

Non-hydrostatic modeling with HARMONIE

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Outline



- Short description of HARMONIE
- Behaviour of AROME in HARMONIE
- Implementation of Finite Element discretization in the vertical



The HARMONIE model

- Collaboration HIRLAM $\leftarrow \rightarrow$ ALADIN
- IFS dynamical core
 - Non-hydrostatic
 - Limited area version
 - Physical parameterizations
 - ALADIN
 - HIRLAM
 - ECMWF
 - ALARO
 - AROME



Brac-HR workshop

- BRain-storming on Advanced Concepts on High Resolution modeling
- Should the convection be parameterized (at least in part) at 2.5 km resolution?
- Implicit or explicit methods?

Convection in HARMONIE

- Deep convection and outflow
 - Impact resolution
 - Impact horizontal diffusion
 - Impact SLHD (Semi-Lagrangian Horizontal Diffusion)



Impact resolution

ALARO physics without deep convection ~AROME

Radar NL 20060430 1030 UTC





• Satellite 2, 1, 0.5 km









Impact of resolution

- Increasing resolution from 2 km to 1 km increases maximum precipitation, but it decreases from 1 -> 0.5 km. What happens at even higher resolution?
- Scales are not going to observed open cell scales
- Impact on onset of deep convection relatively small
- Scales become smaller with increasing resolution
- All resolutions show secondary maxima



Impact horizontal diffusion

- Linear spectral horizontal diffusion in HARMONIE
- Same horizontal diffusion coefficients for all parameters
- Initially too strong horizontal diffusion, and different strength for different parameters
- Caused later onset of convection, larger cells and stronger outflow (French fireworks case).

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Impact SLHD



- SLHD = semi lagrangian horizontal diffusion
- SL-scheme diffusive, strength of diffusion can be chosen, the semi-Lagrangian buffer is the only place in the model where 3-D fields are available and 3D effects can be implemented.
- SLHD applied to reduce the strength of spurious small scale lows in ALADIN





(b) SL3 experiment vertical divergence

1-1









Impact microphysics

- Evaporation
- Removal of hydrometeor species
- Fall speed



Impact evaporation

- Hydrometeors often present in only a part of grid cell, especially at boundary of convective clouds
- Model does not know how large part of grid cell is occupied by hydrometeors
- Fraction of evaporation to take hydrometeor distribution into account?

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Evaporation

- Reduction of evaporation increases the temperature of unsaturated air with hydrometeors
- Cold pools warmer, weaker outflow
- Study against observations: reduction to 50-30% of original values gives best distributions of temperature and wind direction



Removal of graupel

- Strong convection only when graupel (quick parameter) present
- Removal of graupel leads to larger convective cells, lower maximum precipitation intensity
- Probably due to snow becoming more important, lower fall speed, higher water loading effect (reduction updraft speed).



Operational setup for CY36H1.1

- SLHD applied to hydrometeors
- SLHD NOT applied to humidity
- Same (small) spectral horizontal diffusion to all spectral upper-air fields

Forecasts of strong rain event in Canary Islands

Hirlam

aic_ 10m wind and precipitation 01/02/2010 00z HARMON H+ 24 Valid: 02/02/2010 00z





Implementation of the Vertical Finite Element discretization on the non-hydrostatic version



NH DYNAMICAL CORE

- Prognostic variables
 - Horizontal velocity, V
 - Vertical divergence, d
 - Temperature, T
 - non-hydrostatic pressure departure, q
- Stability in presence of orography needs a new divergence variable

$$dl = d + X = \frac{p}{mRT} \left(-g \frac{\partial w}{\partial \eta} + \nabla \varphi \cdot \frac{\partial V}{\partial \eta} \right)$$



NH DYNAMICAL CORE (cont)

- Vertical coordinate
 - Mass based, hybrid terrain following coordinate
 - Time dependent
- Time stepping: semi-implicit
 - Constant and horizontally uniform coefficients
 - Linear system using additional elastic
 reference temperature for acoustic wave terms
 - Iterative centered implicit (ICI) implemented



NH DYNAMICAL CORE (cont)

- Spatial discretization
 - Spectral in the horizontal
 - Finite differences in the vertical
- Vertical operators must verify constraints

C1: $G^* + S^* - G^*S^* - N^* = 0$

C2': $L^* \left(S^* G^* - \frac{c_p}{c_v} \left(S^* + G^* \right) \right)$ must have real negative eigenvalues.

Not easy to implement with VFE

Reading 8-10 Nov 2010 ECMWF workshop on non-hydrostatic modeling

$$m^* \equiv \frac{d\pi^*}{d\eta}$$
$$\partial^* X = \frac{\pi^*}{m^*} \frac{\partial X}{\partial \eta}$$
$$G^* X = \int_{\eta}^{1} \frac{m^*}{\pi^*} X d\eta$$
$$S^* X = \frac{1}{\pi^*} \int_{0}^{\eta} m^* X d\eta$$
$$N^* X = \frac{1}{\pi^*_s} \int_{0}^{1} m^* X d\eta$$
$$L^* X = \partial^* (\partial^* + 1) X$$

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Other prognostic variables?

