



Progress achieved on assimilation of satellite data in NWP over the last 30 years

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Progress achieved on assimilation of satellite data in NWP over the last 30 years



sub-title:

Ancient Developments in the Use of Satellite Observations in NWP

Structure of talk

- Satellite soundings (passive IR/MW soundings of temp/humidity profiles)
 - Early instruments
 - Assimilation experience: 1970s and 1980s
 - Problems with assimilation of retrievals
 - Direct assimilation of radiances: 1990s
- Atmospheric Motion Vectors (AMVs)
- Scatterometry
 - Early instruments
 - Early assimilation experience
- More recent advances
 - TOVS → ATOVS, AIRS and IASI, other data types
 - Radio occultation
- Strategies for various data types

Weather satellites – early milestones



TIROS-1	1960	1st satellite giving images of Earth
NIMBUS-1	1964	1st meteorological research satellite
ATS-1	1966	1st geostationary weather satellite
ESSA-1	1966	1st operational weather satellite
NIMBUS-3	1969	1st temperature sounders
ITOS-1	1970	1st APT system – improved imagery
NOAA-2	1972	1st operational temperature sounder
SMS-1	1974	1st USA operational geostationary satellite
GMS-1	1977	1st Japanese operational geostationary satellite
Meteosat-1	1977	1st European operational geostationary satellite
TIROS-N	1978	New generation of operational polar satellites
FGGE	1979	First GARP Global Experiment

Satellite soundings

- passive infra-red/microwave soundings of temperature/humidity profiles

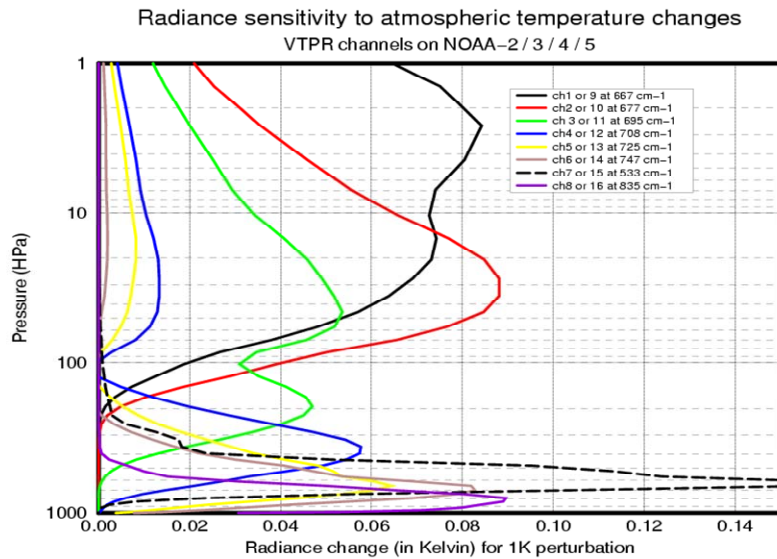
Nimbus series – temperature/humidity sounders

- Nimbus-3 1969-70 **SIRS**, IRIS
- Nimbus-4 1970-71 **SIRS**, IRIS, SCR
- Nimbus-5 1972 ITPR, SCR
- Nimbus-6 1975 **HIRS**, **SCAMS**, PMR, LRIR
- Nimbus-7 1978-94? LIMS, SAMS

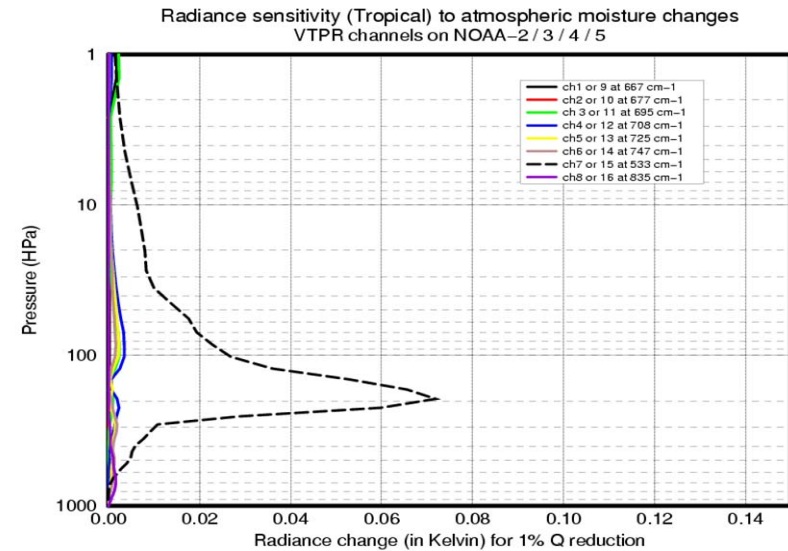
NOAA series – temperature/humidity sounders

- NOAA 2-5 1972-79 **VTPR**
- TIROS-N 1978-80 **TOVS = HIRS, MSU, SSU**
- NOAA-6/14 1979- **TOVS = HIRS, MSU, SSU**
- NOAA-15+ 1998- **ATOVS = AMSU-A, AMSU-B, HIRS**

VTPR Radiance Sensitivity



temperature



humidity

TOVS – weighting functions

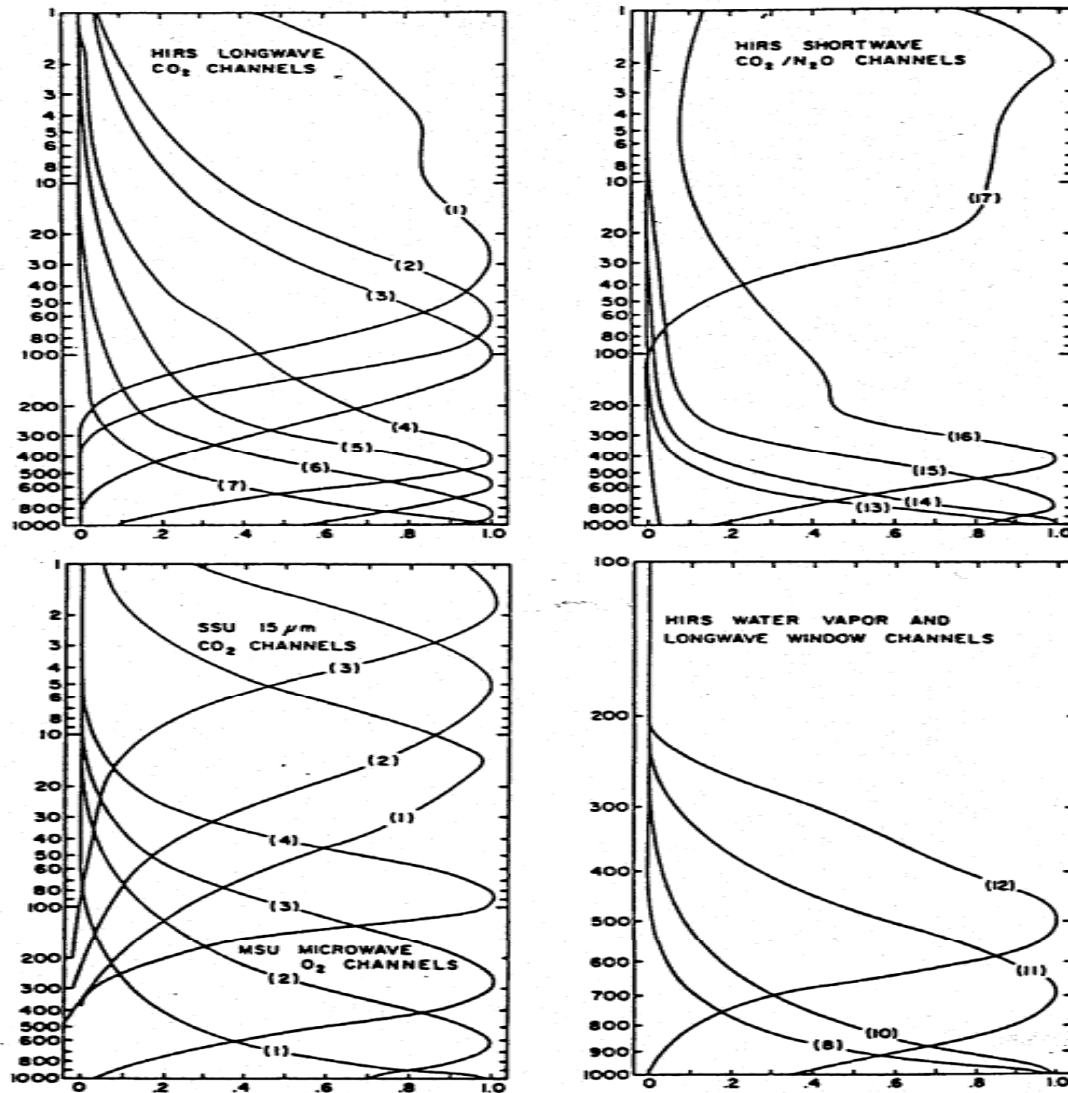
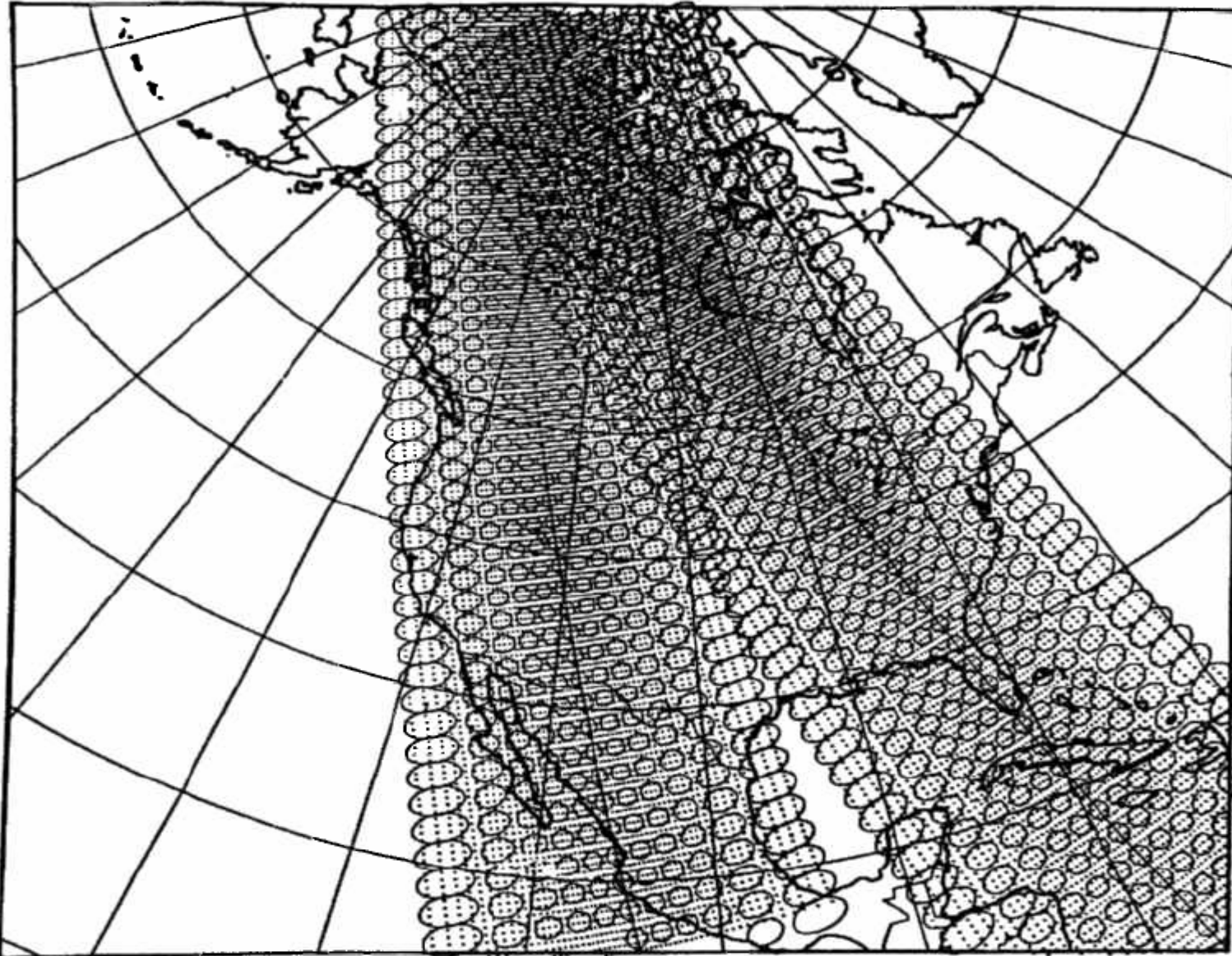


Fig. 3 TOVS normalised weighting functions (from Smith et al., 1979).



