The use of airborne and ground based atmospheric observations in carbon cycle research

Christoph Gerbig Max-Planck-Institute for Biogeochemistry



Presented at the ECMWF Seminar on Global Earth-System Monitoring 5-9 September 2005

Overview

- Introduction: Why Carbon Cycle Research
- Global Observations of atmospheric CO₂: from remote islands to places nearby from decades to seconds
 Challenge: continental boundary layer, closeness to strong sources & sinks
- Airborne measurements of tracer distributions
 - Surface fluxes on regional scales (biosphere-atmosphere exchange, ...)
 - Transport processes (tropospheric mixing, convection, ...)
- Hypothesis: Airborne intensives provide paradigm datasets to
 - help design & test (falsify/validate) modeling frameworks
 - help integrate ground based data into data assimilation systems

for the year 2000, relative to 1750



Level of Scientific Understanding

[IDCC Third Assassment Penart Climate Change 2001



Level of Scientific Understanding

[IDCC Third Assassment Penart Climate Change 2001

Fundamental Carbon Cycle Questions

- Where and through which process is the excess anthropogenic carbon being taken up by land and ocean?
- What and how large are the key feedback links between the carbon cycle and the physical climate system?
- What is the carbon budget of a particular region (continent, country)?

First **Scenarios** Calculated with Coupled Carbon Cycle - Climate Models

> Hadley — IPSL —





Estimating Reginal Carbon Balances: Top-Down vs. Bottom-Up Approach **Atmospheric Observing System** Atmospheric Transport **Regional Experiment:** Top-Down Inverse Modelling Multiple Constraint at Very High Resolution European Flux Estimates Past Future Biomass, Bottom-Up Modelling Soil Carbon **Remote Sensing** Inventories **Ecosystem Flux Process Studies** Measurements

Carbon Cycle Observing Systems







GLOBAL NETWORK - 1980



1008E 1108E 1008 1108U 1008U 008U 008U 008E 008E 1008E



GLOBAL NETWORK - Future



10005 11005 1000 11000 10000 0000 0005 0005 10005



Year

Continental boundary layer: diurnal cycle for different heights at a tall tower



July average diurnal cycle (top level 24h mean substracted)



80

90 100

Eddy covariance



acada including annual sums for 2100 ha





- hourly CO₂, HF 30 m
- ? midday CO₂
- 10 day medians
- Bermuda CO₂
- Mauna Loa \overline{CO}_2

Difference Harvard Forest CO₂ – Bermuda CO₂

10-day median of the daily mean CO_2 flux at Harvard Forest (µmole m⁻²s⁻¹)



COBRA-2000 Northern Transect 18-19 August 2000



What models don't need to resolve



"unresolvable" eddies ⇒ "resolvable" mixed laver mean

What models don't need to resolve





distance of pair [km]

Grain size of atmospheric CO₂: Variogram

Variogram for a given "distance bin" (*h*=average distance): $2g(h) = var(CO_2(s_i) - CO_2(s_j))$ with $h = \overline{|s_i - s_j|}$

classical Variogram:

$$2\hat{\boldsymbol{g}}(h) = \frac{1}{N(h)} \sum_{N(h)} \left(CO_2(s_i) - CO_2(s_j) \right)^2$$

with N(h): Number of pairs robust Variogram:

$$2\bar{g}(h) = \frac{\left\{\frac{1}{N(h)}\sum_{N(h)} |CO_2(s_i) - CO_2(s_j)|^{1/2}\right\}^4}{0.457 + 0.494 / N(h)}$$

Fitting of power variogram model



'ariogram: 'ariance of ifferences of airs of mixed ayer profiles neasured ithin 3 hours

Spatial simulation for CO₂



CO₂ [ppm]

Spatial simulation for CO₂











<u>Receptor Oriented Atmospheric Model "ROAM"</u>





































<u>Receptor Oriented Atmospheric Model "ROAM"</u>







GBP Terrestrial Vegetation Map: 17 classes ~10-15 useful eddy flux sites (AmeriFlux) (mostly NE and SE forests)

> NEE = $I_{ir} b_i T + I_{ip} A_i SWR / (G_i + SWR)$ (b, A, Γ)_i from eddy flux data, T and SWR from EDAS λ factors for upscaling, with a priori uncertainty. [Gerbig et al. 2003]

The GSB (Greatly Simplified Biosphere)

- + captures dominant patterns of variability in space (vegetation cover) and time (light sensitive)
- Not very detailed, only diagnostic







<u>Receptor Oriented Atmospheric Model "ROAM"</u>

- Does it work? Is this realistic?
- Compare spatial tracer distribution with observations to validate (or falsify)



Large-scale biospheric CO₂ distribution







Large-scale biospheric CO₂ distribution



Large-scale biospheric CO₂ distribution





Constraints on Convective Fluxes COBRA-2003, June 2003

CO_2 Measurements over the U.S.

CO2 Pease-Terre Haute-Jeffco, 6/27/03-6/28/03



STILT-BRAMS: convection

Stochastic Time Inverted agrangian Transport Model coupled to Brazilian Regional Atmospheric Modeling System

ackward particle motion, receptor at top of COBRA profile 6/28/03, ~18:00

BRAMS_STILT modeled trajectories Arrival time: 2003x06x28x17x39.30Nx089.10Wx08550 hours before arrival: 1 2 2 Altitude [km] ° 8 60 55 50 45 40 N 35 30 0 -120-130410 -90 180 -70100 60 BRAMS starting time: 1/5/2003

MPI BGC Jena

+ fluxes + boundary condition

Forecasting of airmass history: Lagrange Experiment







Concluding Remarks

Future observational network:

- More continental sites that are closer to processes
- Vertical distribution:
 - CMDL: rental aircraft
 - IAGOS: Integration of routine Aircraft measurements into a Global Observing System
 - remote sensing (ground based and satellite based)
- Airborne intensive data can provide
- Constraints on fluxes / terrestrial processes
 - Tight constraint on regional scale: Lagrangian experiments
- "Testbed" for a modeling framework
 - we can only learn from discrepancies models vs. measurements (mixing, convective redistribution)

How we can learn: Interplay between modeling and experiment

- Model => Measurement: utilize the little flexibility we have in the experiments (many constraints by sensors/physics, platforms)
- Measurement => Model (example: grain size): models are more flexible than we often think, need to design models to match measurements (thus they become falsifiable)