

# Upper Ocean Model Physics Development

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Paris, France

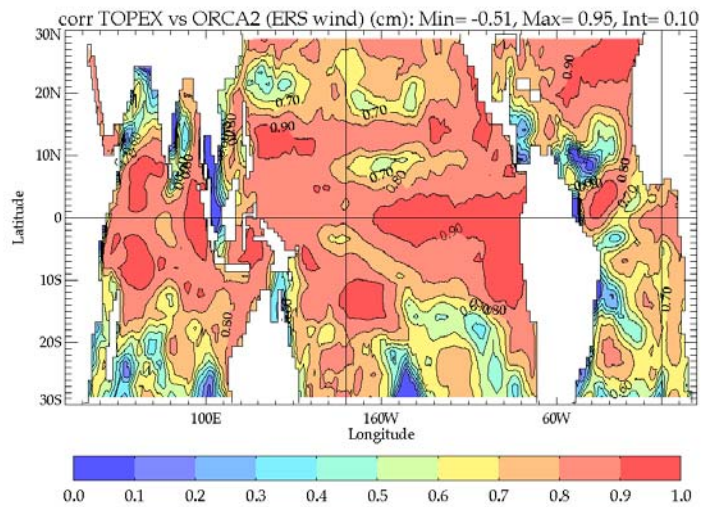
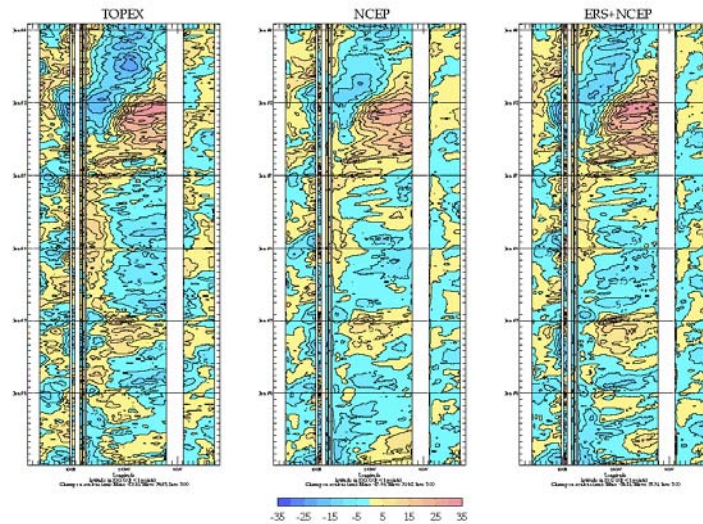
Gurvan.madec@lodyc.jussieu.fr

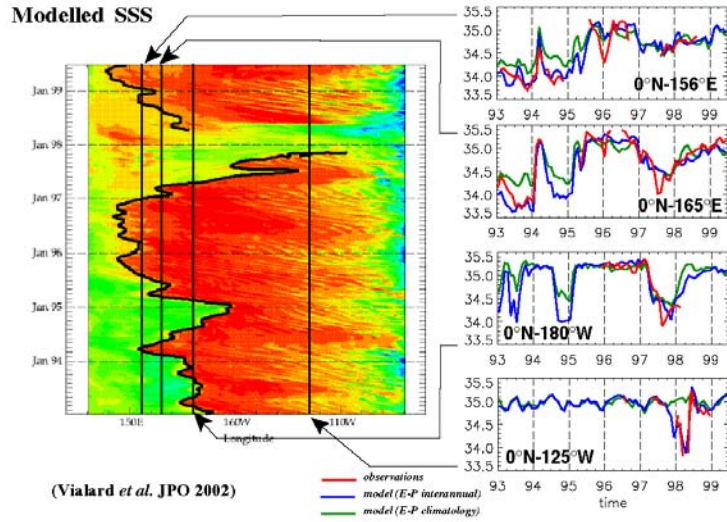
## 1 - Tropical ocean dynamics

- Now-a-day OGCM qualities and why (ORCA2 as a example)
- Incoming Problems with increasing resolution
- Recent development: isopycnal momentum mixing

## 2 - Mid and high latitudes dynamics

- Lessons from HadOPA Coupled simulations  
=> Need of mixed layer improvement
- Deep convection : the Labrador Sea example in CLIPPER projet  
=> Need of efficient way to better mimic the eddy effects even for "eddy resolving" configurations