HIRS and
MSU scan
patterns

Australian experience

See W.Bourke, “History of NWP in Australia – 1970 to the present”, BMRC Workshop, October 2004

- Importance of satellite cloud imagery interpretation for analysis of surface pressure (PAOBs) and 1000-500 hPa thickness in SH.
- From 1972, Kelly used NOAA-2,3,4 VTPR data – retrievals from cloud-cleared radiances.
- 1976, Kelly demonstrated within a continuous data assimilation system benefits of assimilating VTPR and PAOBs.
- Kelly, Mills and Smith (BAMS, 59, 393-405, 1978) “Impact of Nimbus-6 temperature soundings on Australian regional forecasts”:
 - 14 days assimilation. Average improvement of >5 skill scores points on 24h geopotential forecasts (surface → 200 hPa)

In this story, this chap Kelly appears everywhere!

UK experience

Atkins and Jones, “An experiment to determine the value SIRS data in numerical forecasting”, Meteorol Mag, 104, 125-142, 1975.

- SIRS data impact study
- Used operationally, at discretion of Chief Forecaster.

Summary paper:

Ohring G, "Impact of satellite temperature soundings on weather forecasts"
(BAMS, 60, 1142-1147, 1979).

Summarised results from several OSEs

- Desmarais et al (1978) VTPR + Nimbus 6
- Halem et al (1978) VTPR + Nimbus 6
- Bonner et al (1976) VTPR
- Atkins and Jones (1975) SIRS
- Druyan et al (1978) VTPR
- Kelly (1977) VTPR
- Kelly et al (1978) Nimbus 6

Summary paper:

Ohring G, “Impact of satellite temperature soundings on weather forecasts” (BAMS, 60, 1142-1147, 1979).

Summary:

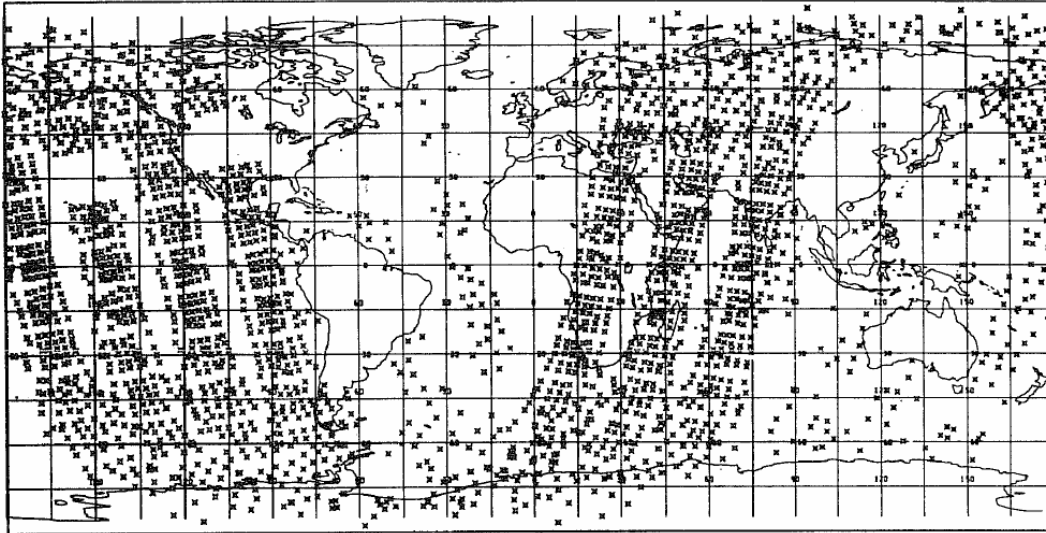
- “on average, a small improvement in numerical forecasts”
- “beneficial but modest impacts”
- “hesitate to claim that satellite data changed a poor forecast to an accurate one”
- Greater improvements in forecasts in S Hem.

Problems:

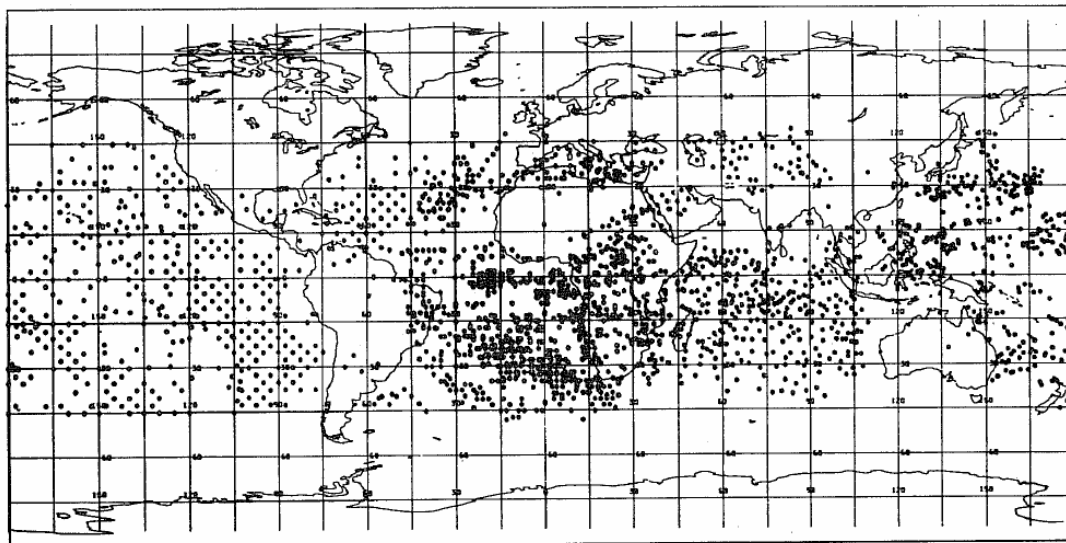
- Differences between retrievals and collocated radiosondes of 2-3 deg
- Analyses using satellite data have lower eddy potential energy
- Satellite soundings not point observations – have their own error characteristic – improved analysis schemes may enhance impact
- Improvements in retrieval methods likely - but basic problem is poor vertical resolution – “the statistical/climatological nature of retrieval techniques may suppress horizontal structure”

FGGE – satellite data coverage

POSITION OF SATEM OBSERVATIONS
0Z ON 27/2/1979



POSITION OF SATOB OBSERVATIONS
0Z ON 27/2/1979



FGGE:

First GARP Global
Experiment

(GARP = Global
Atmospheric Research
Programme)

General observational period:
01.12.1978 - 30.11.1979

Special observational
periods:
05.01.1979 - 05.03.1979
01.05.1979 - 30.06.1979

Halem M, E Kalnay, W E Baker and R Atlas,

“An assessment of the FGGE Satellite Observing System during SOP-1”
BAMS, 63, 407-426, 1982

- OSEs for several obs types
- 6-hour forecast errors reduced downstream of data sparse areas by including satellite observations
- over N.America and Europe, small improvements in forecast skill
- over Australia, positive impact of satellite data is much larger

Exeter Workshop 1982. Report of “JSC Study Conference on Observing System Experiments” (Gilchrist A, 1982).

From the summary:

- 4 centres, 11 experiments, 85 forecast-days
- 3 periods: SOP-1, SOP-2, Nov 79 (2 NOAA satellites)
- **ECMWF**: NOSAT: “useful predictability reduced from 5.5 to 4.5 days in NH and from 5 to 3 days in SH
- **GLAS**: NOSAT: Large impact over S.America and Australia. Smaller but +ve impact over N.America and Europe
- **ANMRC**: NO-SATEM: Substantial +ve impact in SH
- **GLAS**: NO-SATEM: +ve impact over Australia. Europe and N.America, less impact and variable
- **NMC**: NO-SATEM: +ve impact on one cycle at T+3.5 over E.USA
- **ECMWF**: space-based only. “surprisingly good skill at T+4”, SH: small differences

ECMWF Seminar 1984. “Data Assimilation and observing system experiments, with particular emphasis on FGGE”.

Summary:

- Accuracy of satellite temperature soundings ... 2-3 deg below 850 hPa, 1.5-2 deg above ... satisfactorily assimilated ... important role in analysing large scale weather systems at high and mid latitudes, in particular in SH
- “(satellite) atmospheric soundings ... are an essential element of the GOS”
- Uppala et al
 - AMVs important for analysis of tropics
 - SATEMs of paramount importance for extra-tropical analysis over ocean areas

Kelly and Pailleux(1988). "Use of satellite vertical sounder data in the ECMWF analysis system". ECMWF Tech Mem 143.

- Layering of retrievals:
 - Change from 14 layers: 1000-850, 850-700, 700-500, 500-400, 400-300, 300-250, 250-200, 200-150, 150-10, 100-70, 70-50, 50-30, 30-20, 20-10 hPa
 - To 11 layers in 1985,
 - To 7 layers in 1987: 1000-700, 700-500, 500-300, 300-100, 100-50, 50-30, 30-10
- SH: +ve impact, NH: mixed
- QC problems (cloud and rain)
- Improvements in stratosphere
- Reduced impact in NH compared with Uppala et al (1984)