- Test other prognostic variables
 - horizontal and vertical velocity, temperature, geopotential and logarithm of hydrostatic pressure at surface
- It is SHB stable
 - but instability appears when orography is present
- Constraints
 - Vertical differential and integral operators must satisfy constraints

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Other prognostic variables

 In sigma mass based vertical coordinate the prognostic equations are

$$\frac{dV}{dt} + \frac{RT}{p} \nabla p + e^{-q} \frac{\partial p}{\partial \sigma} \nabla \varphi = F$$
$$\frac{dw}{dt} + g \left(1 - e^{-q} \frac{\partial p}{\partial \sigma} \right) = F_z$$
$$\frac{dT}{dt} + \frac{RT}{c_v} D_3 = \frac{Q}{c_v}$$
$$\frac{d\varphi}{dt} - gw = 0$$
$$\frac{\partial q}{\partial t} + \int \nabla \cdot V d\sigma + (\nabla q) \cdot \int V d\sigma = 0$$

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Other prognostic variables?

 \mathbf{A}

This set of constraints must be satisfied

C1:
$$\partial^* N^* = 0$$

C2: $(1+\partial^*)S^* = 1$
C3: $[\partial^*\partial^*] = \partial^* \cdot \partial^*$
 $A^*[f(\sigma)] = \frac{1}{\sigma}\int f(\sigma')d\sigma'$

- The constraints can be fulfilled with VFE
- Strong instability when orography is

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Other vertical coordinate?

- Both mass and height based coordinates has been shown to be suitable for NH modeling.
- Why to introduce height based coordinate?
 - Is a time independent coordinate.
 - It has an easier mathematical treatment of covariant velocity than the mass based coordinate.
- Covariant derivative can be used in the discretization of differential operators.

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Other vertical coordinate?



 Covariant variables and covariant derivative makes the expression of divergence operator simpler, without the non-linear term which is unstable in presence of orography.

Covariant : div
$$V = |g|^{-\frac{1}{2}} \sum \frac{\partial}{\partial X_i} \left(|g|^{\frac{1}{2}} V^i \right)$$

Non Covariant, mass based : div $V = \nabla_{\pi} \cdot V_{H} - \rho g \frac{\partial w}{\partial \pi} + \rho \nabla_{\pi} \varphi \cdot \frac{\partial V}{\partial \pi}$

Covariant differential operators (gradient, divergence, curl and laplacian) are Reastraightforwardly found operators (gradient, Reastraightforwardly found operators (gradient, Reastraightforwardly found operators (gradient, straightforwardly foun

2D (X,Z) model



- The set of prognostic variables used are
 - Covariant **V**, In(T) and In(p)
- SHB numerical linear stability analysis shows in the 3TL case:
 - amplification factors under 1.01 in most cases.
 - stability for residual term temperature T in the range 0.5 T*< T < 1.5 T*.
- Vertical discretization:
 - VFE has been implemented and it is more accurate than VFD with almost equal stability.
- Advection:
- Covariant Eulerian and covariant semi-Lagrangian
 schemes has been applied successfully
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2D (X,Z) model



 Covariant vertical velocity variable W is the prognostic variable used in the Helmholtz structure equation of the semi-implicit solver:

Helmholtz
$$(I - \beta^2 c_*^2 (\nabla^2 + L_Z) - \beta^4 c_*^2 N_*^2 \nabla^2) W_{n+1} = RHS$$

where vertical Laplacian $L_Z = \frac{1}{H_*^2} \left(\frac{H_*}{H_T} \partial_Z \cdot \frac{H_*}{H_T} \hat{\partial}_Z + \frac{H_*}{H_T} \hat{\partial}_Z \right)$
 $c_*^2 = R_d T^* \frac{C_{pd}}{C_{vd}}; \quad H_* = \frac{R_d T^*}{g}; \quad N_*^2 = \frac{g^2}{C_{pd} T^*}$

- Robust boundary condition implementation
 - W=0 at boundaries is included in the semi-implicit solver.

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- Two layer atmosphere with two different stabilities
- Waves propagating downstream
- Good agreement with the linear analytical solution





- Warm bubble.
- Good agreement with other published results.



Janjic Z. I., Gerrity Jr. J. P., Nickovic S., 2001: An Alternative Approach to Nonhydrostatic Modeling. Monthly Weather Review Volume 129, Issue 5 pp. 1164-1178



- Linear and non linear Hydrostatic waves.
- Good agreement with the analytical solutions and other results.



Bubnová R., Hello G., Bénard P., Geleyn J.F., 1995: Integration of the Fully Elastic Equations Cast in the Hydrostatic Pressure Terrain-Following Coordinate in the Framework of the ARPEGE/Aladin NWP System, Monthly Weather Review Volume 123, pp. 515-535



- Linear and non linear Non-Hydrostatic waves.
- Good agreement with the analytical solutions and other results.







- Acoustic waves.
- It is a full elastic model and therefore can reproduce accurately acoustic waves.



Covariant SL scheme



- Eulerian and semi-Lagrangian produce similar results although semi-Lagrangian is more dispersive (due in part to the low order bilinear interpolation)
- Semi-Lagrangian (left) allows bigger time steps than eulerian (right)







Thank you for your attention