# Diagnosing Model Systematic Error for Clouds and Precipitation

#### **Richard Forbes (ECMWF)**

With Thanks to Maike Ahlgrimm, Peter Bechtold, Martin Köhler and ECMWF colleagues, and Graeme Stephens (CSU), Julien Delanoë (Univ. Reading)

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# **Cloud Validation: The issues**





Sounds easy.....



• How much of the 'error' derives from observations?



# **Cloud Validation: The problems**



• Which Physics is responsible for the error?



# Cloud Validation: The problems





1. Methodology for diagnosing errors and improving parametrizations

# **Cloud Validation: The problems**





# A strategy for cloud parametrization evaluation





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- ECMWF Global Atmospheric Model (IFS)
  - T159 (125km) Monthly and Seasonal Prediction, (+Model Testing)
  - T399 (50km) Ensemble Prediction System (EPS)
  - T511 (40km) Previous NWP
  - T799 (25km) Current Deterministic Global NWP 10 day f/c
  - T1279 (16km) NWP (soon!)
  - 62 and 91 levels in use
- Need a model with a "climate" that is robust to resolution.



- 1. Choice and formulation of microphysical processes
  - As resolution increases, the range of scales of dynamical forcing increases. Nonlinear microphysical processes can respond differently.
  - Parametrization schemes are based on time/space scale separation. As resolution increases, require change from diagnostic to prognostic variables.
- 2. Representation of sub-gridscale inhomogeneities
- 3. Numerical techniques for efficient implementation
  - Long timesteps used for computational efficiency.
  - Explicit vs implicit formulations.

# TOA Net Radiation (T511 vs T159)



-30

135 E



\*\*Top solar radiation ey0u-exw9 200009 nmon=12 nens=4 Diff: 1.52 Stdev: 6.616



30 24 18 12 6 -6 -12 30.8 -18 -24 60 ° S

\*\*Top thermal radiation ([W m\*\*-2 s)] exw9 200009 nmon=12 nens=4 Mean: -241.7

TOA Net SW

TOA Net LW

45.°E

dd+E

135°W

det = VA

\*\*Top thermal radiation ey0u-exw9 200009 nmon=12 nens=4 Diff: -1.693 Stdev: 3.577



# Sensitivity to resolution Total Cloud Cover (T511 - T159)





\*\*Total cloud cover ey0u-exw9 200009 nmon=12 nens=4 Diff: -0.01878 Stdev: 0.04294



# Sensitivity to resolution Ice Water Content (T511 - T159)





## Vertical resolution sensitivity Ice Sedimentation

- Forward-in-time upstream implicit solver.
- Small sensitivity to vertical resolution / timestep.

- At earlier cycles, the problem was **much** worse.
- Caused by the "exact" solver of Tiedtke in combination with ice sedimentation acting as a proxy for autoconversion if ice fell into clear sky.

#### **Adrian Tompkins**





# A strategy for cloud parametrization evaluation





- Step 1 : identify major problem areas
- Step 2 : identify major problem regimes
- Step 3 : identify typical
- Step 4 : identify detailed
  - parametrization



# Isolating the source of error



- We want to isolate the sources of error. Focus on particular phenomena/regimes, e.g.
  - Extra tropical cyclones
  - Stratocumulus regions
- An individual case may not be conclusive: Is it typical?
- On the other hand general statistics may swamp this kind of system
- Can use compositing technique (extra-tropical cyclones)
- Focus on distinct regimes if can isolate (SCu, Trade Cumulus)

## Composites – Extra-tropical cyclones







ECMWF clouds



Overlay about 1000 cyclones, defined about a location of maximum optical thickness

Plot predominant cloud types by looking at anomalies from 5-day average

- High Clouds too thin
- Low clouds too thick

High tops=Red, Mid tops=Yellow, Low tops=Blue

Klein and Jakob, 1999, MWR

### Model Climate: Regime dependent error?

CY30R1 ·S Mean err -6.7 50N-S rms 17 CY31R1 Difference lean err -7.58 50N-S rms 15.3 CY32R2 Wm<sup>\*2</sup> 30°1 **ERA-I** cycle 30° (almost) Difference 3 Mean err -11.4 50N-S rms 17.7 CY32R3 Wm<sup>-2</sup> 60°N McICA SW radiation 30°N Convective param. and vertical diffusion 135°W 45°E 90°E 135°E 90°W

Diagnosing Cloud Error, Forbes ECMWF Seminar 2009

TOA net SW radiation vs. CERES: Maike Too much reflectance from TCu, not enough from SA hlgrimm