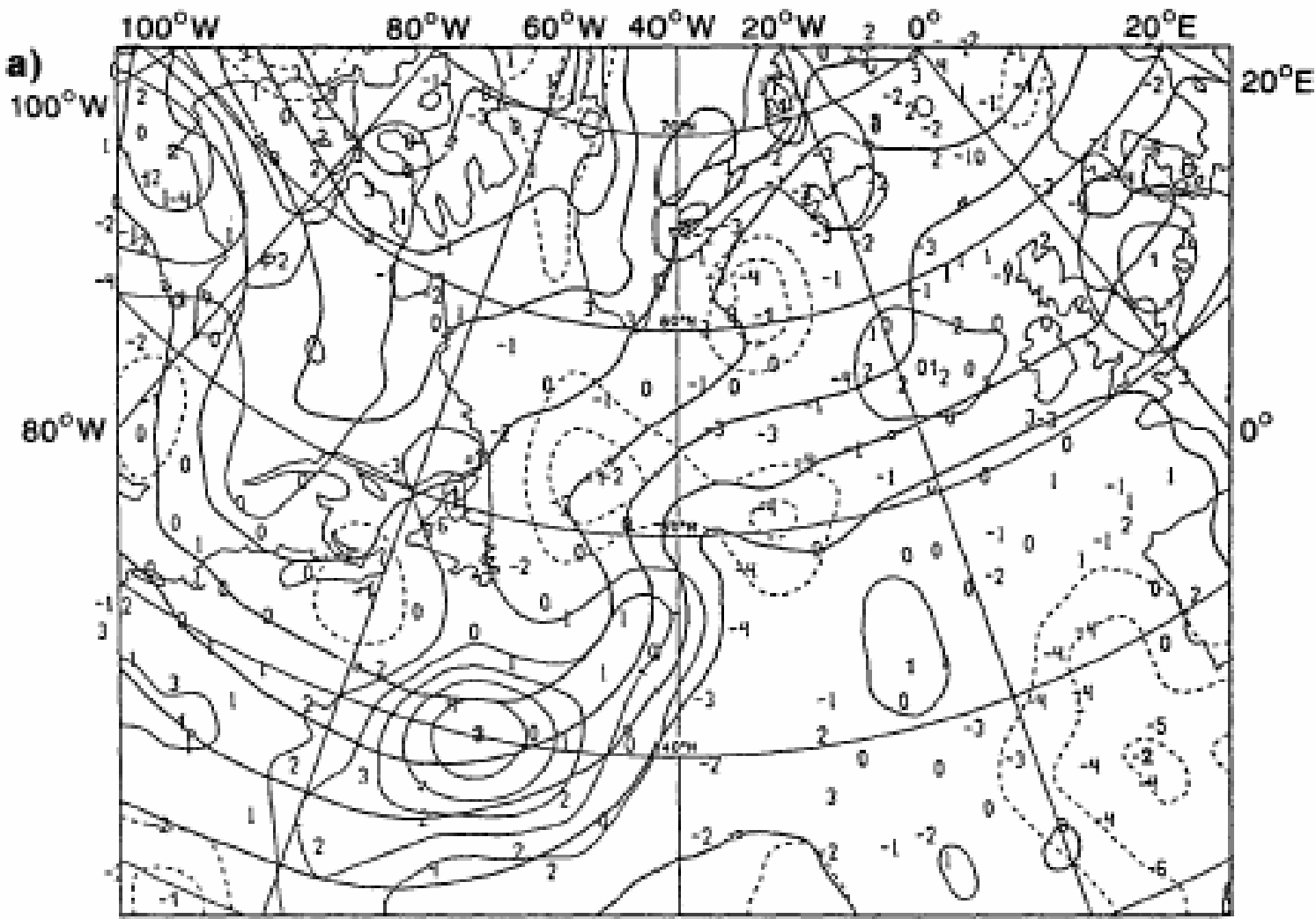
Andersson et al. “Global observing system experiments on operational statistical retrievals of satellite sounding data”, MWR, 119, 1851-1864 (1991)

- The neutral impact of SATEMs with the 1987 system gave way to a negative impact in the 1988 system. “In the present study the overall impact of SATEM data in the NH is negative”.
- Synoptically correlated biases

Kelly et al. “Quality control of operational physical retrievals of satellite soundings”, MWR, 119, 1866-1880 (1991)

- “the new physical retrievals have much the same problems of bias and noise that were noted with the statistical retrievals”
- Improved QC to mitigate the worst problems

Late 1980s: problems - synoptically correlated biases



From Andersson et al (1991). Analysis increments and background, 1000-700 hPa

Problem No.1 - RADIOSONDES

Suomi's 11th commandment:
"Thou shalt not worship the radiosonde"

- early NWP systems designed to make use of sondes
- satellite sounders and sondes have opposite strengths and weaknesses
- treating satellite soundings as "poor-quality sondes" is flawed

The history and future of data assimilation (1)



... backwards ... and in 2 slides

Bayesian:

- What is the probability of atmospheric state, x , given observations, y^o ?
- Evaluate: $P(x|y^o) = P(y^o|x).P(x)/P(y^o)$

Variational (VAR):

- What is the most probable atmospheric state, x , given observations, y^o ?
- To maximise $P(x|y^o)$,
 - **maximise: $\ln\{P(x|y^o)\} = \ln\{P(y^o|x)\} + \ln\{P(x)\} + \text{constant}$**
- If PDFs are Gaussian, then minimise a PENALTY FUNCTION,
 - **$J[x] = \frac{1}{2} (x-x^b)^T B^{-1} (x-x^b) + \frac{1}{2} (y^o-H[x])^T (E+F)^{-1} (y^o-H[x])$**

x^b : background

B : background error covariance

$H[x]$: observation operator

E, F : error covariances of observations and observation operator

Optimal Interpolation (OI)

- Linearising the VAR problem →
 - $x^a = x^b + K \cdot (y^o - H[x])$
 - where $K = B \cdot H^T \cdot (H \cdot B \cdot H^T + E + F)^{-1}$
- H is the Jacobian of the observation operator $H[x]$

Empirical

- $x^a = x^b + K \cdot (y^o - H[x])$
 - but with K as empirically-derived weights
-

Key issues for satellite soundings

- VAR provides method on handling large numbers of observations
- ... linked to analysis variables in a non-linear way

- Linearized retrieval equation: $x^a - x^b = K.(y^o - H[x^b])$
- Linearized forward equation: $y^o - H[x^b] = H.(x - x^b) + \varepsilon$
- Combine: $x^a - x^b = K.H.(x - x^b) + K.\varepsilon$

or

$$x^a - x^t = (I - K.H).(x^b - x^t) + K.\varepsilon$$

retrieval background measurement
error error error

where x^t denotes truth, I = unit matrix, $H = \nabla_x H[x]$

- This equation shows why assimilating retrieved temperature/humidity profiles into NWP models is more problematic than assimilating radiances directly

Variational equations: for 1D-Var, 3D-Var, 4D-Var

Minimize:

$$\mathcal{J}[x] = \frac{1}{2} (x-x^b)^T B^{-1} (x-x^b) + \frac{1}{2} (y^o-H[x])^T (E+F)^{-1} (y^o-H[x])$$

where x contains the NWP model state
 x^b is background estimate of x (short-range forecast)
 B is its error covariance,
 y^o is vector of measurements
 $H[...]$ is “observation operator” or “forward model”,
mapping state x into “measurement space”
 E is error covariance of measurements,
 F is error covariance of forward model.

$$\nabla_x \mathcal{J}[x]^T = B^{-1} (x-x^b) - \nabla_x H[x]^T (E+F)^{-1} (y^o-H(x)) = 0$$

TOVS in NWP via 1D-Var

- [Eyre et al](#), QJRMS, 119, 1427-1463 (1993) "Assimilation of TOVS radiance information through one-dimensional variational analysis"
- main advance over assimilation of SATEMs: 1D-Var produces no analysis increments when measured radiances agree with forecast radiances
- still needs care over assimilation because of use of forecast background in 1D-Var retrieval
 - operational ECMWF, June 1992

TOVS in 3D-Var

- [Derber and Wu](#), MWR, 126, 2287-2299 (1997). "The use of TOVS cloud-cleared radiances in the NCEP SSI analysis system".
 - operational at NCEP, October 1995
- [Andersson et al](#), QJRMS, 120, 627-653 (1994). "Use of cloud-cleared radiances in three/four-dimensional variational data assimilation"
 - operational at ECMWF, January 1996