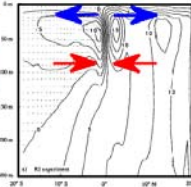




**WHY ? (1/2)**

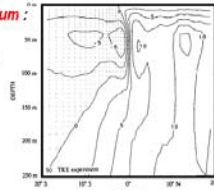
**Equatorial Atlantic MSF**

*Ekman Drift divergence*  
*Geostrophic convergence.*



$Kz = F(Ri)$   
(Pacanowski & Philander JPO81)

**Vertical mixing on momentum:**  
Reduction by factor of 2  
of the equatorial upwelling  
=> reduction of  
the cold tongue pb



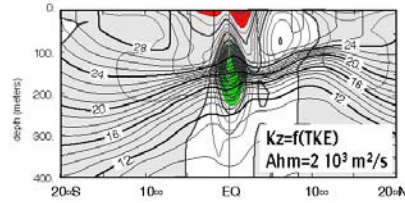
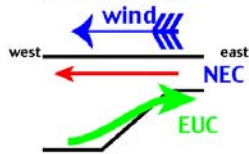
$Kz = F(TKE)$   
(Gaspar et al. JPO90)

FIG. 11. Annual mean (1983) seasonally integrated circulation for (a) the RI experiment and (b) the TKE experiment. The 1000-hPa interval is 2.5-hPa. Dotted negative values correspond to an anticyclonic winter circulation.

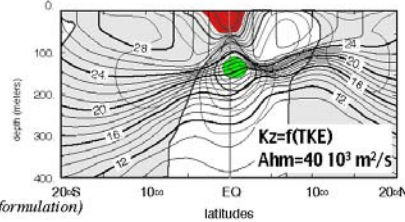
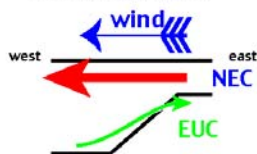
(Blanke & Delecluse JPO93)

**WHY ? (2/2)**

Reference case



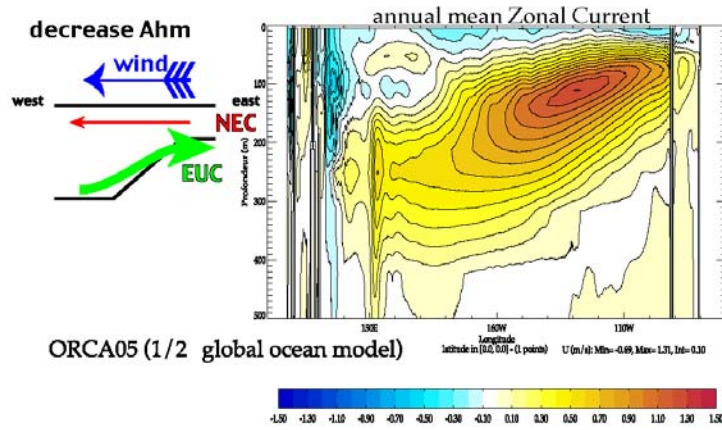
increase of Ahm



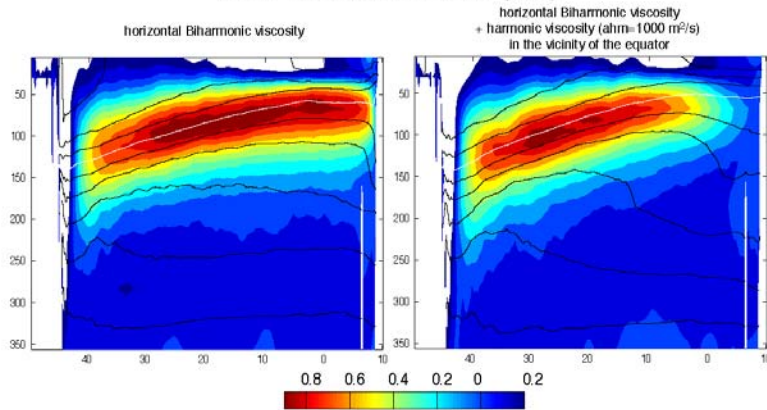
(Delecluse & Madec Elsevier 99)

(see also Large et al. 2001 anisotropic formulation)

