales
- Stochastic physics in ocean models
- → Linking theory/obs /idealized studies with global climate models is crucial

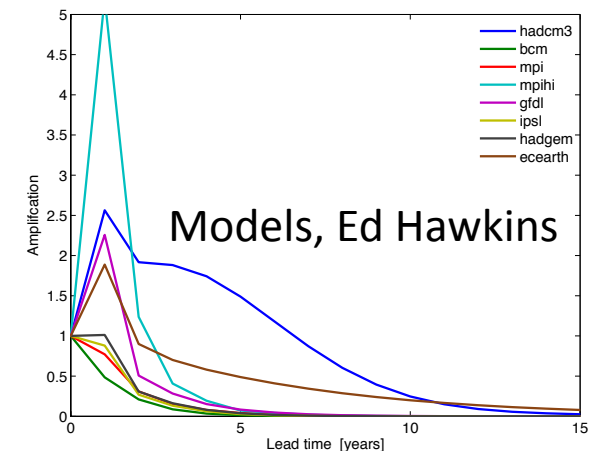
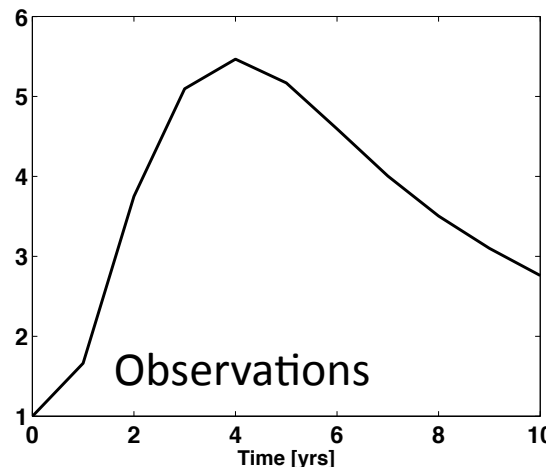
# Using observations to reducing uncertainties

- Ocean heat content, ARGO & altimetry: large uncertainties with obs, analysis & models; can be used to reduced model uncertainties to increasing CO2



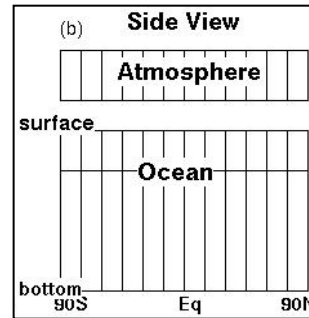
- Regional statistical models based on observations can be used as benchmark for IPCC models

Annual averaged Atlantic SSTs, maximum amplification curves  
(Zanna 2011)

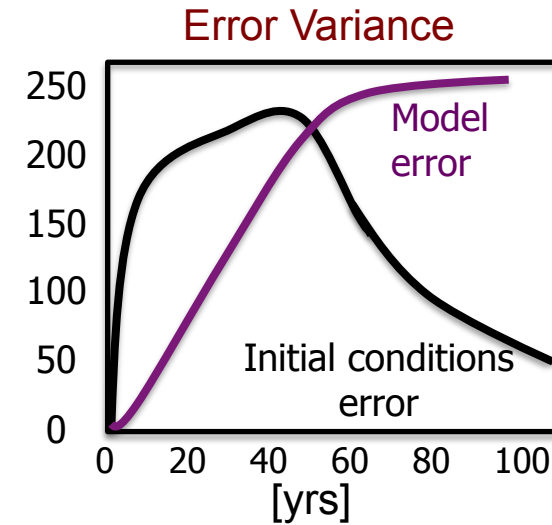


# Role of Stochasticity

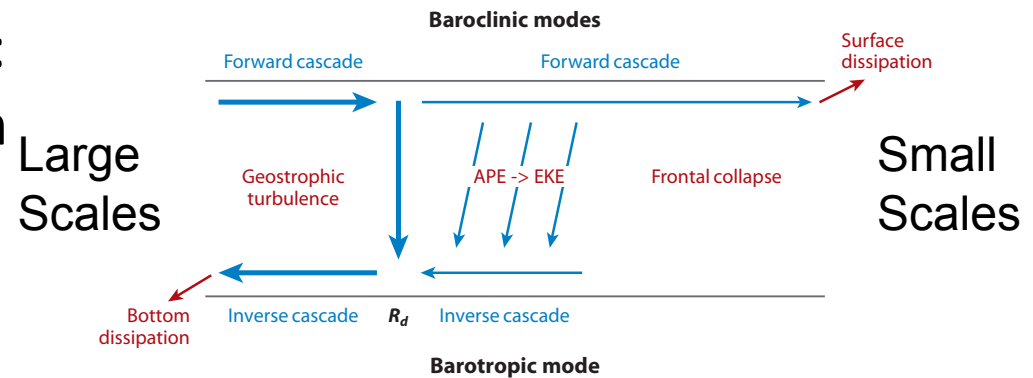
- Stochastic Physics in simple model of the ocean circulation



Zanna Tziperman 2008



- Stochastic parameterization: turbulent mixing & convection



- Implementation of stochastic physics in ocean models and coupled ensemble data assimilation