TOVS in 4D-Var

- Operational at ECMWF, November 1997

Atmospheric motion vectors

- winds derived by tracking features in imagery

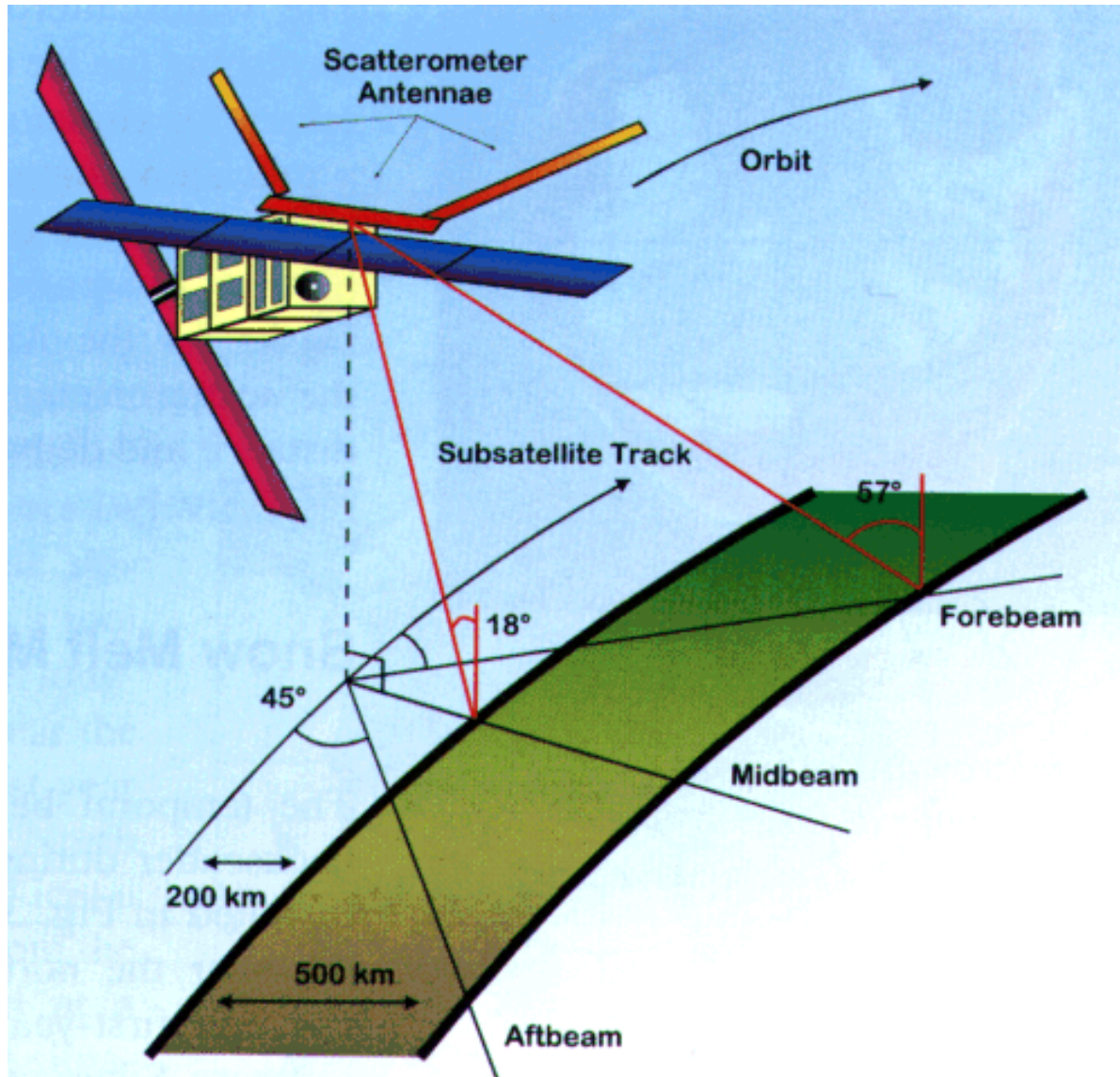
Scatterometry

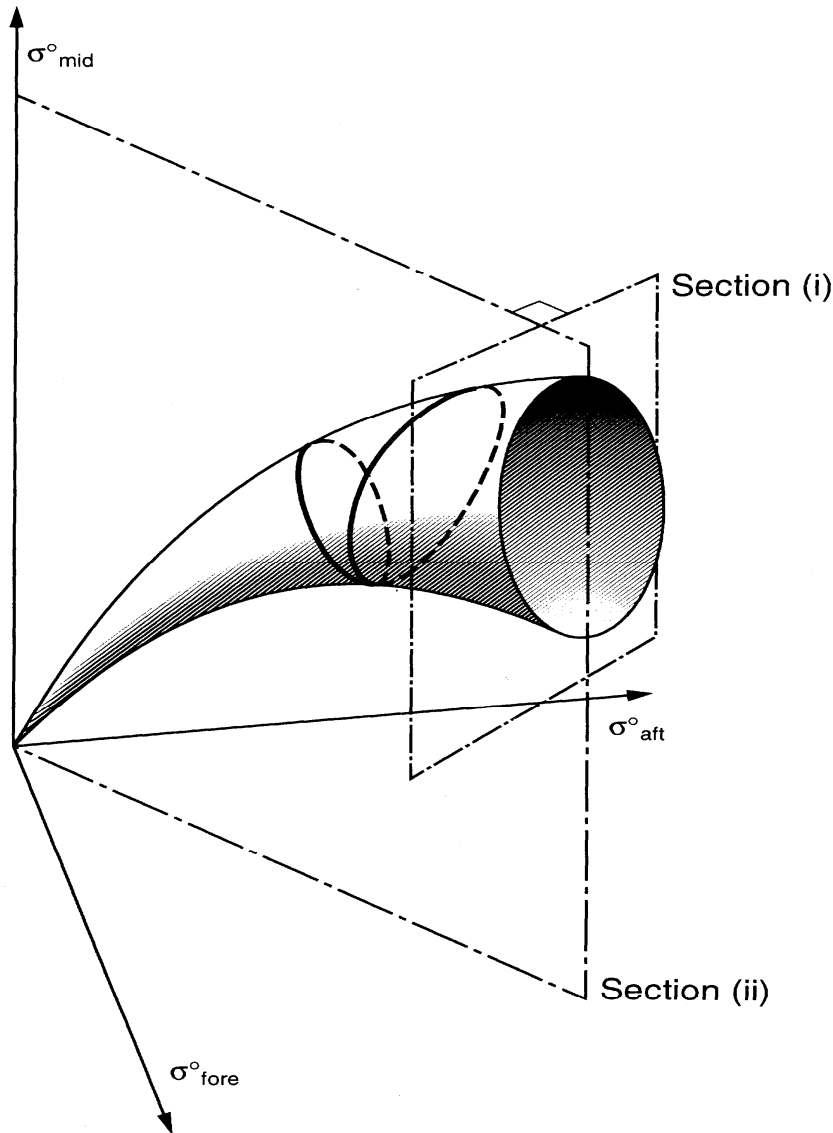
Scatterometry: satellites and instruments



year	satellite	instrument	freq GHz	views	res km	swath km
73-74	Skylab	MRSA	13.9 (Ku-band)	1*	15	185
78	SEASAT	SASS	14.6 (Ku-band)	2	25	1000
91-00	ERS-1	AMI	5.3 (C-band)	3	25	500
95-	ERS-2	AMI	5.3 (C-band)	3	25	500
96-97	ADEOS-I	NSCAT	14.0 (Ku-band)	3	50	1000
99-	Quikscat	Seawinds	13.4 (Ku-band)	4	25	1800
06-	Metop	ASCAT	5.3 (C-band)	3	25	1000

* dual polarisation





ERS scatterometer:

the measurement cone
in σ^0 -space

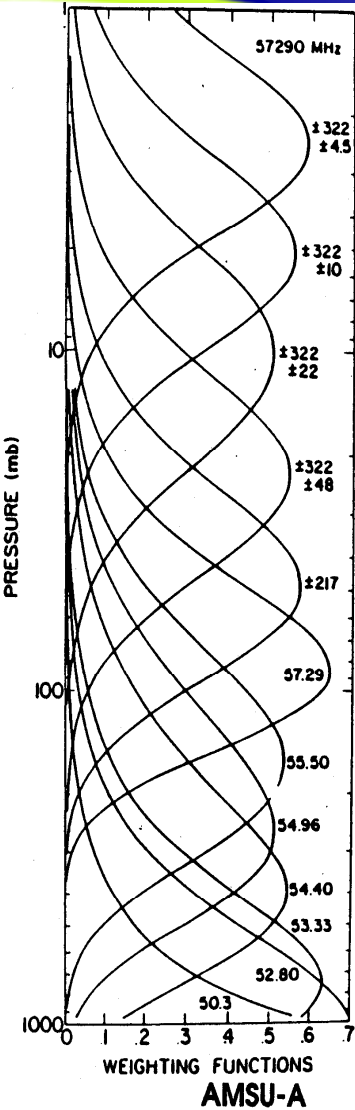
Wind speed increases along
the cone

Wind direction changes
through 360° for twice around
the cone

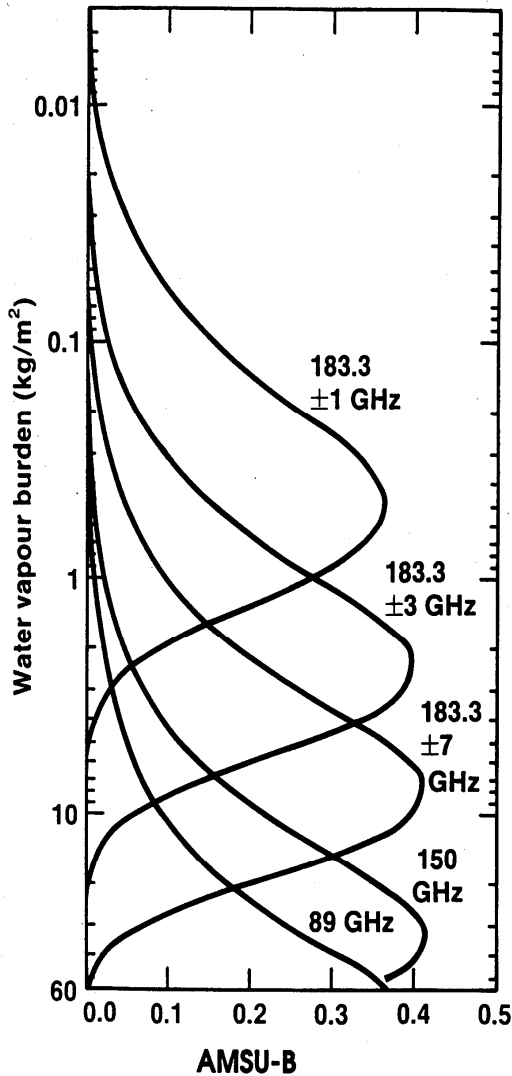
- Baker et al (JGR, 89, 4927-, 1984) SEASAT
 - Impact negligible in NH (~2% in skill score in PMSL). Impact +ve in SH removed when VTPR included. (Low-resolution model with no PBL scheme.)
- Yu and McPherson (MWR, 112, 368-, 1984) SEASAT
 - Significant impact in SH, but not possible to assess if impact is positive.
- Andersson et al (JGR, 96, 2653-, 1991) SEASAT
 - Neutral
- Stoffelen and Cats (MWR, 119, 2794-, 1991) SEASAT
 - LAM, QE-2 storm. **Positive impact.**
- Hoffman (JGR, 98, 10233-, 1993) ERS-1
 - Neutral
- Breivik et al (DNMI Tech Rep 104, 1993) ERS-1
 - Norwegian LAM. **Small positive impact.**
- Bell (Proc 2nd ERS-1 Symp, 1994) ERS-1
 - **Positive in SH at T+120**
- Stoffelen and Anderson (QJ, 123, 491-, 1997) ERS-1
 - **Positive in short-range.**
 - Operational at ECMWF ? (with 3D-Var, January 1996?)

More recent advances

More recent advances: TOVS → ATOVS



AMSU-A



AMSU-B

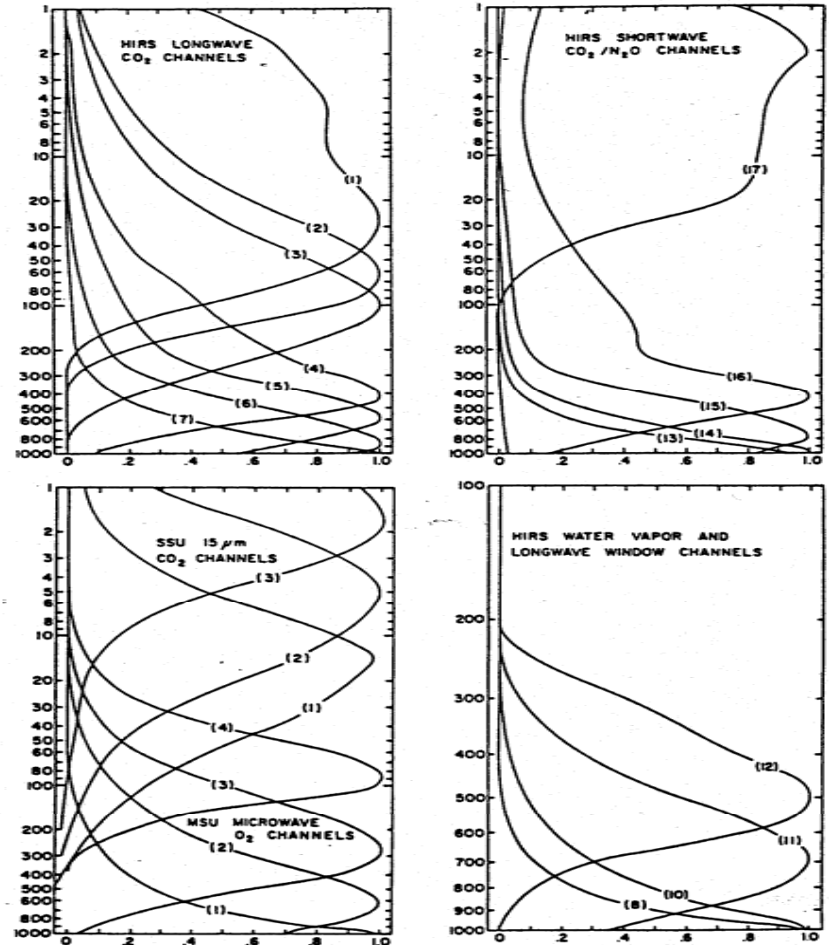
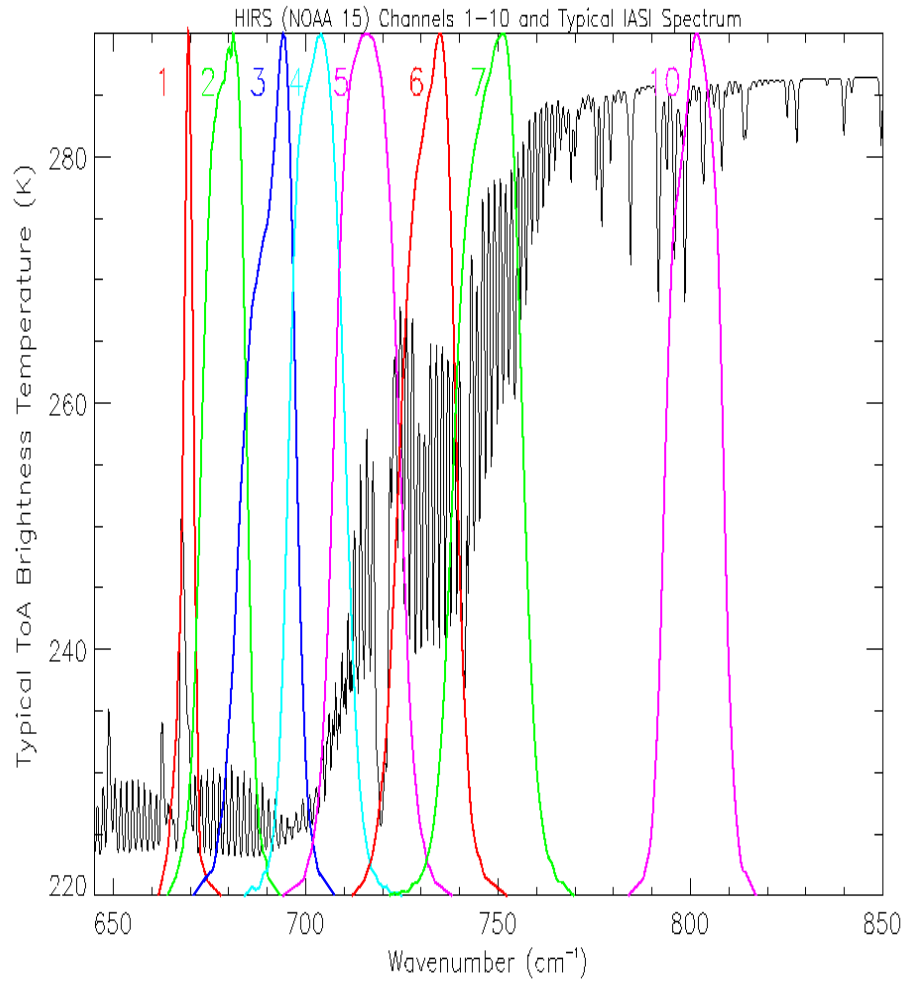