Incoming problem with increasing resolution:  
**too low viscosity !**

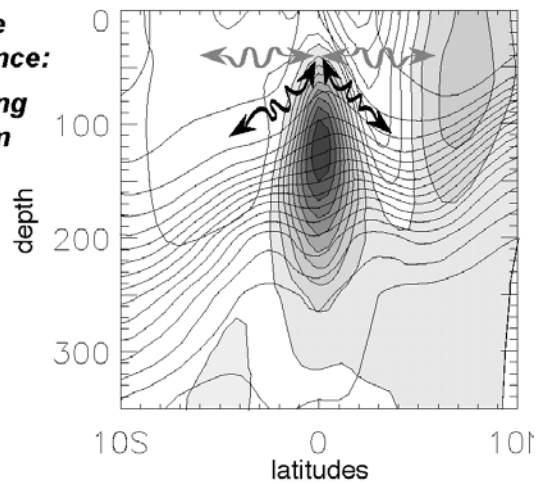


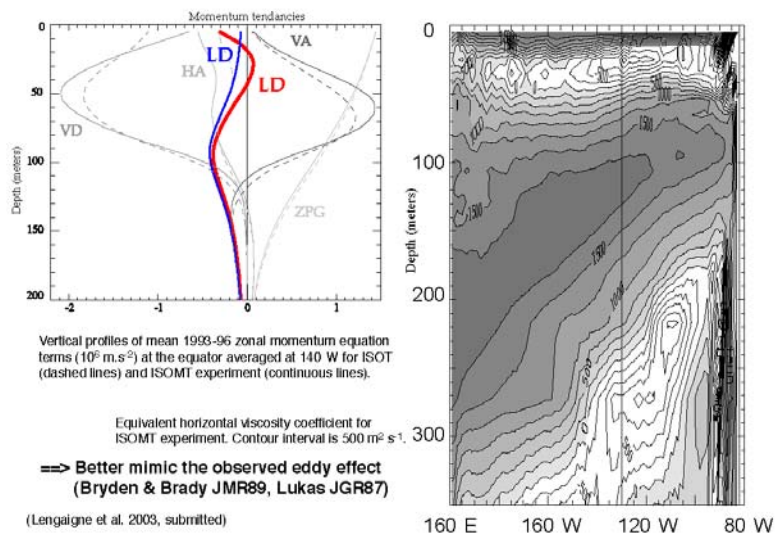
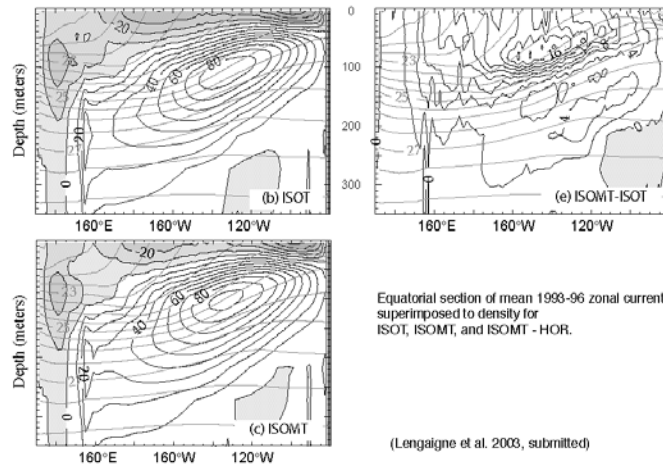
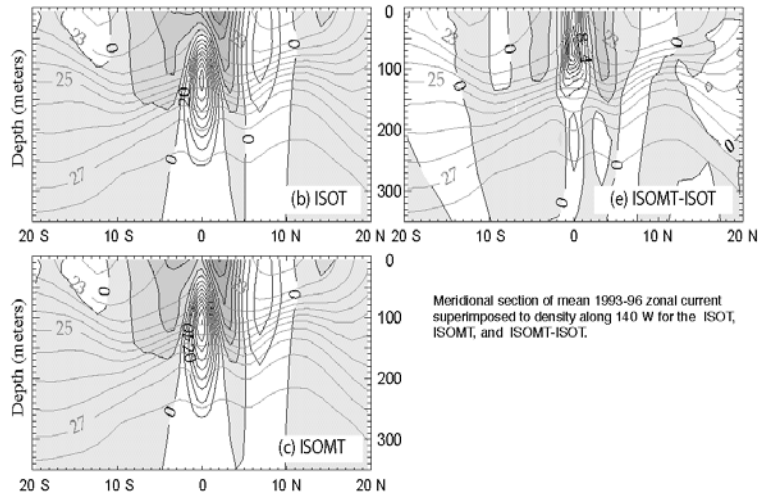
**CLIPPER: Atlantic Ocean (1/6)**