# Does the model have "correct" trade cumulus cloudiness?



Three aspects:

Cloud amount when present (AWP)

helps identify cloud type



Cloud frequency of occurrence (FOO)

with AWP gives total cloud cover

Radiative properties

radiative balance ultimately drives the system

Maike Ahlgrimm



Identify cloud samples as:

- with less than 50% cloud fraction
- cloud top below 4km
- over ocean
- between 30S and 30N

#### Maike Ahlgrimm

#### TCu frequency of occurrence (Foigesig Cloud Error, Forbes ECMWF Seminar 2009



# Cloud amount when present (AWP)





Cloud fraction is subject to representativity error. Observations have not been corrected!

Most of the additional TCu samples have very small cloud fractions

Maike Ahlgrimm

# Cloud top height

Diagnosing Cloud Error, Forbes Skewed distrived ar 2009 with low peak: Majority of TCu clouds are very shallow, few

Model clouds have higher cloud tops than observed





# A strategy for cloud parametrization evaluation





C.Jakob



#### GCSS: Validation of CRMs

Redelsperger et al QJRMS 2000 SQUALL LINE SIMULATIONS





#### GCSS: Comparison of many SCMs with a CRM CCM / = **Diagnosing Cloud Error, Forbes** Bechtold et al QJRMS 2000 SQUALL LINE SIMULATIONS ECMWF Seminar 2009



Total mass flux (kg/m²/s)

column models (see Tables 1 and 2 for explanations of the acronyms).

# Summary



- Long term climatologies:
  - Climate systematic errors we want to improve the basic state/climatology of the model
  - But which physics is responsible for the errors? Non-linear interactions.
  - Long term response vs. transient response.
  - We want to remove sensitivity to resolution = the parametrization problem!
- Isolating regimes: Composites and focus on geographical regions.
- Case studies
  - Detailed studies with Single Column Models, Cloud Resolving Models, NWP models
  - Easier to explore parameter space.
  - Are they representative? Do changes translate into global skill?

# 2. Comparing model and obs: Uncertainty and limitations

# Cloud Validation: The problems



Diagnosing Cloud Error, Forbes ECMWF Seminar 2009

# What is a cloud ?







# What is a cloud ?

- Different observational instruments will detect different characteristics of clouds.
- A cloud from observations may be different to the representation in models

- Understanding the limitations of different instruments
- Benefit of observations from different sources
- Comparing like-with-like (physical quantity, resolution)

## Verification Annual average T159 Ice Water Path vs. Obs



Widely varying estimates of IWP from different satellite datasets!



#### New 5 prognostic cloud microphysics Ice vs. Snow



#### Model Ice Water Path (IWP) (1 year climate)

IWP from prognostic cloud ice variable



#### IWP from cloud ice + precipitating snow



#### Observed Ice Water Path (IWP)

CloudSat 1 year climatology



#### CloudSat/CALIPSO/EarthCare Meeting- 2008



#### A-Train: 28th April 2006



CloudSat: Cloud profiler radar 94GHz CALIPSO: Cloud profiler lidar 532, 1064nm + Infra Red Imager AQUA: radiometers MODIS, AIRS, CERES, AMSR-E

#### **Global coverage:**

Radar :2.5 km along track X 1.2 km across track / 500m =>250m Lidar: 333 m / 30 m Our merged product: CloudSat footprint/vertical resolution 60m

#### Julien Delanoë/Robin Hogan

# Example of mid-Pacific convection



**MODIS 11 micron channel** 







### Why combine radar, lidar and radiometers?



- Radar  $Z \propto D^6$ , lidar  $\beta' \propto D^2$  so the combination provides particle size
- Lidar: sensitive to particle concentration, can be extinguished
- Radar: very sensitive to the particle size, not very sensitivity to liquid clouds and small ice particles



Radiances ensure that the retrieved profiles can be used for radiative transfer studies -Single channel: information on extinction near cloud top -Pair of channels: ice particle size information near cloud top

We use "unified" variational scheme to retrieve ice cloud properties, thin and thick ice clouds Delanoë and Hogan 2008, JGR (doi:10.1029/2007JD009000)

#### Julien Delanoë/Robin Hogan

### Formulation of the retrieval scheme



We know the **observations** (instrument measurements) and we would like to know **cloud properties** : visible extinction, Ice water content, effective radius...

lacksquare

Observation vector

State vector (which we want to retrieve)