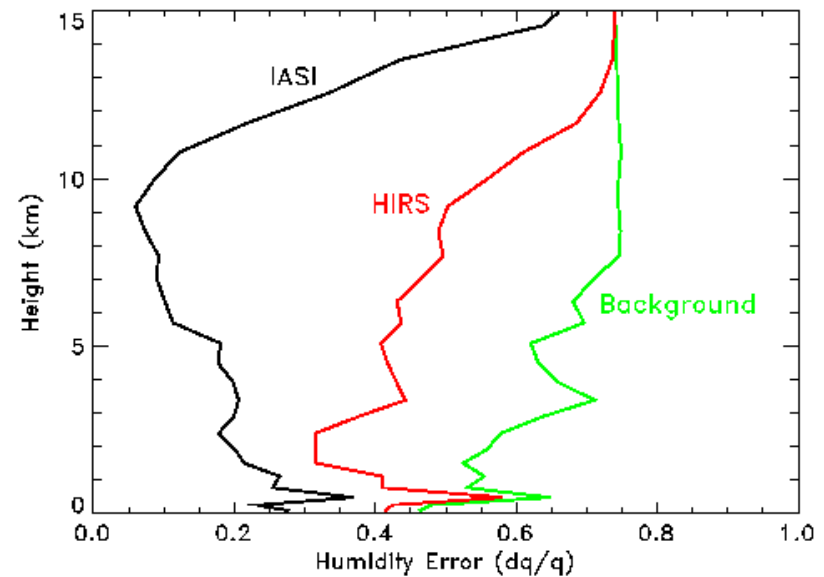
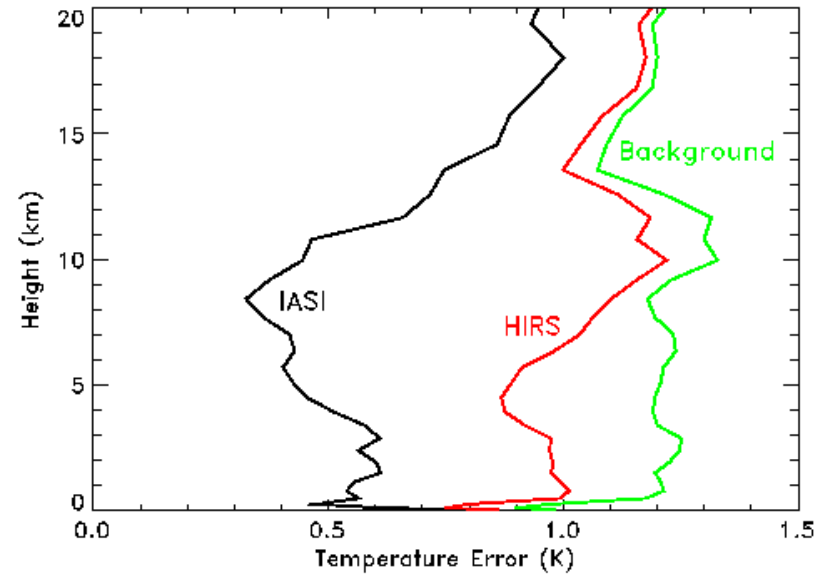
Fig. 3 TOVS normalized weighting functions (from Smith et al., 1979).

TOVS = HIRS + MSU + SSU

More recent advances: AIRS and IASI



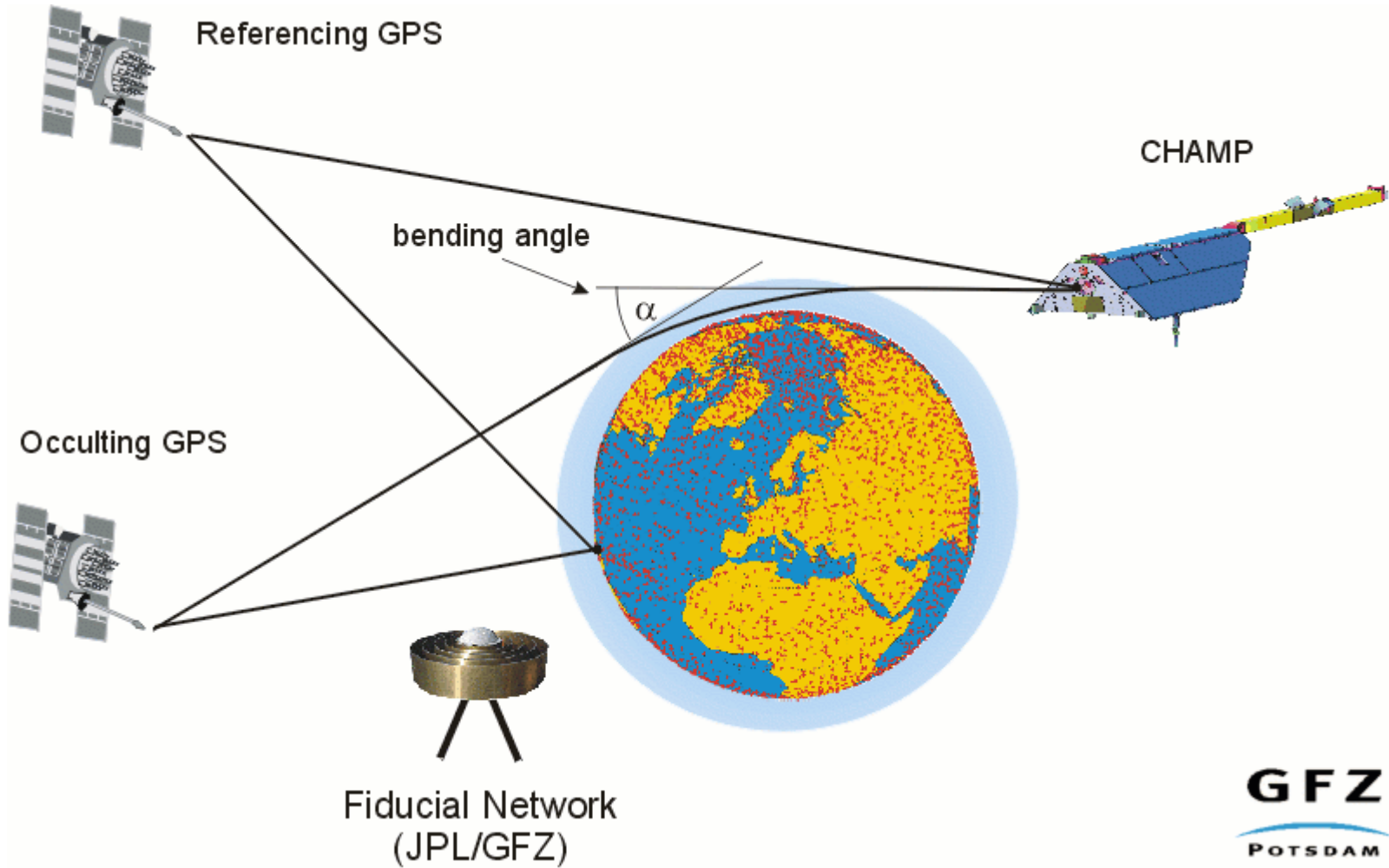
IASI v. HIRS



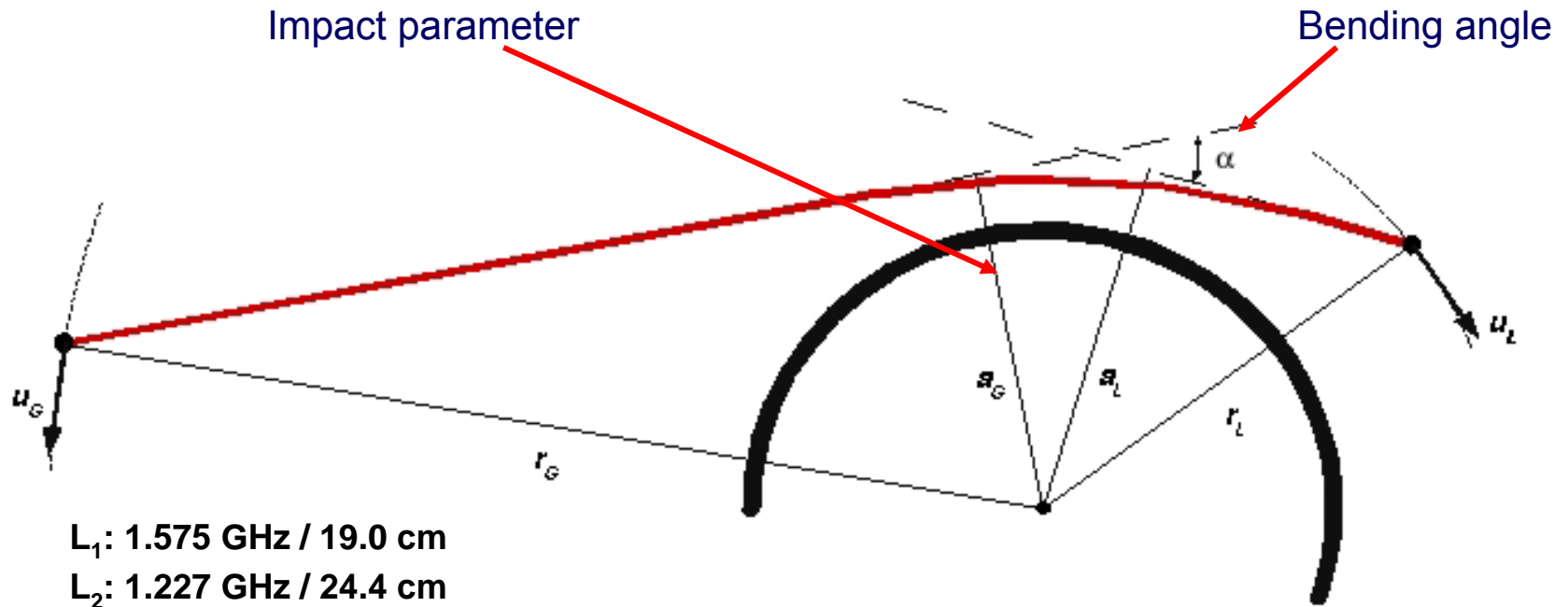
Other satellite data now assimilated in NWP:

- SSMI MW imagery (for surface wind, water vapour, cloud water)
- SSMI cloud-affected radiances (for precipitation)
- geo WV radiances
- geo retrieved cloud
- ozone (SBUV, SCIAMACHY)
- MIPAS limb radiances
- SSMIS MW sounder radiances
- GPS-WV (satellite-to-ground)
- **GPS-RO (satellite-to-satellite)**

Radio occultation: the technique (1)



Radio occultation: the technique (2)



$$\ln(n(a)) = \frac{1}{\pi} \int_a^\infty \frac{\alpha(a')}{\sqrt{a'^2 - a^2}} da'$$

Refractive index

$$N = (n - 1) \times 10^6 = \kappa_1 \frac{p}{T} + \kappa_2 \frac{e}{T^2}$$

Refractivity

Refractivity gradients caused by gradients in:

- density (pressure and temperature)
- water vapour
- electron density
- (liquid water)

$$N = \kappa_1 p/T + \kappa_2 e/T^2 + \kappa_3 n_e/f^2 + \kappa_4 W$$

“dry” “moist” ionosphere “scattering”