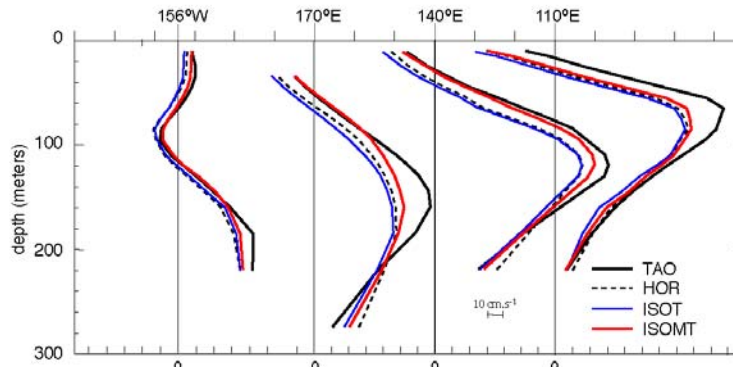


⇒ Need a accurate parameterization of  
 sub-meso-scale equatorial dynamics (K. Richards ...)

**improving the  
 momentum balance:  
 Isopycnal mixing  
 on momentum**



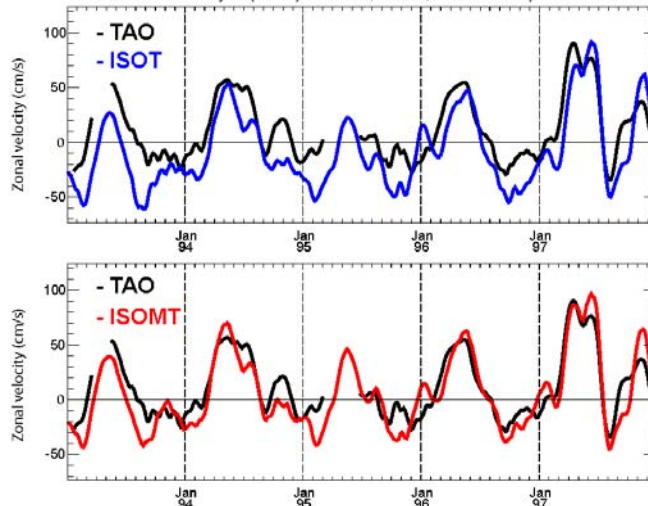




mean 1993-96 zonal velocity profiles for TAO data and the three experiments (HOR, ISOT, and ISOMT) at TAO mooring locations: 156°E, 170°W, 140°W, 110°W.

(Lengaigne et al. 2003, submitted)

Time evolution of zonal currents at 140°W integrated over the mean mixed layer (55 m) for TAO, ISOT, and ISOMT.