- Elements may be missing





Radar reflectivity x = 1 factor profile

Infrared radiance Radiance difference  $\mathbf{x} = \frac{\ln \alpha_1}{\ln \alpha_p}$ 

 $\ln N_{0n}^{\rm ice}$ 

Ice visible extinction coefficient profile

Ice normalized number conc. profile

Extinction/backscatter ratio for ice

Iterative process: compare predicted observations and measurements, with an **a-priori** and **measurement errors** as a constraint

#### Julien Delanoë/Robin Hogan

# Combining radar and lidar...



#### Global-mean cloud fraction



#### Frequency of occurrence of IWC vs temperature

-80

-70

-60

-50

-40

-30

-20

-10

0

-8

-80

-70

-60

-50

-40

-30

-20

-10

0

-8

-7

-7

S

T [deg (



Radar+lidar



IWC increases with temperature:

- but spread over 2 to 3 orders of magnitude at low temperatures
  - reach 5 orders of magnitude close to 0° C

Advantage of the algorithm: Deep ice clouds: radar Thin ice clouds: lidar

When radar and lidar work T [deg C] well together very good confidence in the retrievals

> Obvious  $\Rightarrow$ complementarity radar-lidar

### **CloudSat/CALIPSO Model Verification**

**GLOBAL Ice Water Content vs. T distributions** 



**Diagnosing Cloud Error, Forbes** 

ECMWF Seminar 2009

When comparing a model with observations, we need to compare like-with-like



# Spatial resolution mis-match



- Need to address mismatch in spatial scales in model (50 km) and obs (1 km)
- Sub-grid variability is predicted by the IFS model in terms of a cloud fraction and assumes a vertical overlap.
- Either:
  - (1) Average obs to model representative spatial scale
  - (2) Statistically represent model sub-gridscale variability using a Monte-Carlo multiindependent column approach.





# MODEL to OBSERVATION



#### Radar Reflectivity: Cross-section through a mid-latitude front











# **Radar Reflectivity Statistics**





# **Radar Reflectivity Statistics**



STATISTICS: Frequency of occurrence (Radar Reflectivity vs. Height) Tropics over ocean 30S to 30N for February 2007





# Summary



- Limitations and uncertainty:
  - Observations have limitations, provide a partial picture.
  - We need to know the error characteristics (including systematic errors). But we don't always know this!
- Synergy:
  - Different observation sources have different strengths and weaknesses Make the most of this complementary information (e.g. CloudSat, CALIPSO, MODIS)
- Need to compare like-with-like
  - Compare in "model-space", or "obs space"
  - Assumptions required for both
  - Spatial and temporal resolution differences
- Diagnose model problems from different angles
  - Retrieval of model variables from observations ("obs-to-model"), e.g. IWC
  - Forward modelling of observed variables ("model-to-obs"), e.g. Z
  - Different approaches help to diagnose model problems

# 3. Understanding Physical Processes





#### **Temperature vs. Relative Humidity**





#### Precipitation validation with CloudSat (in collaboration with Graeme Stephens)





- Model overestimates frequency of low precipitation rates (< 1 mm/hr)
- Model underestimates frequency of high precipitation rates (> 5 mm/hr?) (but representativity?)
- Still an issue of uncertainty in the obs work in progress.





- Radar and lidar used to derive drizzle rate below stratocumulus
- Important for cloud lifetime in climate models

- Met Office uses Marshall-Palmer distribution for all rain
  - Observations show that this tends to *overestimate* drop size in the lower rain rates
- Most models (e.g. ECMWF) have no explicit raindrop size distribution

From Robin Hogan



# Drizzle 1-year comparison with models



- ECMWF, Met Office and Meteo-France overestimate drizzle rate
  - Problem with auto-conversion and/or accretion rates?
- Larger drops in model fall faster so too many reach surface rather than evaporating: drying effect on boundary layer?



From Ewan O'Connor, Robin Hogan (Reading Univ.)

# Warm-rain processes

 Autoconversion – conversion of cloud droplets to raindrops

- Accretion sweep out of cloud droplets by rain
- Evaporation below cloud base







## Process-validation with CloudSat: Autoconversion/accretion





From Graeme Stephens







# Summary



#### 1. Methodology

**Different approaches** to verification (climate statistics, case studies, composites), resolution

2. Comparing model and obs: Uncertainty and limitations Need to understand the limitations of observational data. Different techniques (model-to-obs, obs-to-model) and a range of observations are required to validate and improve cloud parametrizations.

#### 3. Processes

Can observations be used to test model's **physical relationships** between variables and to understand physical processes.