N = refractivity = $(n - 1) \times 10^6$; n = refractive index

p = pressure

T = temperature

e = water vapour pressure

n_e = electron density

f = frequency

W = liquid water density

- globally distributed
- temperature in stratosphere and upper troposphere, and ...
- humidity on lower troposphere
- high vertical resolution: 0.5 - 1 km
- low horizontal resolution: ~ 200 km
- high accuracy:
 - random errors ~1K
 - systematic errors <0.2K (to be demonstrated in practice)
- “all-weather”
- space/time sampling determined by number of GPS receivers
- relatively inexpensive

Radio occultation missions (1)



Past:

- GPS/MET: 1995 - 1997 experimental, selected periods only

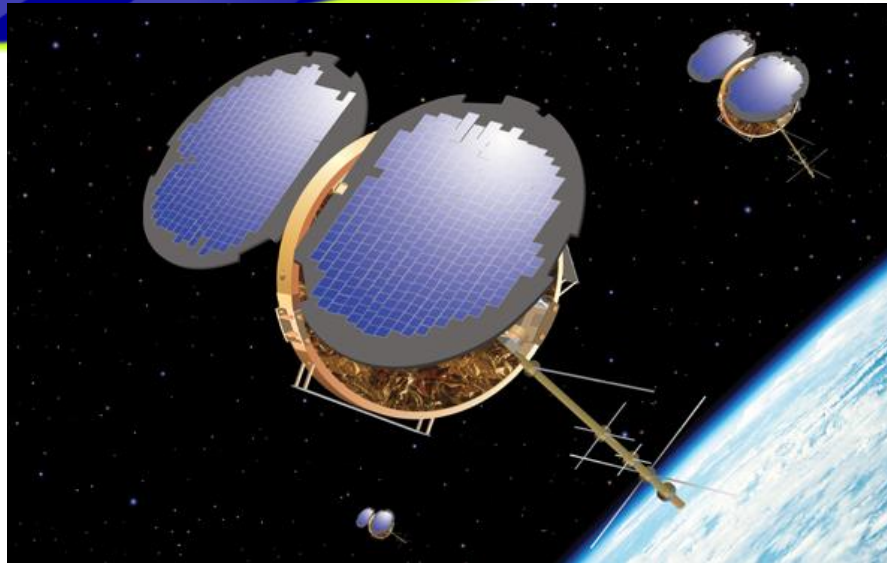
Present:

- CHAMP 2000 - ... exptl, continuous since 2001; NRT since 2006
- SAC-C 2000 - ... sporadic measurements, experimental
- GRACE-A 2002 - ... exptl, continuous since 2003; NRT since 2006
- COSMIC 2006 - ... demonstration mission, **6 satellites**
- MetOp/GRAS 2006 - ... operational from 2007
- TerraSAR-X 2007 - ...

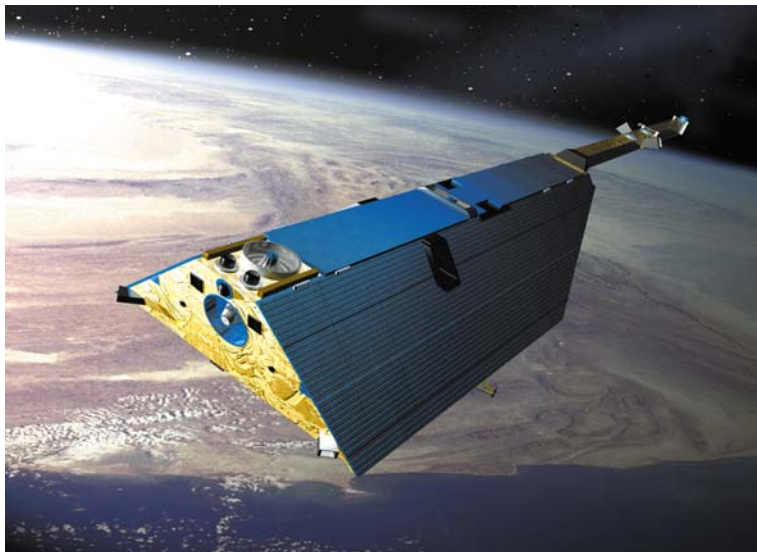
Future:

- EQUARS 2007? emphasising equatorial region
- OCEANSAT-ROSA 2009? Italian / Indian mission
- COSMIC-2 ?
- CICERO ? 20-100 satellites

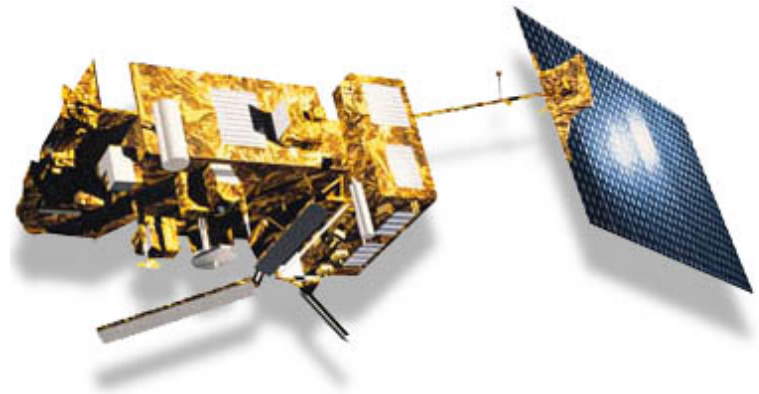
COSMIC

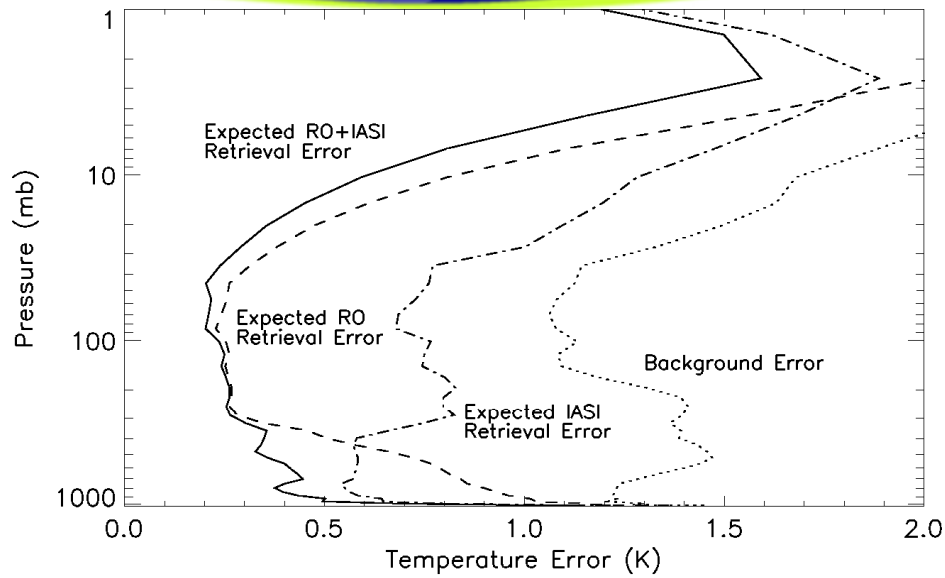


CHAMP

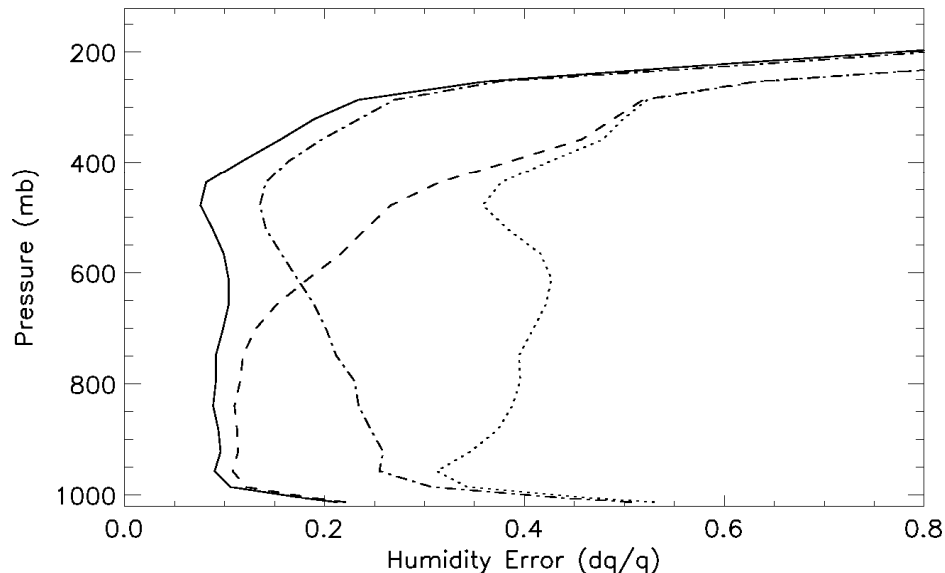


Metop/GRAS





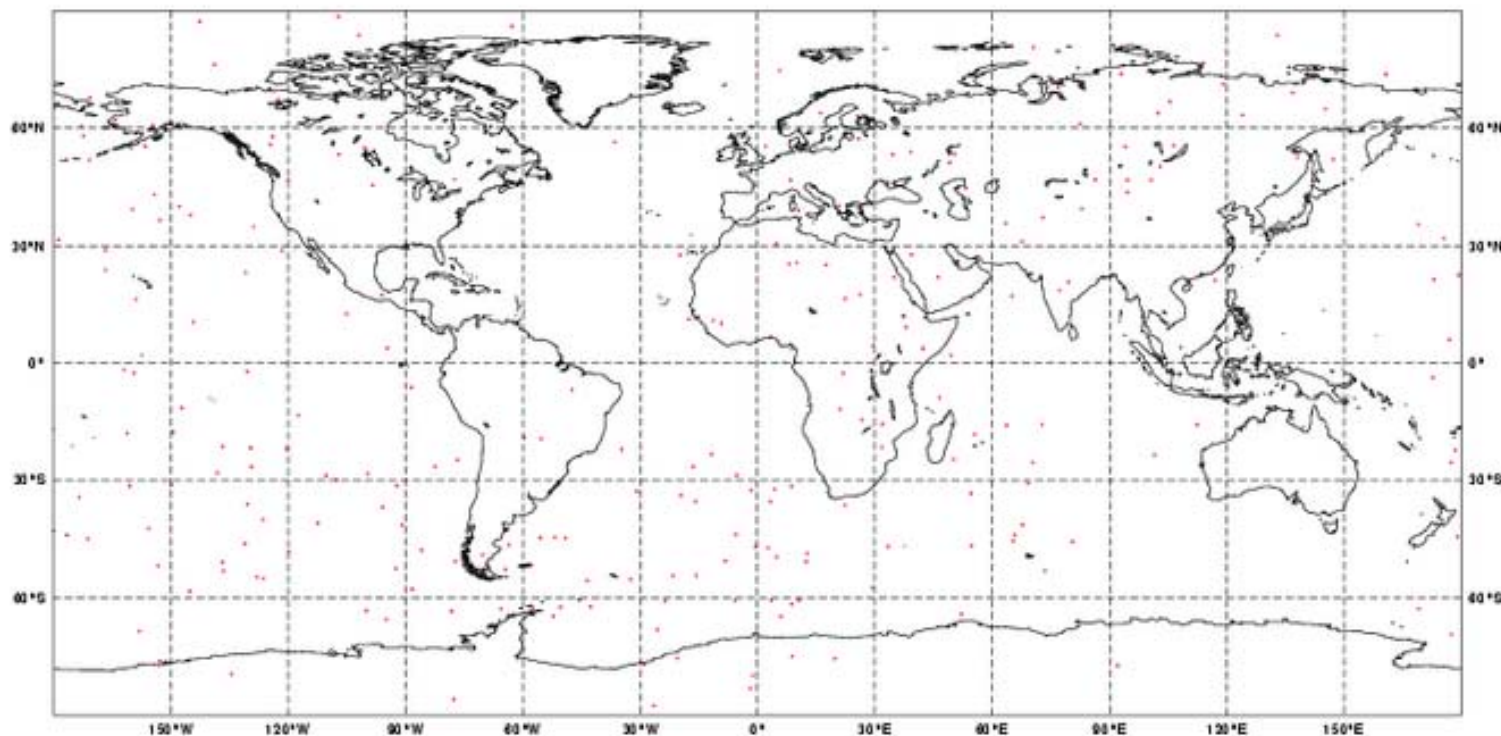
Collard and Healy, 2003



Data Coverage: GPSRO (8/8/2007, 0 UTC, pu00)
Total number of observations assimilated: 246