## Surface layer Ocean Physics Developments

Gurvan Madec

ECMWF, nov. 2002

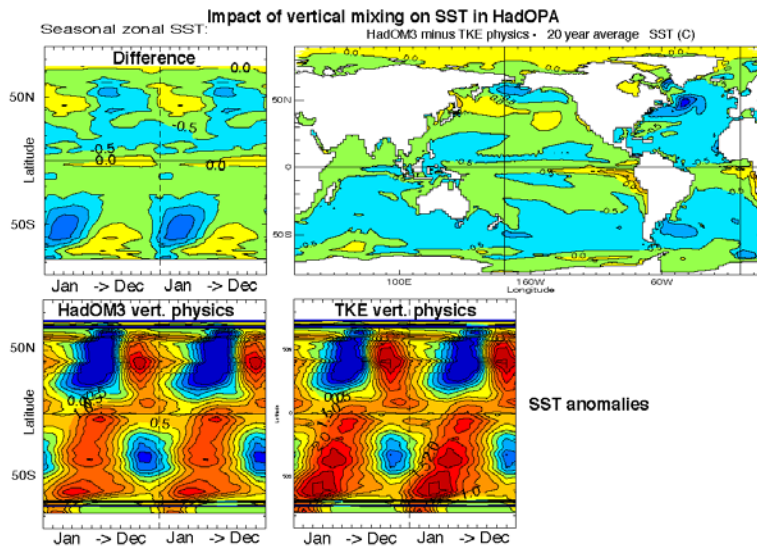
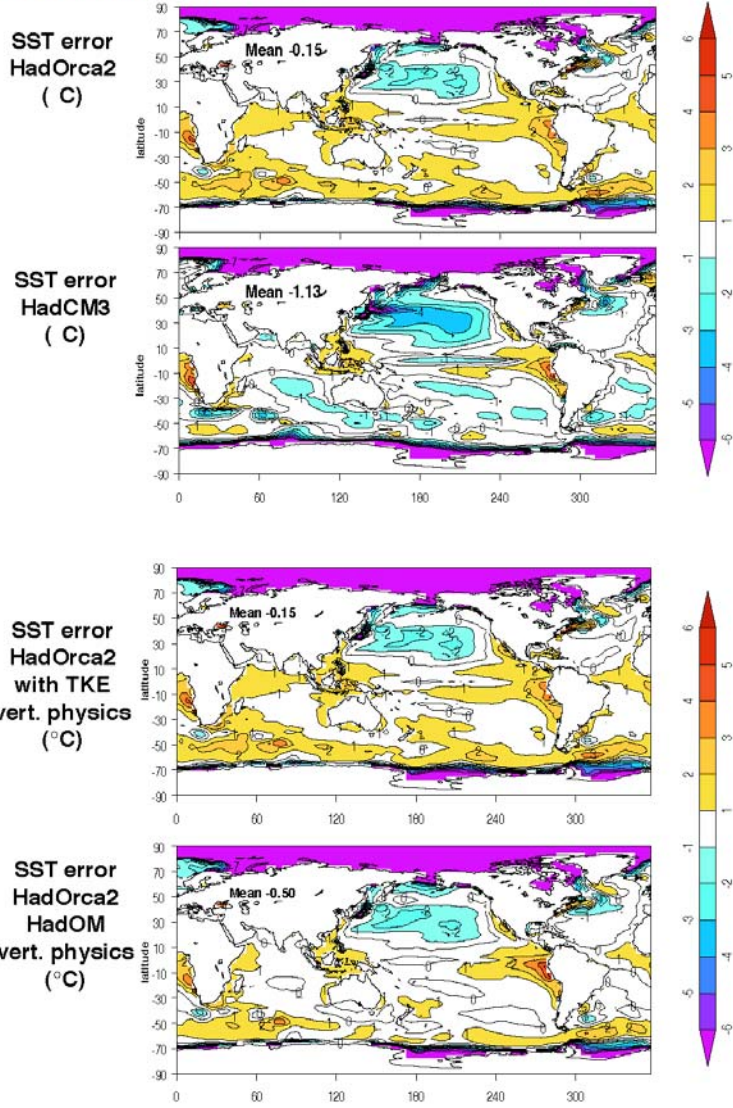
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**Lessons for HadORCA / HadCM3 coupled simulations**



**Problem:** *summer warm biases associated with a too shallow mixed layer*  
*TKE does not represent the effect of wind steering in heating condition*  
*Need to improve the Mixed layer formulation*

**Solution:**  
*set of a minimum mixed layer depth based on Monin-Oboukov length*  
*(as in KPP or Timmermann & Beckmann 2002))*

- but
- bound by the Ekman depth
  - taking into account the ocean vertical motions (upwelling)