GPSRO (246)

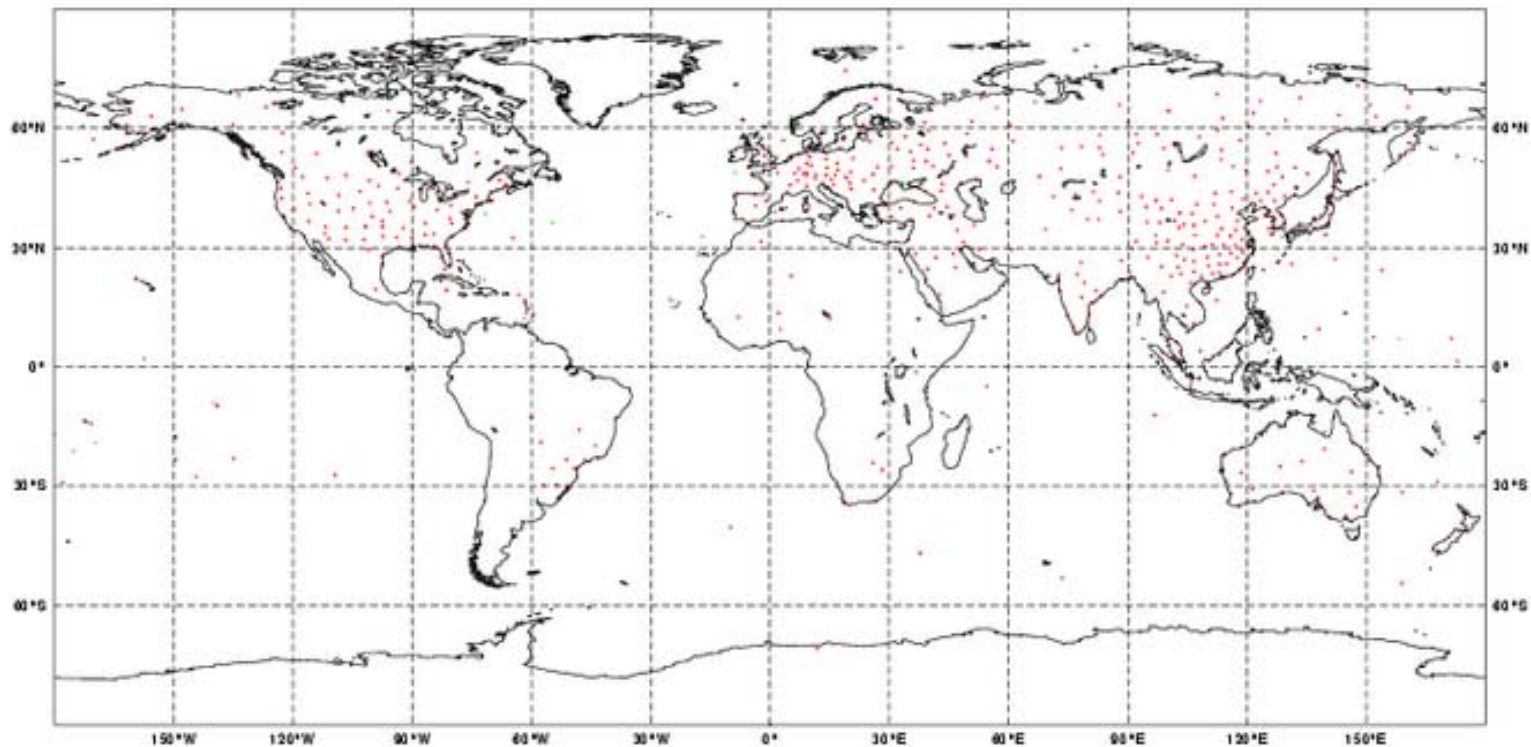


Data Coverage: Sonde (8/8/2007, 0 UTC, pu00)

Total number of observations assimilated: 619



TEMP LAND (613) TEMP SHIP (5) TEMP MOBILE (1)



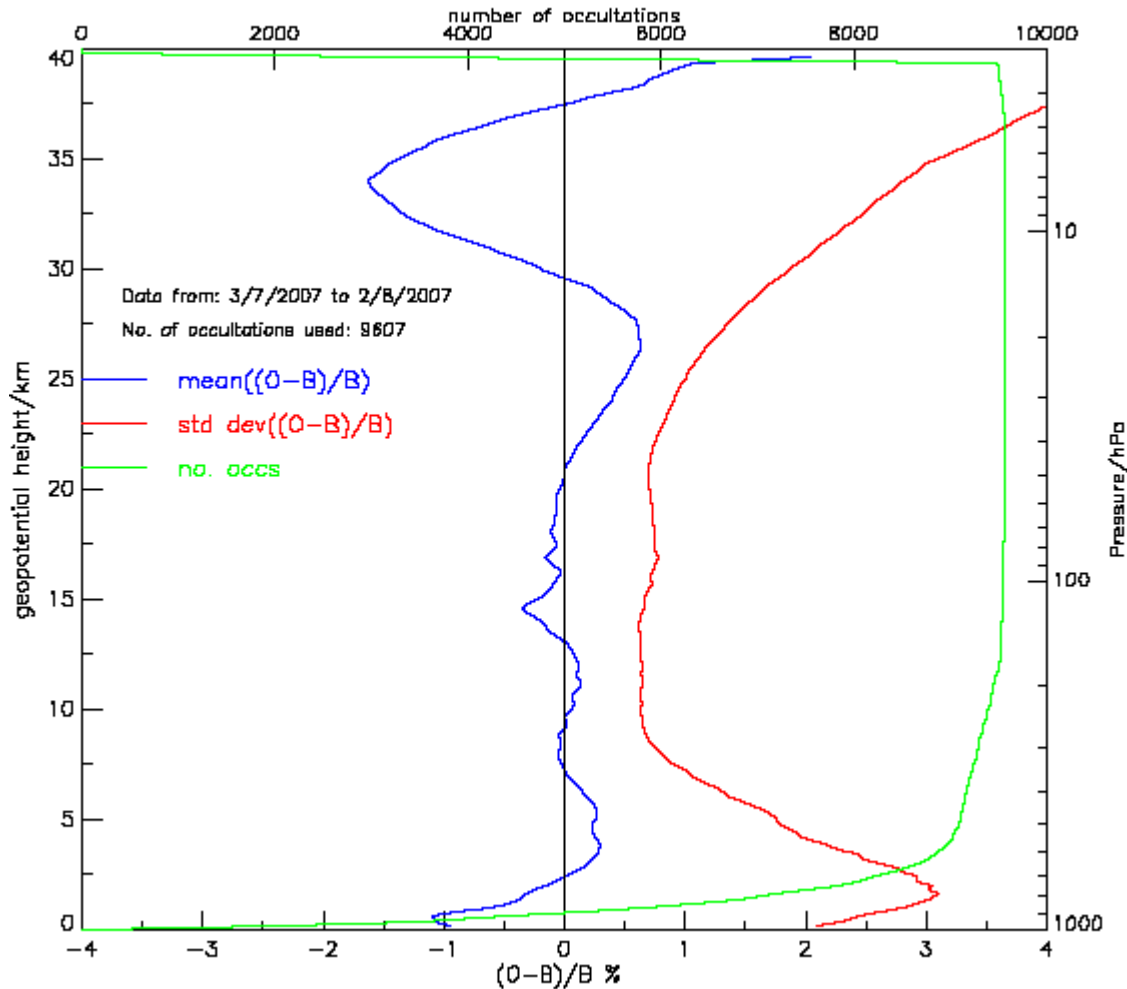
Options:

- (1) assimilate retrieved profiles of temperature and humidity
- (2) assimilate retrieved profile of refractivity, $N(z)$
- (3) assimilate measured refracted angles, $\alpha(a)$, directly

Special problems with RO data:

- non-separability of temperature and humidity
 - addressed by (2) and (3)
- limited horizontal resolution / problems of horizontal gradients
 - partially addressed by (3)

UCAR processed Cosmic 1 data, global
Refractivity bias and standard deviation plots



Plotted at: 13:43 2-Aug-2007

COSMIC-1

3 Jul -2 Aug 2007

Statistics of observation
increments in % refractivity

Statistics are remarkably
stable:

- day to day
- satellite to satellite

More recent advances: radio occultation



Recent results

(M.Rennie, Met Office)

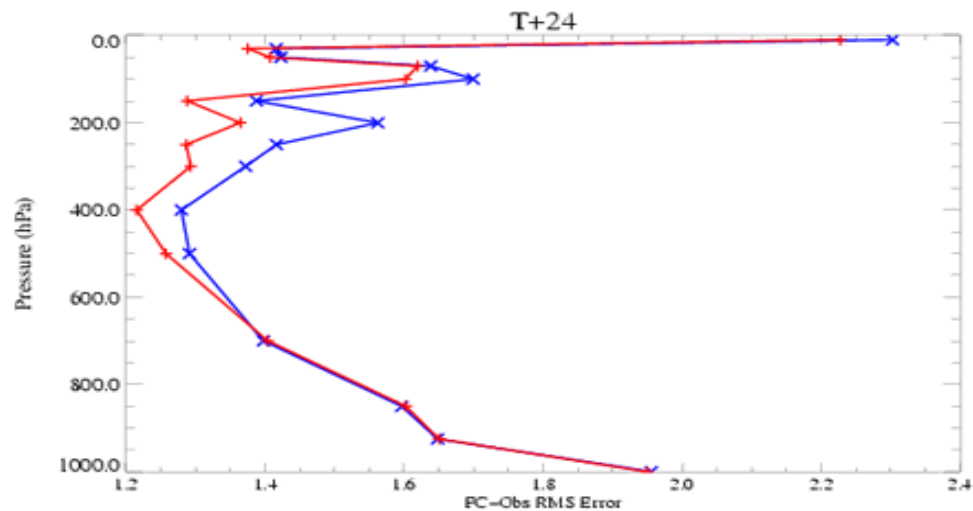
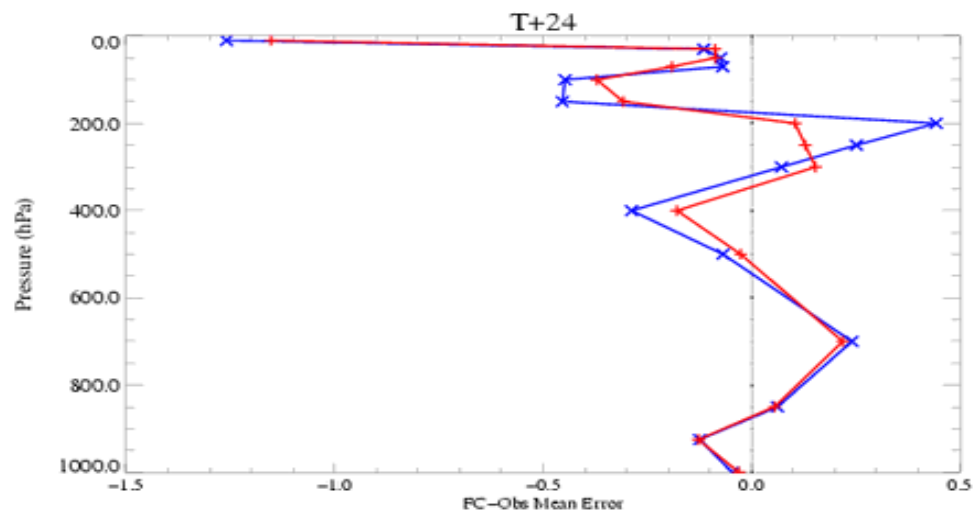
Temperature: mean difference (top)
and RMS difference (bottom) from
sondes, SH, T+24
CONTROL, COSMICx6

The assimilation of GPSRO reduces
RMS errors in the upper troposphere
and corrects model biases.