*Promising tests in forced mode already performed*



## Surface layer Ocean Physics Developments

**Gurvan Madec**

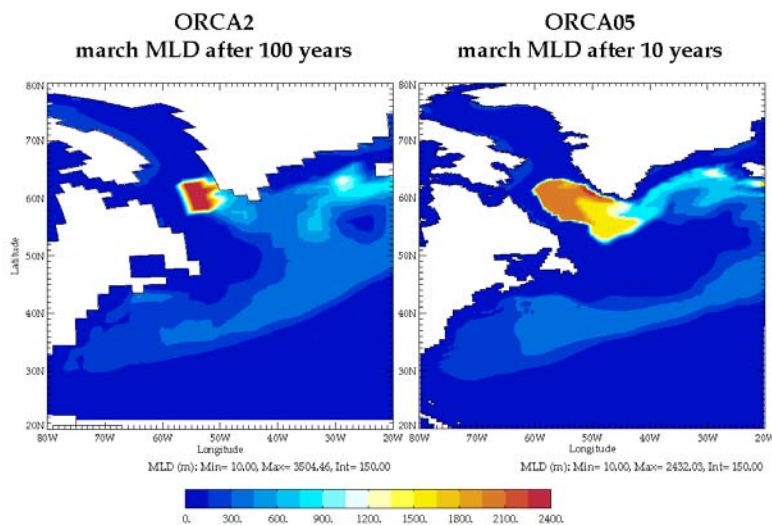
ECMWF, nov. 20002

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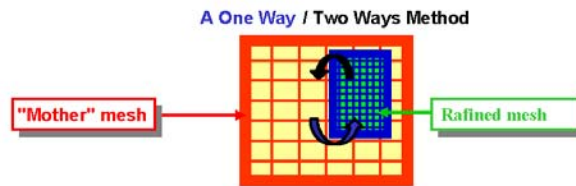
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## AGRIF: Adaptive Grid Refinement In Fortran

L. Debreu, E. Blayo

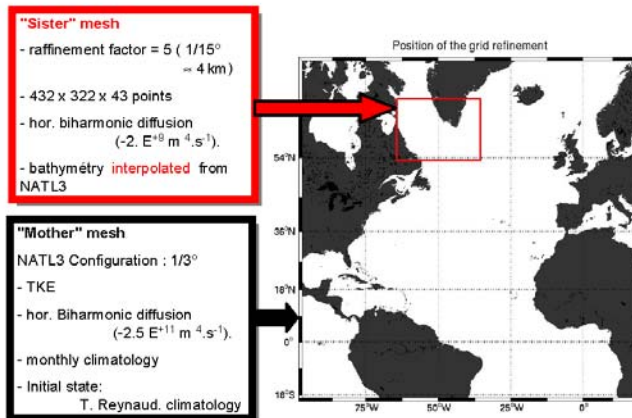


The two models runs simultaneously with information exchanges

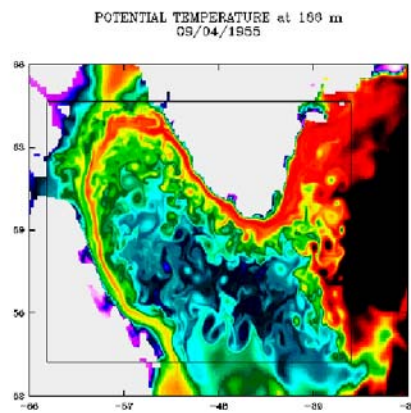
From the mother model to the refined model (One way)

In both directions (Two ways)

### Model Configuration

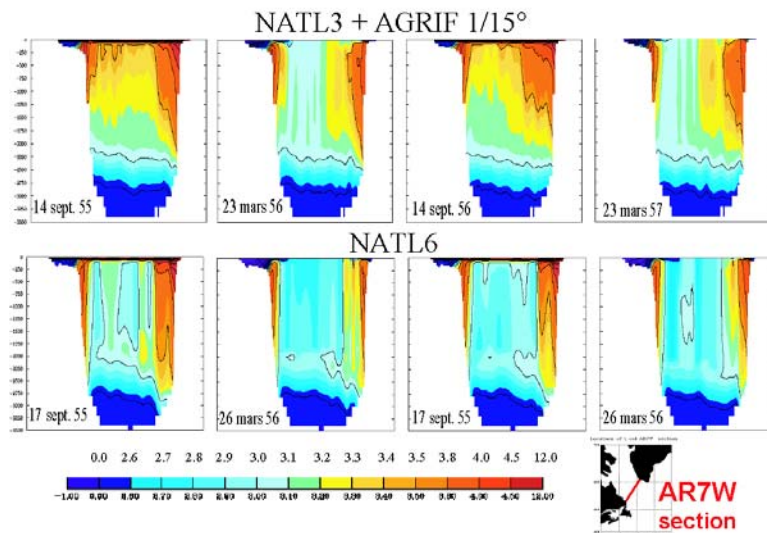
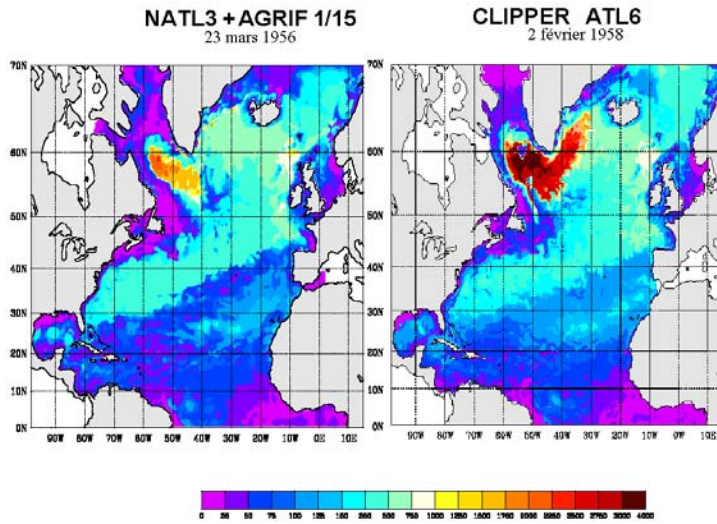


### Time evolution of the Mixed layer

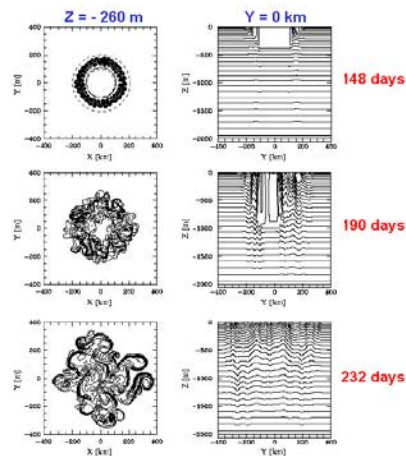


J. Chanut, B. Barrier, J.-M. Molines, A.-M. Tréguier & G. Madec





**Process Study : eddy resolving case**

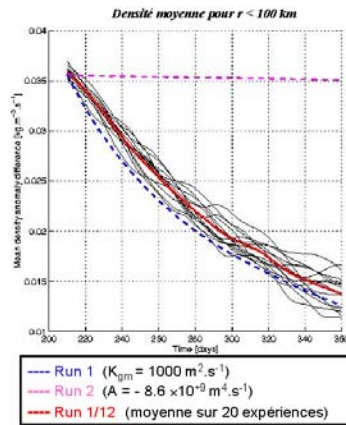


**Process Study : non eddy resolving case**

- Résolution 10 km (1/6 )
- Conditions initiales : moyenne azimutale du profil de densité des 20 expériences au 1/12 à t = 210 jours + vitesses géostrophiques.