Similar patterns in NH and TR, but
smaller impact

Temperature (Kelvin): Sonde Obs
Southern Hemisphere (CBS area 20S-90S)
Meaned from 27/11/2006 12Z to 27/12/2006 12Z

Cases: + COSMIC trial for Dec 2006 x Control for Dec 2006



More recent advances: radio occultation



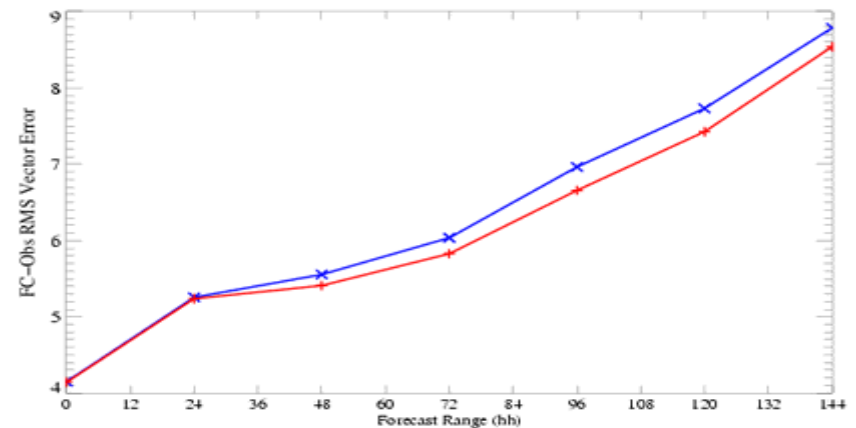
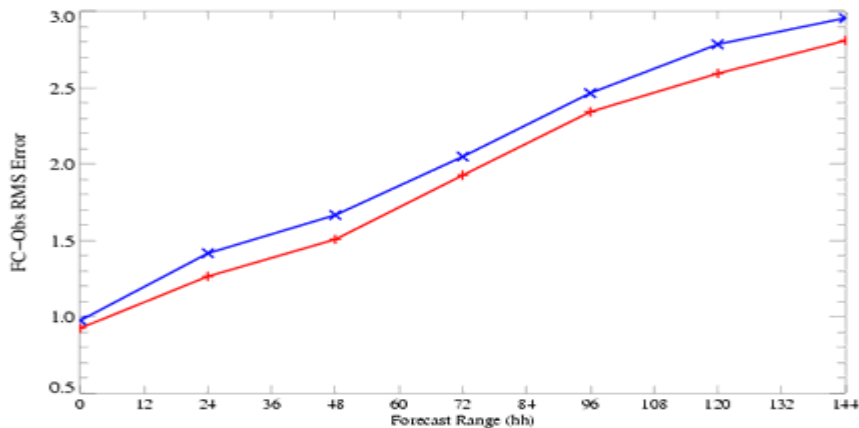
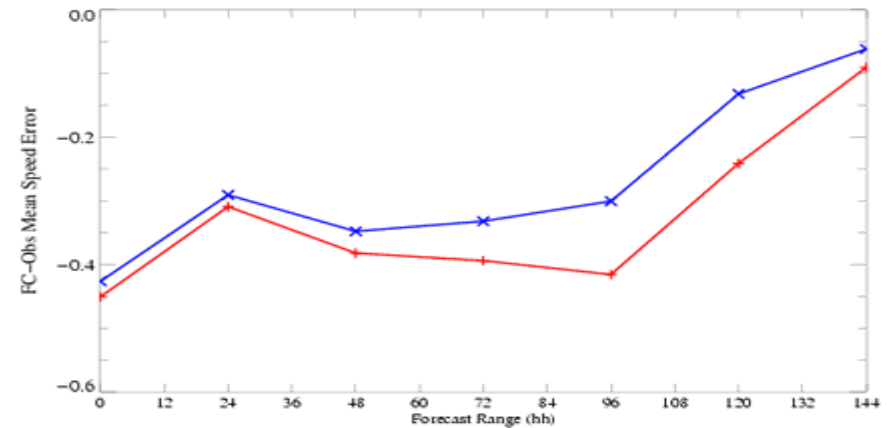
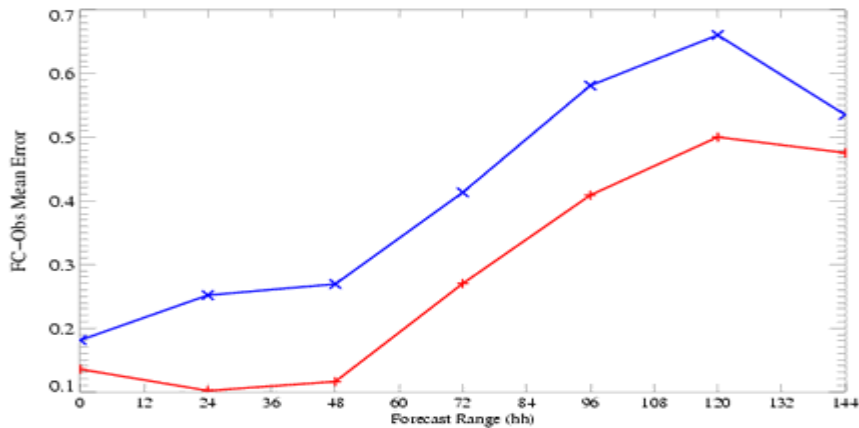
Recent results – bias and RMS v. forecast range

Temperature (Kelvin) at 250.0 hPa: Sonde Obs
Southern Hemisphere (CBS area 20S–90S)
Meaned from 27/11/2006 12Z to 27/12/2006 12Z

Wind (m/s) at 100.0 hPa: Sonde Obs
Southern Hemisphere (CBS area 20S–90S)
Meaned from 27/11/2006 12Z to 27/12/2006 12Z

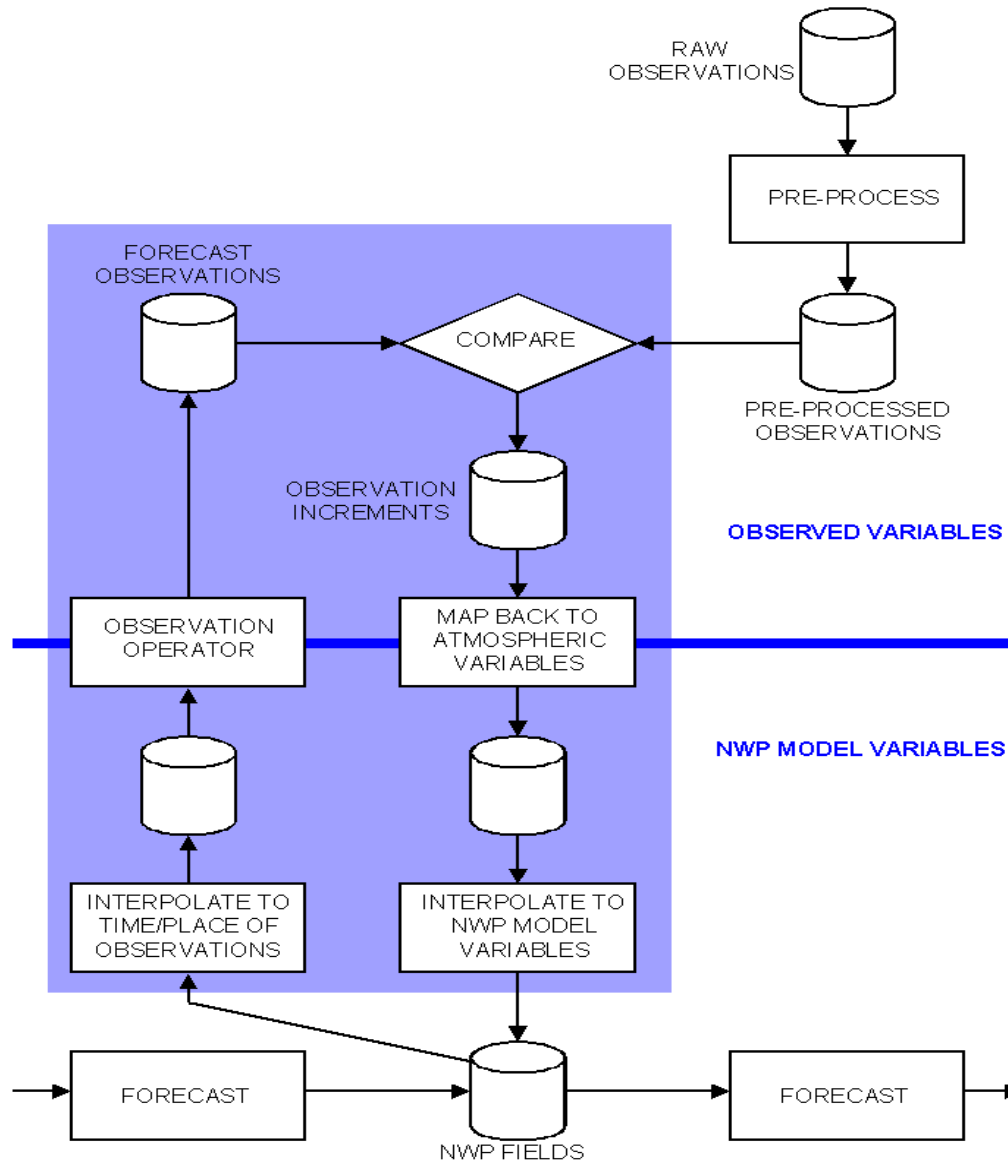
Cases: + COSMIC trial for Dec 2006 x Control for Dec 2006

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Strategies for various data types

Direct assimilation of observations



Assimilation of satellite data: strategies for various data types



Ref.: Eyre, "Variational assimilation of remotely sensed observations of the atmosphere", J Meteorol Soc Japan, 75, 331-338 (1997).

Direct assimilation of "raw" observations

Advantages

- Within variational schemes, the "observation operator", $H(x)$, can be nonlinear - important for many remotely-sensed observations
- In principle, we can use "raw" measurements - in the space of the observed variables - e.g. radiances, backscatter coefficients - **simpler errors**

Limitations

- $H(x)$ must simulate observation **in the form in which it is presented to the system** - $H(x)$ must be matched to any pre-processing
- Raw observations have **more complex operators**
- Some obs are affected by physical variables NOT contained in the control variable
- Logistical problem - need to develop/maintain expertise on all satellite observation operators and associated errors - **STRATEGY NEEDED**: improved links between "assimilation centres" and "satellite centres" → **NWP SAF, JCSDA**

Summary - Needs careful consideration for each obs type

- Passive temperature/humidity soundings
 - as radiances
- Winds
 - small-scale as AMVs, large-scale as radiances?
- Scatterometry
 - as retrieved “ambiguous” wind vectors
 - not backscatter, for subtle reasons - high degree of nonlinearity of obs operator
- MW imagery (water vapour, cloud water, precip, wind speed)
 - complex issues:
 - nonlinearity of multi-variate operators,
 - low vertical resolution (dependence on B-matrix)
- Cloud imagery
 - as retrieved cloud or as radiances?
- Radio occultation
 - as retrieved refractivity or bending angle