- Run 1 : Gent McWilliams
- Run 2 : diffusion biharmonique

**==> Gent McWilliams OK with  $A_{eiv} = 1000 \text{ m}^2/\text{s}$**

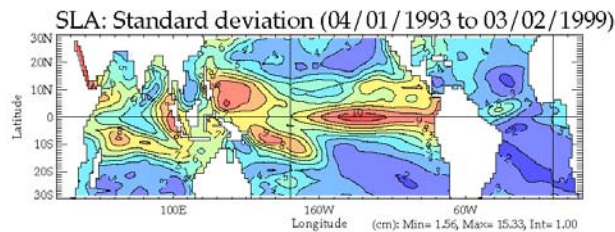


**Conclusion**

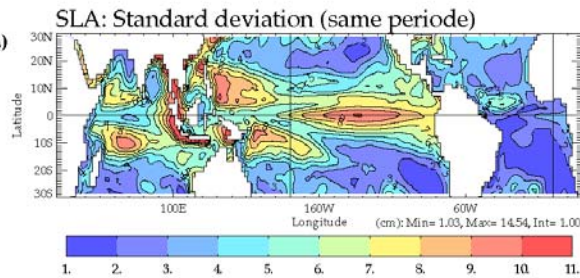


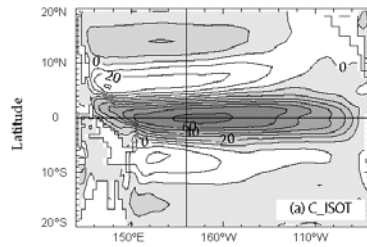
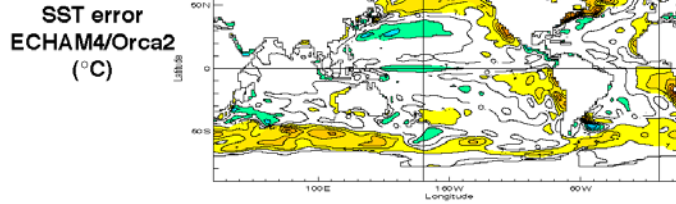
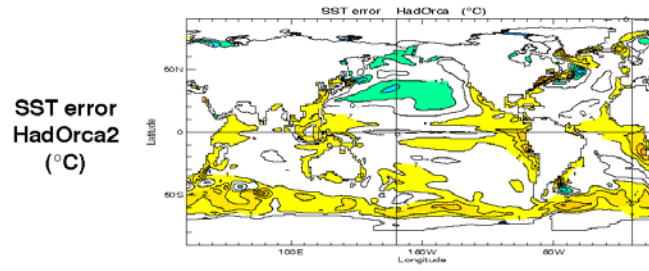
- 1 - Observed wind stress + TAO moorings ==> improved model physics in the tropics
- 2 - With Now-a-day Climate OGCMs resolution :
  - lateral isopycnal physics for both dynamics and tracers (at least at the equator)
  - Eddy induced velocity in extra-tropics and better adjusted for deep convective area
  - vertical physics: TKE + wind steering parameterization (for mid/high latitudes)
- 3 - Foreseen increase in resolution (up to 1/10 ) :
  - Sub-meso-scale physics :  
Need a physically based lateral physics at the equator (==> K. Richards....)
  - still required eddy effect parameterisation :  
Need a accurate parameterisation of restratification phase

TOPEX



ORCA 2 (ERS winds)

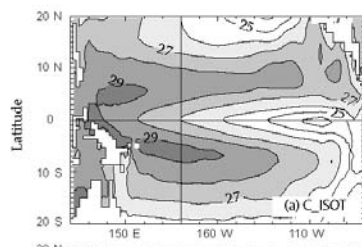
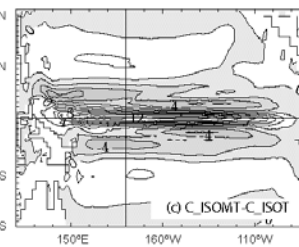
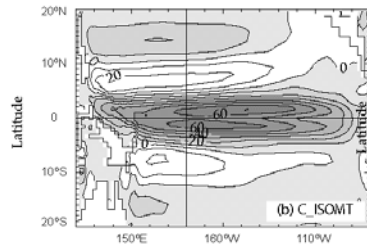




**Coupled experiment ECHAM4/ORCA2**

200-year mean zonal surface currents for (a) C-ISOT, (b) C-ISOMT, and their differences (c) C-ISOMT - C-ISOT for tropical Pacific. Contour interval is 10 cm s<sup>-1</sup> for currents and 2 cm s<sup>-1</sup> for their differences.

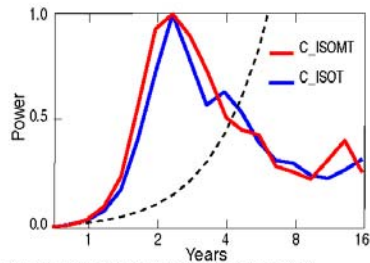
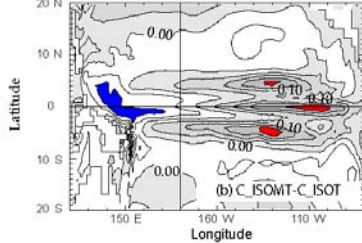
(Lengaigne et al. 2003, submitted)



**Coupled experiment ECHAM4/ORCA2**

200-year mean SSTs for (a) C-ISOMT, (b) C-ISOMT - C-ISOT for tropical Pacific.

(Lengaigne et al. 2003, submitted)



Normalized spectra and standard deviation of Niño 3 SST monthly anomalies for C-ISOT and C-ISOMT experiments.

MADEC, G.: UPPER OCEAN MODEL PHYSICS DEVELOPMENTS

